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DEGREE FOR WHICH THESIS WAS PRESENTED	Master of Science
YEAR THIS DEGREE GRANTED	1981

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A FINITE-ELEMENT METHOD APPLIED TO A
TWO-LEVEL QUASI-GEOSTROPHIC MODEL ATMOSPHERE

by



TIMOTHY OTTO GOOS

A THESIS

SUBMITTED TO THE FACULTY OF GRADUATE STUDIES AND RESEARCH
IN PARTIAL FULFILMENT OF THE REQUIREMENTS FOR THE DEGREE
OF MASTER OF SCIENCE

IN

METEOROLOGY

DEPARTMENT OF GEOGRAPHY

EDMONTON, ALBERTA

SPRING, 1981

THE UNIVERSITY OF ALBERTA

FACULTY OF GRADUATE STUDIES AND RESEARCH

The undersigned certify that they have read, and recommend to the Faculty of Graduate Studies and Research, for acceptance, a thesis entitled "A Finite-Element Method Applied to a Two-level Quasi-geostrophic Model Atmosphere" submitted by Timothy Otto Goos in partial fulfilment of the requirements for the degree of Master of Science in Meteorology.

ABSTRACT

A finite-element method (using the Galerkin formulation) and a truncated spectral method are applied to find numerical solutions to the three non-linear partial differential equations describing a two-level quasi-geostrophic model atmosphere on a β -plane. This work was undertaken to provide a vehicle with which the author could study the finite-element method and its application to problems of meteorological interest. A relatively simple meteorological problem was chosen to allow concentration on the method and its application rather than complex physical interactions.

With the finite-element method, bi-linear basis functions defined on a variable resolution grid are used while sinusoidal basis functions are used with the spectral method. The domain of the problem is a channel of length $2.8 \times 10^7 \text{ m}$ and width $4.4 \times 10^6 \text{ m}$ with a free-slip wall boundary condition applied at the north and south boundaries and periodicity assumed in the x-direction. The grid consists of a central portion with uniformly high resolution and a uniformly changing resolution away from this sub-domain. A second-order Adams-Bashforth time integration scheme is used with both numerical techniques.

Parallel integrations of up to 48 hours duration for a set of four cases are presented and compared, using the spectral solution as a highly accurate standard. In the uniform resolution sub-domain, the finite-element solution achieves a maximum S1 score of under 25 for both

height and thickness fields in three of the cases. Values of lower than 30 are generally considered to be near perfect forecasts in operational weather forecasting. However, it is to be noted that in this study, a highly simplified model is used and integrations are done from initial conditions with no inherent error. In the fourth case, numerical instability occurred due to the rapid growth of spurious short-wave-length waves generated near the boundaries. Investigations revealed that these waves were being produced by the inaccurate evaluation of normal derivatives near and on the boundaries. A possible method for overcoming this problem is discussed.

ACKNOWLEDGEMENTS

I am grateful to the people and organizations who helped me with this study.

I am indebted to my departmental supervisor, Dr. E. P. Lozowski, for his guidance, suggestions, and his thorough review of this manuscript. In particular, I wish to thank him for the confidence he placed in me in allowing me to study this particular area.

I wish to thank Dr. K. D. Hage and Dr. A. Craggs who, with Dr. Lozowski, served on my examining committee.

I wish to thank Dr. A. Staniforth of RPN in Dorval, P.Q. for supplying me copies of computer programs which have been developed in that office.

Thank you to the Department of Geography for providing me with the necessary computer resources to do this study.

This study was undertaken while I was on educational leave from the Atmospheric Environment Service of the Department of the Environment.

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LIST OF SYMBOLS

SYMBOL		PAGE
a	mean radius of the earth	6
a_i	eigenvalue of the i 'th expansion function	23
APE	available potential energy	5
b_{li}	interaction coefficient	25
c_{ljk}	interaction coefficient	25
c_p	specific heat for dry air at constant pressure	11
C	speed	60
D	domain	15
e_G	error in forecast gradient	63
$e^i(x,y)$	i 'th two-dimensional basis function	38
$e^m(x)$	m 'th one-dimensional basis function	41
f	Coriolis parameter	7
f_0	Coriolis parameter at latitude ϕ_0	8
$F_i(x,y)$	i 'th spectral expansion function	20
g	gravitational constant	17
$I(\phi)$	variational integral	42

SYMBOL		PAGE
$I^*(\phi)$	approximation to $I(\phi)$	43
J	Jacobian function	12
KE	kinetic energy	5
l_y	y-direction cosine	44
L_x	length of domain in x-direction	7
L_y	length of domain in y-direction	7
M	total mass of model atmosphere	16
N	total number of spectral expansion functions	20
NI	number of grid points in x-direction	40
NJ	number of grid points in y-direction	40
p	pressure	6
PE	potential enstrophy	16
q_i	potential vorticity at level i	18
Q_i	vorticity at level i	11
Q_g	geostrophic vorticity	9
\overline{Q}	vorticity of mean height field	13
\hat{Q}	vorticity of thickness field	13

SYMBOL		PAGE
R	gas constant for dry air	11
S_1	verification score	62
$S(x_i)$	smoothing parameter at point x_i	64
t	time	9
T	temperature	11
τ	streamfunction of the thickness field	109
u_x	x-derivative of $u(x,y)$	46
u_g	x-component of geostrophic wind	9
v_g	y-component of geostrophic wind	9
\vec{V}_g	geostrophic velocity	9
x	east-west coordinate in horizontal plane	6
y	north-south coordinate in horizontal plane	6
y_0	y-coordinate at ϕ_0	6
z	vertical coordinate	17
α	specific volume	11
β	variation of Coriolis parameter with latitude	3

SYMBOL		PAGE
Γ	boundary off domain	17
δ_{ij}	Kronecker delta	22
Δp	pressure difference between levels 1 and 3	6
θ	potential temperature	11
ρ	density	17
σ	static stability parameter	10
τ_i	i'th spectral expansion coefficient for the thickness field	23
ϕ	latitude	6
ϕ_0	latitude where tangent plane touches the earth	6
$\phi_{B,1}$	value of the function ϕ on the boundary	55
ψ_i	(1) stream function at level i (2) i'th spectral expansion coefficient for the height field	6 23
$\overline{\psi}$	stream function for the mean height field	12
$\hat{\psi}$	stream function for the thickness field	12
ψ_Z	zonal stream function	16
ψ_E	eddy stream function	16
ω	over-relaxation factor	59

SYMBOL		PAGE
$\tilde{\omega}$	vertical wind component in isobaric coordinates	6
ω_i	i'th spectral expansion coefficient for vertical wind component	23
$\tilde{\omega}_i$	vertical wind component at level i	10
Ω	angular speed of rotation of the earth	7
XCI	streamfunction of the mean height field	109

CHAPTER 1

INTRODUCTION

An accurate method for producing a weather forecast has been an elusive dream throughout much of man's history. The abundance of folk beliefs (Hornstein, 1978) available for predicting the weather indicates the importance man has attached to weather forecasts. During this century, much attention has been focussed on using numerical methods to solve the hydro-dynamical equations, thereby producing a weather forecast (Bjerknes, 1904; Richardson, 1921; Charney et. al., 1950; Cressman, 1958, 1963; Shuman and Hovermale, 1968; Bourke, 1972, 1974; Machenhauer and Daley, 1972).

In recent years (Wang et. al., 1972; Cullen, 1973, 1974a, 1974b, 1976, 1979; Staniforth and Mitchell, 1977, 1978; Staniforth and Daley, 1977, 1978, 1979), the finite-element method (FEM) has been applied to the problem of numerical weather prediction. The development of larger and faster computers has made possible the use of the FEM for the prediction of large-scale atmospheric flows. The FEM enables one to use a grid with variable spacing between grid points with relative ease. This is a distinct advantage over the finite-difference methods which have been used previously.

Numerical weather prediction began when Bjerknes (1904) recognized that the primitive equations of meteorology formed a system of non-linear partial differential equations which could, in principle,

be solved to forecast the subsequent states of the atmosphere from a known initial state. He also recognized that the system did not have an analytic solution and that the available data were inadequate to specify the initial conditions. Holton (1972) shows that the primitive equations consist of the momentum equations in the horizontal (plane), the thermodynamic energy equation, the continuity equation, the hydrostatic approximation, and the equation of state. This system of equations forms the basis of large-scale numerical weather prediction. Richardson (1921) attempted the first numerical integration of these equations using a finite-difference method. Unhappily, his forecast took many months to produce and was in error by several orders of magnitude.

The development of the electronic computer in the 1940's allowed Charney et. al. (1950) to perform the first successful numerical weather forecast. They integrated, in time, the barotropic vorticity equation which is a simplified equation derivable from the primitive equations. By simplifying the primitive equations using certain assumptions about the atmosphere, one may obtain a system of equations which is more easily solved. Holton (1972) describes the assumptions necessary to eliminate sound, gravity and inertia waves. Charney et. al. (1950) used the most simplified form of the primitive equations for their integration.

The field of numerical weather prediction has expanded rapidly since the first successful integration, with many more complicated finite difference models being developed (Cressman, 1958, 1963;

Shuman and Hovermale, 1968; Howcroft, 1971). Beginning in the late 1960's, the use of the spectral method for integrating the primitive equations was investigated (Robert, 1966, 1970; Elsaesser, 1966; Eliassen and Machenhauer, 1970; Baer and Alyea, 1971; Bourke, 1972). Efficient spectral models were not possible until Orszag (1970) described the use of transforms to allow the efficient evaluation of the product terms in the primitive equations. Since then, the spectral method has been used with success in numerical weather prediction (Daley et. al., 1976; Bourke, 1974).

In the 1970's, the FEM was first introduced to numerical weather prediction. Early work (Wang et. al., 1972; Cullen, 1973, 1974a), considered highly simplified atmospheric models. Recently, models using the primitive equations defined over the northern hemisphere have been integrated using the FEM (Staniforth and Daley, 1978, 1979; Cullen, 1979). Staniforth and Daley (1979) have found the finite-element model produces forecasts as accurate as those of the operational spectral model even though the finite-element model does not contain many of the physical processes present in the spectral model.

In this thesis, a FEM is used to solve a set of partial differential equations. These equations mathematically describe a two-level baroclinic model of the atmosphere on a β -plane. (The model is fully described in Chapter 2.) The equations of this model atmosphere are also solved using a spectral method. This solution is highly accurate and is used to evaluate the accuracy of the solution by the

FEM. In Chapter 5, the solutions using both methods are presented and compared. Although the meteorological problem under consideration is relatively simple, this approach provides a good framework within which my supervisor and I could learn about and apply the FEM.

CHAPTER 2

A DESCRIPTION OF THE MODEL ATMOSPHERE

2.1 The Model Equations and the Domain

The intent of this thesis is to study the application of the FEM to the prediction of synoptic scale atmospheric flows. Consequently, a model atmosphere has been chosen in which such flows are possible. The quasi-geostrophic two-level model was selected as it is the simplest synoptic scale model which includes baroclinic effects. This model has been used in the past to study atmospheric flows (e.g. Phillips (1951), Holton (1972), Stone (1974), Held (1975)) as although it is a relatively simple atmospheric model, it simulates many atmospheric processes well.

A model atmosphere which allows baroclinic effects has been employed because the baroclinic conversion of potential to kinetic energy is a major process in the development of synoptic scale storms. The total potential energy of a system is the sum of its internal energy and its gravitational potential energy. Lorenz (1955) defines the available potential energy (APE) of a column of the atmosphere to be the difference between the total potential energy and the minimum total potential energy which could be achieved by an adiabatic redistribution of the mass of that air column. The APE is that portion of the total potential energy which is available for possible conversion to kinetic energy (KE). This conversion is often an unstable process, i.e. baroclinic instability.

In the quasi-geostrophic two-level model, the atmosphere is divided into levels as shown in Figure 2.1. Pressure (p) is used as the vertical coordinate. The streamfunction (ψ) is defined at levels 1, 2 and 3 and the vertical motion ($\tilde{\omega}$) at levels 0, 2 and 4 is considered. Levels 1 and 3 are the two levels of prime importance, from which the model derives its name. The pressure difference between levels 1 and 3, Δp , is 500 mb for this model. Pedlosky (1979) has shown that this model, with one small additional assumption, has the same dynamic equations as a two layer model, in which the layer between levels 0 and 2 is assumed to have a certain uniform density and the layer between levels 2 and 4 is assumed to have another uniform density.

Prior to presenting the model equations, a few definitions will be given. When the β -plane approximation is used, the earth is approximated by a plane tangent to the earth at some latitude ϕ_0 , as shown in Figure 2.2. On this plane, a three-dimensional right-handed Cartesian coordinate system is defined. The unit vector \hat{i} points in the x-direction, which is east, while the unit vector \hat{j} points in the y-direction, which is north. The unit vector \hat{k} points vertically in the direction of lower atmospheric pressure. The origin of the coordinate system is defined to be at the point $x = 0$, $y = 0$, $p = 1000$ mb. A relation between latitude, ϕ , and distance along the y-axis may be written:

$$y - y_0 = a(\phi - \phi_0) \quad (2.1)$$

where a is the mean radius of the earth and y_0 is the value of y at ϕ_0 .

The plane in Figure 2.2 is assumed to have a width of $L_y/2$ so the range of y is $[0, L_y/2]$. This definition for the width of the plane is made for consistency with the expansion functions used with the spectral method discussed in Chapter 3. The plane is assumed to have a length L_x . More precisely, it is assumed that the solution to the model equations, derived later in this chapter, is periodic in x with wavelength L_x . This length of the domain is also chosen for consistency with the expansion functions.

The particular values of L_x and L_y used are chosen by considering the type of atmospheric motions to be studied. As mid-latitude synoptic scale motions will be considered in this thesis, ϕ_0 is chosen to be 45°N latitude. The domain is chosen so that $L_y/2 = 4.4 \times 10^6\text{m}$, which is approximately 55° of latitude on the earth. This width is chosen by considering the meridional extent of typical mid-latitude long waves on the earth. It is similar to that used by others (Cullen, 1976; Grammelvedt, 1969). The chosen length of periodicity is $L_x = 2.8 \times 10^7\text{m}$ which is essentially the length of the latitude circle at 45°N .

The Coriolis parameter, f , is defined by:

$$f = 2\Omega \sin\phi \quad (2.2)$$

where Ω is the earth's angular velocity. If only the first two terms of a Taylor expansion of f about y_0 are retained, one obtains:

$$f = f_0 + \beta(y-y_0) \quad (2.3)$$

where:

$$\beta = \left. \frac{df}{dy} \right|_{\phi=\phi_0} = \frac{2\Omega}{a} \cos\phi_0 \quad (2.4)$$

and f_0 is the Coriolis parameter at ϕ_0 .

The horizontal del operator is defined by:

$$\vec{\nabla} = \hat{i} \frac{\partial}{\partial x} + \hat{j} \frac{\partial}{\partial y} \quad (2.5)$$

while the horizontal Laplacian operator is defined by:

$$\nabla^2 = \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2} \quad (2.6)$$

Jeffreys (1922) defines the geostrophic velocity to be the velocity field resulting from a balance between the Coriolis force and the pressure gradient force. When pressure is used as the vertical coordinate, the pressure gradient force term is transformed into a geopotential height gradient term. If the flow is quasi-non-divergent, the geopotential height gradient term may be replaced by a stream function gradient term.

The horizontal geostrophic velocity is quasi-non-divergent and, therefore, it may be written:

$$\vec{V}_g = \hat{k} \times \vec{\nabla} \psi_g \quad (2.7)$$

where \vec{V}_g is the horizontal geostrophic velocity and ψ_g is the geostrophic stream function. In component form, this is:

$$u_g = - \frac{\partial \psi_g}{\partial y} \quad (2.8)$$

$$v_g = \frac{\partial \psi_g}{\partial x} \quad (2.9)$$

where $\vec{V}_g = \hat{i} u_g + \hat{j} v_g$.

The vertical velocity in pressure coordinates, $\tilde{\omega}$, is defined by:

$$\tilde{\omega} = \frac{dp}{dt} \quad (2.10)$$

where t is time. The geostrophic vorticity, Q_g , is defined by:

$$Q_g = \hat{k} \cdot \vec{\nabla} \times \vec{V}_g = \nabla^2 \psi_g \quad (2.11)$$

The quasi-geostrophic vorticity equation may be written as (Holton, 1972; Hoskins, 1975):

$$\frac{\partial Q_g}{\partial t} = - \vec{V}_g \cdot \vec{\nabla} \{Q_g + f\} + f_o \frac{\partial \tilde{\omega}}{\partial p} \quad (2.12)$$

It is paradoxical that although the geostrophic wind is non-divergent, the second term in (2.12) is a divergence term. Holton (1972) has shown that this term is important in keeping temperature changes hydrostatic and vorticity changes geostrophic in synoptic scale systems. The temperature and vorticity fields must remain consistent with the original assumptions of the quasi-geostrophic vorticity equation. In the derivation of this equation, it is assumed that the atmosphere is in hydrostatic balance and that the flow is quasi-geostrophic.

Eqn. (2.12) is applied at levels 1 and 3 using the following finite-difference approximations for the vertical gradient of $\tilde{\omega}$:

$$\left(\frac{\partial \tilde{\omega}}{\partial p} \right)_1 \approx \frac{\tilde{\omega}_2 - \tilde{\omega}_0}{\Delta p} \quad (2.13)$$

$$\left(\frac{\partial \tilde{\omega}}{\partial p} \right)_3 \approx \frac{\tilde{\omega}_4 - \tilde{\omega}_2}{\Delta p} \quad (2.14)$$

The vertical velocities at the top, $\tilde{\omega}_0$, and bottom, $\tilde{\omega}_4$, of the atmosphere are assumed to be zero. This assumption eliminates external gravity waves from the model.

Holton (1972) shows that the adiabatic quasi-geostrophic thermodynamic energy equation may be written:

$$\frac{\partial}{\partial t} \left(- \frac{\partial \psi}{\partial p} \right) = - \vec{V}_g \cdot \vec{\nabla} \left(- \frac{\partial \psi}{\partial p} \right) + \frac{\sigma}{f_0} \tilde{\omega} \quad (2.15)$$

where the static stability parameter is defined by:

$$\sigma = - \frac{\alpha}{\theta} \frac{\partial \theta}{\partial p} \quad (2.16)$$

and the potential temperature is defined by:

$$\theta = T \left(\frac{1000}{p} \right)^{R/c_p} \quad (2.17)$$

In these equations, α is the specific volume and T is the temperature of the air. R is the specific gas constant for dry air and c_p is its specific heat at constant pressure. Eqn. (2.15) is applied at level 2 using the following finite-difference approximation:

$$\left(\frac{\partial \psi}{\partial p} \right)_2 \approx \frac{\psi_3 - \psi_1}{\Delta p} \quad (2.18)$$

Holton (1972), applying Eqns. (2.12) and (2.15) as indicated, finds the three model equations to be:

$$\frac{\partial Q_1}{\partial t} = - J(\psi_1, Q_1) - J(\psi_1, f) + \frac{f_o}{\Delta p} \tilde{\omega}_2 \quad (2.19)$$

$$\frac{\partial Q_3}{\partial t} = - J(\psi_3, Q_3) - J(\psi_3, f) - \frac{f_o}{\Delta p} \tilde{\omega}_2 \quad (2.20)$$

$$\frac{\partial}{\partial t}(\psi_1 - \psi_3) = - J(\psi_2, \psi_1 - \psi_3) + \frac{\sigma \Delta p}{f_o} \tilde{\omega}_2 \quad (2.21)$$

where the Jacobian operator, J , is defined by:

$$J(A,B) = \frac{\partial A}{\partial x} \frac{\partial B}{\partial y} - \frac{\partial A}{\partial y} \frac{\partial B}{\partial x} \quad (2.22)$$

In Eqns. (2.19) - (2.21), the stream function is assumed to be the geostrophic stream function although the subscript denoting the geostrophic value has been dropped here and will be for the remainder of this thesis. The static stability parameter is assumed to be horizontally homogeneous and independent of time.

The stream function for the mean flow, $\overline{\psi}$, and the thickness, $\hat{\psi}$, are defined to be:

$$\overline{\psi} = \frac{\psi_1 + \psi_3}{2} \quad (2.23)$$

$$\hat{\psi} = \frac{\psi_1 - \psi_3}{2} \quad (2.24)$$

Note that $\hat{\psi}$ is really half the thickness between levels 1 and 3.

However, for convenience, $\hat{\psi}$ will be simply referred to as the thickness in this thesis. The stream function at level 2, ψ_2 , is obtained by linearly interpolating between levels 1 and 3 giving:

$$\psi_2 \approx \frac{\psi_1 + \psi_3}{2} \quad (2.25)$$

Thus, in Eqn. (2.21), ψ_2 may be approximated by $\overline{\psi}$.

Forming the sum and difference of Eqns. (2.19) and (2.20) and applying Eqns. (2.23) - (2.25) to (2.19) - (2.21), one obtains:

$$\frac{\partial \bar{\psi}}{\partial t} = - J(\bar{\psi}, \bar{Q}) - J(\hat{\psi}, \hat{Q}) - J(\bar{\psi}, f) \quad (2.26)$$

$$\frac{\partial \hat{Q}}{\partial t} = - J(\hat{\psi}, \bar{Q}) - J(\bar{\psi}, \hat{Q}) - J(\hat{\psi}, f) + \frac{f_0}{\Delta p} \tilde{\omega}_2 \quad (2.27)$$

$$\frac{\partial \hat{\psi}}{\partial t} = - J(\bar{\psi}, \hat{\psi}) + \frac{\sigma \Delta p}{2f_0} \tilde{\omega}_2 \quad (2.28)$$

Finally, Eqns. (2.27) and (2.28) may be combined to yield:

$$(\nabla^2 - \lambda^2) \frac{\partial \hat{\psi}}{\partial t} = - J(\hat{\psi}, \bar{Q}) - J(\bar{\psi}, \hat{Q}) - J(\hat{\psi}, f) + \frac{2f_0^2}{\sigma \Delta p^2} J(\bar{\psi}, \hat{\psi}) \quad (2.29)$$

$$(\nabla^2 - \lambda^2) \tilde{\omega}_2 = - \frac{4f_0}{\sigma \Delta p} J(\hat{\psi}, \bar{Q}) - \frac{2f_0}{\sigma \Delta p} J(\hat{\psi}, f) \quad (2.30)$$

where:

$$\lambda^2 = \frac{2f_0^2}{\sigma \Delta p^2} \quad (2.31)$$

Equations (2.26), (2.29) and (2.30) form a system of three non-linear partial-differential equations in the three unknowns: $\bar{\psi}$, $\hat{\psi}$ and $\tilde{\omega}_2$. Given boundary and initial conditions, a solution to this system of equations may be found. Baer (1970) has found an

analytical solution for a particular set of boundary and initial conditions. For more general solutions, however, numerical methods must be used. Chapters 3 and 4 describe the two numerical methods employed in this thesis.

2.2 The Boundary Conditions

A free-slip wall boundary condition was imposed at the northern and southern boundaries of the domain. Accordingly, there can be no flow perpendicular to the wall and there is no viscous boundary layer at the wall. Using Eqns. (2.7) - (2.9), the first of these conditions implies that:

$$\frac{\partial \psi}{\partial x} = 0 \quad (2.32)$$

at $y = y_0 + L_y/4$ and at $y = y_0 - L_y/4$. The second of these conditions means that no additional terms must be used with the model equations found in Sec. 2.1 at the boundary of the domain. These extra terms would be necessary to account for the turbulence generated near the boundary if free slip was not assumed.

These conditions do not completely specify the solution as ψ is specified by them only to within an arbitrary additive constant. Pedlosky (1979) and Phillips (1954) have shown that, as a consequence of the vanishing of the normal velocity component on the boundary:

$$\int_0^L \frac{\partial^2 \psi}{\partial y \partial t} dx = 0 \quad (2.33)$$

where the integration is made along the northern and southern boundaries separately. Pedlosky (1979) has shown that Eqn. (2.33) is equivalent to:

$$\int_D \frac{\partial \psi}{\partial t} dx dy = 0 \quad (2.34)$$

where the integration occurs over the domain, D.

As the boundary conditions, Eqn. (2.32) and (2.33) or (2.34), apply to the stream functions at each level, they also apply to linear combinations of these stream functions. Thus, they may be applied to the derived stream functions $\bar{\psi}$ and $\hat{\psi}$. Equation (2.33) is the form of the second boundary condition which is used in this thesis. This form requires less computer time when numerically integrated, as it must be when applied to find a solution by the FEM.

2.3 The Energy and Potential Enstrophy Relations

The study of the energetics of the atmosphere and, in particular, of the exchanges of energy among its various forms, among the various scales of motion, and among the sources and sinks of energy,

provides an important tool for understanding atmospheric motion.

Also, the potential enstrophy (PE) may be used as an aid in understanding atmospheric motion. The PE is defined to be the mean squared potential vorticity. Holton (1972) shows that the potential vorticity, q , is a measure of the ratio of the vorticity of a vortex to the depth of the vortex.

These quantities also provide a useful tool for studying a particular atmospheric model, and, they may be used as a means of comparing an atmospheric model with other models or with the real atmosphere. For the chosen model, there are three quantities of interest:

- 1) the kinetic energy (KE)
- 2) the available potential energy (APE)
- 3) the potential enstrophy

The KE and APE are considered to be present in two forms: zonal (ZKE and ZAPE) and eddy (EKE and EAPE). The atmospheric flow is considered to have a basic zonal component with disturbances called "eddies" superimposed on that flow. Thus, any stream function, ψ , may be written as the sum of a zonal, ψ_Z , and an eddy, ψ_E , stream function.

The KE of the horizontal flow is defined as:

$$KE = \int_M \frac{1}{2} (\vec{V} \cdot \vec{V}) \, dM \quad (2.35)$$

where dM is an element of mass and the integration is over the total mass, M , of the atmosphere. For the two-level model under consideration, this may be written as:

$$KE = \frac{\Delta p}{g} \int_D \int (\nabla \bar{\psi} \cdot \nabla \bar{\psi} + \nabla \hat{\psi} \cdot \nabla \hat{\psi}) \, dx dy \quad (2.36)$$

where the hydrostatic relation, $dp/dz = -g\rho$, and $dM = \rho \, dx \, dy \, dz$ have been used. ρ is the atmospheric density, g is the acceleration due to gravity and z is the vertical coordinate parallel to the $-p$ direction.

Differentiating Eqn. (2.36) with respect to time and using the divergence theorem and the boundary conditions, one may show that:

$$\frac{dKE}{dt} = - \frac{2f_0}{g} \int_D \int \tilde{\omega}_2 \hat{\psi} \, dx dy \quad (2.37)$$

This shows that when, in the mean, warm air is ascending and cold air is descending, there is an increase in the kinetic energy, because the centre of gravity is being lowered.

For this model, Holton (1972) shows that the APE is given by:

$$APE = \frac{\Gamma}{2} \int_D \int \hat{\psi}^2 \, dx dy \quad (2.38)$$

where $\Gamma = 4 f_o^2 / g \sigma \Delta p$. Differentiating Eqn. (2.38) with respect to time, one finds that:

$$\frac{dAPE}{dt} = \frac{2f_o}{g} \int_D \int \tilde{\omega}_2 \hat{\psi} \, dx dy \quad (2.39)$$

Combining Eqns. (2.37) and (2.39), one has:

$$\frac{d}{dt}(KE + APE) = 0 \quad (2.40)$$

Thus, the total energy, $TE = KE + APE$, of the model is conserved but an exchange of energy between KE and APE may occur.

Relations similar to Eqns. (2.36) and (2.38) are used to calculate the eddy and zonal components of the KE and APE. Conservation relations, similar to Eqn. (2.40), yield the intuitive result that the total KE (APE) is simply the sum of the ZKE (ZAPE) and the EKE (EAPE).

The potential vorticities at levels 1 and 3 are given by:

$$q_1 = Q_1 - \frac{f_o^2}{\sigma \Delta p^2} (\psi_1 - \psi_3) \quad (2.41)$$

$$q_3 = Q_3 + \frac{f_o^2}{\sigma \Delta p^2} (\psi_1 - \psi_3) \quad (2.42)$$

and the PE is given by:

$$PE = \frac{\Delta p}{2g} \int_D \int (q_1^2 + q_3^2) \, dx dy \quad (2.43)$$

Differentiating Eqns. (2.42) - (2.44) with respect to time and using the divergence theorem, one finds that:

$$\frac{dPE}{dt} = 0 \quad (2.44)$$

Thus, the potential enstrophy is a conservative property of this model. When an approximate numerical solution is found for the model equations, the two conservative quantities, TE and PE, will provide a test of the conservation properties of that solution.

CHAPTER 3

THE SPECTRAL METHOD

3.1 A Description of the Method

This chapter details the method of solution of the model equations, (2.26) - (2.28), using the spectral method, which is an application of a method developed by Galerkin (1915). When this method is applied to the model equations, a highly accurate (but not analytic) solution is obtained. It is found that the solution by the spectral method conserves total energy and potential enstrophy. This spectral solution will be used to judge the accuracy of the solution by the FEM.

A simple example will be used to demonstrate the application of the spectral method to a boundary-value problem. Consider the following equation:

$$\frac{\partial f(x,y,t)}{\partial t} = g(x,y,t) \quad (3.1)$$

This is similar in form to the model equations, (2.26) - (2.28). The functions $f(x,y,t)$ and $g(x,y,t)$ are assumed to be defined on a two-dimensional domain D and the problem is assumed to be properly posed. The functions f and g are represented using a set of $N + 1$ expansion functions, $F_i(x,y)$, to obtain:

$$f(x,y,t) \approx \hat{f}(x,y,t) = \sum_{i=0}^N f_i(t) F_i(x,y) \quad (3.2)$$

$$g(x,y,t) \approx \hat{g}(x,y,t) = \sum_{i=0}^N g_i(t) F_i(x,y) \quad (3.3)$$

Kreider et. al. (1966) show that if the expansion functions form a basis on the domain, the approximate functions, \hat{f} and \hat{g} , will converge to the true functions, f and g , as N approaches infinity. It is assumed that the expansion functions can satisfy the boundary conditions.

The approximate functions are substituted into Eqn. (3.1) using Eqns. (3.2) and (3.3) and a residual, $R(x,y,t)$ is defined according to:

$$R(x,y,t) = \sum_{i=0}^N \frac{df_i}{dt} F_i - \sum_{i=0}^N g_i F_i \quad (3.4)$$

Note that in Eqn. (3.4) the total derivative with respect to time is used as the expansion coefficients, f_i and g_i , depend only on time.

The explicit dependence of the expansion coefficients on time and of the expansion functions, F_i , on the two space variables have been omitted for the sake of brevity. As the approximate functions will not, in general, satisfy Eqn. (3.1) exactly, the residual is not necessarily zero.

In Galerkin's method, the residual is forced to zero with respect to the expansion functions in an average sense over the domain

(Zienkiewicz, 1977). Thus, we insist that:

$$\int_D \int F_j \left(\sum_{i=0}^N \frac{df_i}{dt} F_i - \sum_{i=0}^N g_i F_i \right) dx dy = 0 \quad (3.5)$$

for $j = 0, 1, 2, \dots, N$. This technique gives a system of $N + 1$ ordinary differential equations for the expansion coefficients, $f_i(t)$. This system may then be numerically integrated, given initial conditions, to find the approximate solution, $\hat{f}(x, y, t)$, to Eqn. (3.1).

3.2 The Application of the Spectral Method to the Model Equations

The expansion functions must be specified and they are usually chosen with properties which are particularly appropriate to the problem under consideration. For the present problem, the following set of properties is chosen:

- 1) F_0 is a constant over the domain
- 2) $\partial F_i / \partial x = 0$ along the north and south boundaries
- 3) the F_i are orthonormal i.e.

$$\frac{1}{A} \int_D \int F_i F_j dx dy = \delta_{ij} \quad (3.6)$$

where δ_{ij} is the Kronecker delta and A is the area of the domain

- 4) the expansion functions are eigenfunctions of the Laplacian operator so that:

$$\nabla^2 F_i = -a_i^2 F_i \quad (3.7)$$

where $a_i^2 > 0$

The first property allows one to represent a field with a non-zero mean using the expansion functions. Property 2 is chosen so that the boundary condition, Eqn. (2.32), may be satisfied. Property 3 is chosen to facilitate the integration of terms in equations similar to Eqn. (3.5). Property 4 is chosen as the Laplacian operator is present in the model equations. This property will also simplify the integration of terms in equations similar to Eqn. (3.5).

Expanding the variables $\bar{\psi}$, $\hat{\psi}$ and $\tilde{\omega}_2$, one has:

$$\bar{\psi}(x,y,t) \approx \sum_{i=0}^N \psi_i(t) F_i(x,y) \quad (3.8)$$

$$\hat{\psi}(x,y,t) \approx \sum_{i=0}^N \tau_i(t) F_i(x,y) \quad (3.9)$$

$$\tilde{\omega}_2(x,y,t) \approx \sum_{i=0}^N \omega_i(t) F_i(x,y) \quad (3.10)$$

Using this expansion and Eqn. (3.7), the first model equation, Eqn. (2.26), becomes:

$$\begin{aligned}
- \sum_{i=0}^N a_i^2 \frac{d\psi_i}{dt} F_i &= \sum_{j=0}^N \sum_{k=0}^N a_k^2 \psi_j \psi_k J(F_j, F_k) + \sum_{j=0}^N \sum_{k=0}^N a_k^2 \tau_j \tau_k J(F_j, F_k) \\
&\quad - \beta \sum_{i=0}^N \psi_i \frac{\partial F_i}{\partial x} + R
\end{aligned} \tag{3.11}$$

Applying the Galerkin procedure to Eqn. (3.11) by multiplying by a particular F_ℓ and integrating over the domain, one has:

$$\begin{aligned}
- a_\ell^2 \frac{d\psi_\ell}{dt} &= \sum_{j=0}^N \sum_{k=0}^N a_k^2 \psi_j \psi_k \overline{F_\ell J(F_j, F_k)} + \sum_{j=0}^N \sum_{k=0}^N a_k^2 \tau_j \tau_k \overline{F_\ell J(F_j, F_k)} \\
&\quad - \beta \sum_{i=0}^N \psi_i F_\ell \frac{\partial F_i}{\partial x}
\end{aligned} \tag{3.12}$$

where the averaging operator, $(\overline{\quad})$, is defined by:

$$\overline{(\quad)} = \iint_D (\quad) dx dy \tag{3.13}$$

The terms in Eqn. (3.12) which involve the averaging operator may only be evaluated once the expansion functions have been chosen. These terms, however, are independent of space and time. They are constants which depend only on the expansion functions and the domain of the problem and they may be written:

$$c_{\ell jk} \equiv \overline{F_{\ell}^J(F_j, F_k)} \quad (3.14)$$

$$b_{\ell i} \equiv F_{\ell} \overline{\frac{\partial F_i}{\partial x}} \quad (3.15)$$

Applying the Galerkin procedure to the three model equations, Eqns. (2.26) - (2.28), and using Eqns. (3.14) and (3.15), one obtains:

$$- a_{\ell}^2 \frac{d\psi_{\ell}}{dt} = \sum_{j=0}^N \sum_{k=0}^N a_k^2 c_{\ell jk} (\psi_j \psi_k + \tau_j \tau_k) - \beta \sum_{i=0}^N b_{\ell i} \psi_i \quad (3.16)$$

$$- a_{\ell}^2 \frac{d\tau_{\ell}}{dt} = \sum_{j=0}^N \sum_{k=0}^N a_k^2 c_{\ell jk} (\tau_j \psi_k + \psi_j \tau_k) - \beta \sum_{i=0}^N b_{\ell i} \tau_i + \frac{f_o}{\Delta p} \omega_{\ell} \quad (3.17)$$

$$\frac{d\tau_{\ell}}{dt} = - \sum_{j=0}^N \sum_{k=0}^N c_{\ell jk} \psi_j \tau_k + \frac{\sigma \Delta p}{2f_o} \omega_{\ell} \quad (3.18)$$

Using Eqns. (3.14) and (2.22), one may show that:

$$\sum_{j=0}^N \sum_{k=0}^N X_k Y_{jk} c_{\ell jk} = \frac{1}{2} \left(\sum_{j=0}^N \sum_{k=0}^N (X_k - X_j) Y_{jk} c_{\ell jk} \right) \quad (3.19)$$

for arbitrary X_k and for all Y_{jk} where $Y_{jk} = Y_{kj}$. X_k and Y_{jk} represent expansion coefficients or combinations of expansion coefficients which depend only on the given indices. Eqn. (3.19) may be applied to Eqns. (3.16) - (3.18) to give:

$$\frac{d\psi_\ell}{dt} = \frac{1}{2} \left(\sum_{j=0}^N \sum_{k=0}^N \frac{(a_j^2 - a_k^2)}{a_\ell^2} c_{\ell jk} (\psi_j \psi_k + \tau_j \tau_k) \right) + \frac{\beta}{a_\ell^2} \sum_{i=0}^N b_{\ell i} \psi_i \quad (3.20)$$

$$\frac{d\tau_\ell}{dt} = \frac{1}{2} \left(\sum_{j=0}^N \sum_{k=0}^N \frac{(a_j^2 - a_k^2)}{a_\ell^2} c_{\ell jk} (\tau_j \psi_k + \tau_k \psi_j) \right) + \frac{\beta}{a_\ell^2} \sum_{i=0}^N b_{\ell i} \tau_i - \frac{f_o}{\Delta p} \frac{\omega_\ell}{a_\ell^2} \quad (3.21)$$

$$\frac{d\tau_\ell}{dt} = \frac{1}{2} \left(\sum_{j=0}^N \sum_{k=0}^N c_{\ell jk} (\tau_j \psi_k - \tau_k \psi_j) \right) + \frac{\sigma \Delta p}{2f_o} \omega_\ell \quad (3.22)$$

Eqns. (3.20) - (3.22) yield a system of 3 (N + 1) ordinary differential equations for the 3 (N + 1) expansion coefficients. Before solving these, the expansion functions must be chosen. They must satisfy the four criteria given earlier. Following Lorenz (1960), the choice of expansion functions is:

$$F_0 = 1$$

$$F_1 = \sqrt{2} \cos \frac{2\pi y}{L_y}$$

$$F_2 = 2 \sin \frac{2\pi y}{L_y} \sin \frac{2\pi nx}{L_x}$$

$$F_3 = 2 \sin \frac{2\pi y}{L_y} \cos \frac{2\pi nx}{L_x}$$

(3.23)

This is referred to as a truncated spectral series expansion.

Applying Eqn. (3.7) to these functions, one finds:

$$a_0^2 = 0 \quad (3.24)$$

$$a_1^2 = (2\pi/L_y)^2 \quad (3.25)$$

$$a_2^2 = a_3^2 = (2\pi/L_y)^2 + (2\pi n/L_x)^2 \quad (3.26)$$

This set of functions allows only a wave with wavenumber n to exist in the x -direction. This particular wavenumber will be left to be chosen later. In the y -direction, only waves of one wavenumber are allowed and this is completely specified with the choice of L_y given earlier. In order to fully describe atmospheric flows, which are characterized by many wavelengths, many more expansion functions would be necessary. Limiting the model to waves of only one wavelength, however, simplifies the solution considerably. This set of functions does, however, allow one to study the non-linear interactions between waves of the same wavelength in the fields of the stream functions for the mean flow (or mean height) and the thickness. The amplitudes of the waves and the phase difference between the waves may be varied to simulate aspects of various atmospheric flow patterns.

The expansion function F_0 is used to specify the mean height of the field. It is irrelevant to the present problem, however, and will be dropped. As proof of its irrelevance, Eqn. (3.14) demonstrates that with the chosen expansion functions, $c_{\ell j k} = 0$ if any of $\ell = 0$,

$j = 0$, or $k = 0$. Eqn. (3.15) shows that $b_{\ell i} = 0$ if $\ell = 0$ or $i = 0$. Thus, the mean of the field does not interact with the perturbation in the field and it is, therefore, not important. Consequently, the expansion coefficients associated with F_0 will be assumed to be 0.

For the chosen domain and the three expansion functions, F_1 , F_2 and F_3 , Eqn. (3.15) gives:

$$b_{1\ell} = 0 \quad \text{for } \ell = 1, 2, 3 \quad (3.27)$$

$$b_{23} = -b_{32} = -2\pi n/L_x \quad (3.28)$$

Eqn. (3.14) shows that $c_{\ell j k}$ is non-zero only when the three indices are distinct and one finds that:

$$c_{321} = -\frac{32\sqrt{2}\pi n}{3L_x L_y} \quad (3.29)$$

The dependent variables, $\bar{\psi}$, $\hat{\psi}$ and $\tilde{\omega}_2$, are expanded using the expansion functions to obtain:

$$\bar{\psi} = \psi_1 F_1 + \psi_2 F_2 + \psi_3 F_3 \quad (3.30)$$

$$\hat{\psi} = \tau_1 F_1 + \tau_2 F_2 + \tau_3 F_3 \quad (3.31)$$

$$\tilde{\omega}_2 = \omega_1 F_1 + \omega_2 F_2 + \omega_3 F_3 \quad (3.32)$$

Finally, Eqns. (3.23) - (3.32) are substituted into Eqns. (3.20) - (3.22) to obtain a system of nine ordinary differential equations in the nine expansion coefficients. These equations are non-dimensionalized as this will give a simpler form for the final equations. Thus, in the equations, an arbitrary expansion coefficient for the mean flow, ψ_i , is replaced by $L^2 f_o \psi_i$. This new coefficient is non-dimensional. Similarly, τ_i is replaced by $L^2 f_o \tau_i$; ω_i is replaced by $\Delta p f_o \omega_i$, and d/dt is replaced by $f_o d/dt$.

Simple algebraic manipulation allows one to rewrite the resulting nine non-dimensional equations as a system of six ordinary differential equations and three diagnostic omega equations:

$$\frac{d\psi_1}{dt} = 0 \quad - \quad (3.33)$$

$$\frac{d\psi_2}{dt} = n^2 \nu \mu (\psi_1 \psi_3 + \tau_1 \tau_3) - \beta^* n \mu \psi_3 \quad (3.34)$$

$$\frac{d\psi_3}{dt} = - n^2 \nu \mu (\psi_1 \psi_2 + \tau_1 \tau_2) - \beta^* n \mu \psi_2 \quad (3.35)$$

$$\frac{d\tau_1}{dt} = - \frac{\omega_1}{\alpha^2} \quad (3.36)$$

$$\frac{d\tau_2}{dt} = n^2 \nu \mu (\tau_1 \psi_3 + \tau_3 \psi_1) - \beta^* n \mu \tau_3 - \mu \omega_2 \quad (3.37)$$

$$\frac{d\tau_3}{dt} = - n^2 \nu \mu (\tau_1 \psi_2 + \tau_2 \psi_1) + \beta^* n \mu \tau_2 - \mu \omega_3 \quad (3.38)$$

$$\omega_1 = \frac{-\nu}{\sigma^* + \frac{1}{\alpha^2}} (\tau_2 \psi_3 - \tau_3 \psi_2) \quad (3.39)$$

$$\omega_2 = \frac{1}{\sigma^* + \mu} \left(\nu \{ n^2 \mu (\tau_1 \psi_3 + \tau_3 \psi_1) - (\tau_3 \psi_1 - \tau_1 \psi_3) \} - \beta^* n \mu \tau_3 \right) \quad (3.40)$$

$$\omega_3 = \frac{1}{\sigma^* + \mu} \left(\nu \{ (\tau_2 \psi_1 - \tau_1 \psi_2) - n^2 \mu (\tau_1 \psi_2 + \tau_2 \psi_1) \} + \beta^* n \mu \tau_3 \right) \quad (3.41)$$

where the following symbol definitions have been used:

$$\alpha = L_x / L_y \quad (3.42)$$

$$L = L_x / 2\pi \quad (3.43)$$

$$\gamma = 8\sqrt{2}/3\pi \quad (3.44)$$

$$\beta^* = \beta L / f_o \quad (3.45)$$

$$\sigma^* = \frac{\Delta p^2 \sigma}{2 f_o^2 L^2} \quad (3.46)$$

$$\nu = \gamma \alpha n \quad (3.47)$$

$$\mu = 1/(\alpha^2 + n^2) \quad (3.48)$$

Thus, the spectral method requires the solution of six non-linear ordinary differential equations, Eqns. (3.33) - (3.38), and three omega equations, Eqns. (3.39) - (3.41), for the nine expansion coefficients. Baer (1970) has found an analytical solution to these equations to be expressible in terms of elliptic functions. The analytical solution was not used for this thesis. Instead, a numerical solution was found because a highly accurate solution could be found numerically with considerably less computational effort than an evaluation of the analytical solution would require.

Equations (3.33) - (3.38) have the general form:

$$\frac{df}{dt} = g(t) \quad (3.49)$$

In order to integrate such an equation numerically, it is written in finite-difference form. For the first time step, an Euler forward difference is used:

$$f^1 = f^0 + \Delta t g^0 \quad (3.50)$$

where f^0 and g^0 are the initial values of the functions f and g and Δt is the time step. Ralston (1965) shows that this scheme has first-order accuracy. For subsequent time steps, the second-order Adams-Bashforth method is used:

$$f^{n+1} = f^n + \frac{\Delta t}{2}(3g^n - g^{n-1}) \quad (3.51)$$

Lilly (1965) found the Adams-Bashforth method to be simple and efficient while giving accuracy on a par with more complicated methods when he compared 8 methods for integrating equations similar to Eqns. (3.37) - (3.42).

As the spectral solution is to be used as an accurate solution with which to compare the finite-element solution, the minimization of time truncation error is very important. Haltiner (1971) indicates that the time step must be a "small" fraction of the period of variation of the spectral amplitudes. As synoptic-scale motions are being studied, the period of the spectral amplitudes is expected to be on the order of days. Thus, the qualitative criterion of Haltiner (1971) indicates that a time step on the order of hours is appropriate. Test integrations were done with time steps of 1 hour and 0.5 hours. After 24 hours of integration, the solutions with the different time steps were compared. The spectral amplitudes were found to vary by less than 1 part in 10^4 . These results suggest a time step of 0.5 hours is sufficiently small and this is used in all spectral integrations presented in Chapter 5.

The initial conditions are determined by choosing values for ψ_1^0 , ψ_2^0 , ψ_3^0 , τ_1^0 , τ_2^0 and τ_3^0 . In choosing these values, one determines not only the amplitudes of the waves but also the phase relationship between the waves in the stream function for the mean height, $\bar{\psi}$, and for the thickness, $\hat{\psi}$. This allows one to simulate a variety of atmospheric flow patterns. The initial conditions for the cases presented in Chapter 5 were chosen to represent a variety of flows. This is

done so that the solution by the finite-element method could be tested with a variety of initial conditions in order to determine its possible strengths or weaknesses.

3.3 The Energy and Potential Enstrophy Relations

The energies and potential enstrophy discussed in Sec. 2.3 can be related to the expansion coefficients. Using the definition of the kinetic energy, Eqn. (2.36), and Eqns. (3.30) and (3.31), one finds that:

$$KE = \frac{2L_x L_y \Delta p L^4 f_0^2}{g} \left(\frac{1}{L_y^2} (\psi_1^2 + \tau_1^2) + \left(\frac{1}{L_y^2} + \frac{n^2}{L_x^2} \right) (\psi_2^2 + \psi_3^2 + \tau_2^2 + \tau_3^2) \right) \quad (3.52)$$

where non-dimensionalized coefficients are used. Using Eqn. (2.38), the available potential energy of the spectral solution is found to be given by:

$$APE = \frac{\Gamma L_x L_y f_0^2 L^4}{4} (\tau_1^2 + \tau_2^2 + \tau_3^2) \quad (3.53)$$

The potential enstrophy for the spectral solution, using Eqns. (2.41), (2.42) and (2.43), is found to be:

$$PE = \frac{\Delta p L_x L_y f_0^2}{2g} \left(\{ \alpha^2 \psi_1 \}^2 + \{ (\alpha^2 + n^2) \psi_2 \}^2 + \{ (\alpha^2 + n^2) \psi_3 \}^2 + \{ (\alpha^2 + 1/\sigma^*) \tau_1 \}^2 + (\alpha^2 + n^2 + 1/\sigma^*) \{ \tau_2^2 + \tau_3^2 \} \right) \quad (3.54)$$

Differentiating Eqns. (3.52) - (3.54) with respect to time and substituting Eqns. (3.33) - (3.41) in, it is found that the total energy and the potential enstrophy are conserved by the exact spectral solution. Eqns. (3.52) - (3.54) allow the calculation of these quantities at each step of the integration. The conservation of these quantities was used as a method of verifying the accuracy of the spectral numerical solution. In practice, it was found that these quantities were conserved to better than 1 in 10^4 for a wide variety of initial conditions and lengths of integrations. In particular, it is true for all cases presented in Chapter 5.

CHAPTER 4

THE FINITE-ELEMENT METHOD

4.1 A Description of the Method

Since its initial development by structural engineers during the 1950's (e.g. Turner et. al., 1956; Argyris, 1960; Clough, 1960), the finite-element method (FEM) has become a popular means of finding approximate solutions to initial and/or boundary value problems. In recent years, (Wang, et. al., 1972; Cullen, 1976, 1979; Staniforth and Daley, 1978, 1979), the FEM has been applied to the initial-boundary-value problems of numerical weather prediction. This chapter describes the application of the FEM to the atmospheric model described in Chapter 2. A brief history of the method is given first and a discussion of the FEM follows, with a simple example used to clarify the discussion. Finally, the application of the FEM to the atmospheric model under consideration is presented.

It is difficult to determine the originator of the FEM, although Clough (1960) seems to have been the first to use the name. Until the early 1960's, the method was developed separately by mathematicians and engineers. On the engineering side, the FEM evolved from the matrix method of structural analysis (Zienkiewicz, 1977). In this method, the analysis of structures proceeded by considering the components of the structures separately. Relations between the displacements and internal forces at the nodal points of individual components were derived in matrix form with the displacements and/or the

forces being unknown. (The nodal points are the places where the components were joined.) The solution for the unknowns proceeded by solving the system of equations, written in matrix form, for the unknowns.

This method provided the exact answer for the unknowns. The only assumption made was that mathematical relations could be used to describe real physical systems. By analogy, McHenry (1943) and Newmark (1949), for example, extended the matrix method to continuum problems, i.e. problems without easily identifiable components. They divided the continuum into a number of hypothetical components called elements. They then proceeded as before by writing a system of equations for the nodal displacements and forces and solving this system of equations. This was found to give a good approximate solution to the original continuum problem.

As the method evolved, it was found that the simplest procedure to ensure that the forces and displacements of the approximate solution represented accurately the true solution, was to introduce the concept of virtual work (Zienkiewicz, 1977). An arbitrary (i.e. virtual) nodal displacement is imposed and the internal and external work done by the various forces and stresses during that displacement are equated. It was recognized that this approach was equivalent to minimizing the total potential energy of the structure under consideration. Argyris (1960), for example, detailed the resulting matrix equations, for a rectangular panel under plane stress, in a comprehensive paper on energy theorems and matrix methods.

During the late 1950's and early 1960's (e.g. Szmelter, 1959; Clough, 1965), this method was recognized as an extension of the Rayleigh-Ritz method, which had been well known to mathematicians since the publications of Rayleigh (1870) and Ritz (1909). With this recognition, it became possible to give a mathematical basis to the largely intuitive developments of the engineering profession. During the last two decades, the use of the FEM has grown rapidly and a bibliography of the FEM by Norrie and deVries (1976) provides a detailed listing of many of the developments.

When applying the FEM, the domain of the problem is divided into a number of elements. The number, size and shape of the elements are chosen after consideration of the domain of the problem and the degree of accuracy required. Zienkiewicz (1977) discusses some of the possible choices for elemental shapes. Triangular and rectangular shapes have been popular as they are relatively easy to work with. There are no hard and fast rules for choosing the elements although there are a few guidelines. For a domain with a curved boundary, triangular elements offer an advantage as they can better approximate the boundary. For rectangular domains, rectangular elements may be advantageous as fewer elements are normally required for a given level of accuracy. Also, increasing the total number of elements and decreasing the size of the elements provides increased accuracy. One has the option of using small elements in areas of detailed interest while using larger elements elsewhere. When all the elements are chosen for a domain, the result is called the mesh or grid of the domain. It is advantageous to

automate the choice of this mesh for a given elemental shape and domain.

In this thesis, rectangular elements are used, and Fig. 4.1 shows the mesh used for the domain defined in Sec. 2.1. The domain is divided into unequal rectangles. The central portion of the mesh has high resolution with the elements having a length of 200 km. on a side. The element length and width varies uniformly away from this central portion. Thus, the ratio of elemental lengths between neighbouring elements is a constant. The elemental length reaches a maximum of 1200 km. and the elemental width reaches a minimum of 60 km. The high resolution portion near the north and south boundaries is used to permit a more accurate implementation of the boundary conditions, Eqns. (2.33) and (2.34). A variable grid length was used so that its effect on the solution could be investigated. The portions of the grid marked A and B in Fig. 4.1 are the portions of the domain in which the FEM solution will be presented (in Chapter 5) for comparison with the spectral solution. The portion marked C is the portion on which the energies and the potential enstrophy of the two solutions will be calculated.

The elemental shape must be chosen in conjunction with the basis functions, $e^i(x,y)$, to be used and the number of nodes per element. The basis functions form an interpolatory basis with which a function may be interpolated on the domain. They are normally chosen to be low-order polynomials and are defined in a piece-wise sense on the domain, i.e. the i 'th basis function is non-zero only over some (small) portion of the domain adjacent to element i . The choice of piece-wise defined

basis functions is a major strength of the FEM. It leads to systems of linear equations which may be written in matrix form. These matrices are normally highly sparse and banded. This gives a significant computational advantage over systems that are nearly full or not banded.

In general, when interpolating a function, one needs the basis functions and some values of the function being interpolated. In the FEM, these values are the value of the function at the nodes. The nodes are specified points on the domain and the number of nodes chosen per element determines the order of the basis functions. Thus, with rectangular elements, if four nodes are chosen per element, the basis functions are bi-linear while the use of sixteen nodes per element would be consistent with bi-quadratic basis functions. Although the nodes are often chosen to be on the boundary of the elements, they may be in the interior of the element. The number of nodes per element, the nodal positions and the basis functions must be chosen together. The reader is referred to Zienkiewicz (1977) and Tong and Rossettos (1977) for examples of elements and their associated nodes and basis functions which have been used.

In this thesis, four nodes per element are used and the nodes are chosen to be in the corners of the elements. Each node is common to four elements.

Bi-linear basis functions are used. One may define the basis functions with respect to each element or, equivalently and more simply

in this case, one may define them with respect to the nodes. If the nodes are numbered $m = 1, 2, \dots, NI$ in the x -direction and $n = 1, 2, \dots, NJ$ in the y -direction, the basis function for node (m,n) may be written:

$$\left(\frac{x-x_{m-1}}{h_{m-1}} \right) \left(\frac{y-y_{n-1}}{k_{n-1}} \right) : x \in (x_{m-1}, x_m), y \in (y_{n-1}, y_n)$$

$$\left(\frac{x_{m+1}-x}{h_m} \right) \left(\frac{y-y_{n-1}}{k_{n-1}} \right) ; x \in (x_m, x_{m+1}), y \in (y_{n-1}, y_n)$$

$$e^{\ell}(x,y) = \left(\frac{x-x_{m-1}}{h_{m-1}} \right) \left(\frac{y_{n+1}-y}{k_n} \right) ; x \in (x_{m-1}, x_m), y \in (y_n, y_{n+1}) \quad (4.1)$$

$$\left(\frac{x_{m+1}-x}{h_m} \right) \left(\frac{y_{n+1}-y}{k_n} \right) ; x \in (x_m, x_{m+1}), y \in (y_n, y_{n+1})$$

$$0 \quad ; \text{ otherwise}$$

where:

$$h_m = x_{m+1} - x_m \quad (4.2)$$

$$k_n = y_{n+1} - y_n$$

and ℓ is the multi-integer (m,n) . The vector $[X_1, X_2, \dots, X_{NI}]$ contains the values of the x -coordinate at the nodes while $[Y_1, Y_2, \dots, Y_{NJ}]$ contains the values of the y -coordinate.

Thus, $e^{\ell}(x,y)$ is non-zero only over the four rectangles which have node ℓ in common. Also, $e^{\ell}(x,y)$ has a value of 1 at the node and diminishes to a value of zero at the other nodes of those four rectangles. With this definition, these basis functions are almost orthogonal, i.e. they interact only locally. Thus, the integral defined by:

$$I = \int_D \int e^{\ell}(x,y) e^k(x,y) dx dy \quad (4.3)$$

where ℓ is the multi-integer (m,n) and k is the multi-integer (M,N) , is non-zero only if $m = M - 1, M$ or $M + 1$ and $n = N - 1, N$ or $N + 1$. Finally, it should be noted that the basis functions are separable, i.e. they may be written $e^{\ell}(x,y) = e^m(x) e^n(y)$. These basis functions may be used to interpolate a function, $g(x,y)$, on a domain in the following way:

$$g(x,y) \approx \sum_i g_i e^i(x,y) \quad (4.4)$$

where the summation extends over all nodes and the g_i are the nodal values of $g(x,y)$.

The FEM is, in fact, a general class of methods and one must choose the particular method(s) to be used for a given problem. Zienkiewicz (1977) provides a discussion of some of the methods which have been used. Two large sub-classes of the FEM are the variational

and the weighted residual approaches. An example of a method from each of these sub-classes is considered in this thesis. First, the Rayleigh-Ritz method (a variational approach) and then the Galerkin method (a weighted residual approach) will be discussed. The Galerkin method is the one actually used for the solution of the model equations.

To illustrate the Rayleigh-Ritz method, let us consider the two-dimensional Poisson equation:

$$\nabla^2 \phi = f(x,y) \quad (4.5)$$

valid on a domain D with the boundary condition:

$$\phi(x,y) = \bar{\phi} \quad (4.6)$$

on the boundary, Γ , of the domain where $\bar{\phi}$ is a constant. A variational principle may be written:

$$I(\phi) = \frac{1}{2} \int_D \int (\nabla \phi)^2 \, dx dy + \int_D \int f(x,y) \phi \, dx dy \quad (4.7)$$

where $(\nabla \phi)^2 = (\partial \phi / \partial x)^2 + (\partial \phi / \partial y)^2$. Tong and Rossettos (1977) have shown that the function, ϕ , which minimizes Eqn. (4.7) and satisfies Eqn. (4.6), is also the solution of Eqns. (4.5) and (4.6). Thus, the solution of the original differential equation is also the extremum of the variational principle.

In the Rayleigh-Ritz method, an approximate solution to Eqns. (4.5) and (4.6) is found by minimizing an approximate form of

the variational principle, Eqn. (4.7). The functions ϕ and f are expanded as in Eqn. (4.4) and are substituted into Eqn. (4.7). This yields an approximate form of the variational, $I^*(\phi)$, and this is minimized with respect to changes in each of the nodal values of ϕ , i.e.

$$\frac{\partial I^*}{\partial \phi_i} = 0 \quad ; \quad i=1,2,\dots,N \quad (4.8)$$

where N is the total number of interior nodes. This procedure yields a system of linear equations which may be written in matrix form as:

$$A\phi = Bf \quad (4.9)$$

where A and B are square $N \times N$ matrices and ϕ and f are vectors of the nodal values of their respective functions. Given the nodal values of f , one may find the nodal values of ϕ by inverting A in Eqn. (4.9). This is the approximate solution to Eqns. (4.5) and (4.6).

Many boundary-value problems, however, do not have a corresponding variational principle and other methods must be used to find solutions to them. The Galerkin method is one of these. The present method is an extension of that given by Galerkin (1915). In the general weighted residual approach, the differential equation, Eqn. (4.5), is multiplied by a test function, $t(x,y)$, and the product is integrated over the domain to yield:

$$\int_D \int_D t \nabla^2 \phi \, dx dy = \int_D \int_D t f \, dx dy \quad (4.10)$$

In the Galerkin method, the test function is chosen to be one of the basis functions.

If the derivatives in Eqn. (4.10) are of higher order than are the basis functions, it is necessary to integrate Eqn. (4.10) by parts. In two-dimensions, this can be accomplished using Green's Theorem. Following Zienkiewicz (1977), Eqn. (4.10) may be written:

$$- \int_D \int_D \left(\frac{\partial t}{\partial x} \frac{\partial \phi}{\partial x} + \frac{\partial t}{\partial y} \frac{\partial \phi}{\partial y} \right) dx dy = \int_D \int_D t f \, dx dy - \oint_{\Gamma} n_x t \frac{\partial \phi}{\partial x} d\Gamma - \oint_{\Gamma} l_y t \frac{\partial \phi}{\partial y} d\Gamma \quad (4.11)$$

where n_x and l_y are the direction cosines between the outward normal to the boundary and the x and y axes respectively. With the boundary condition Eqn. (4.6), the last two terms of Eqn. (4.11) are zero.

Now, ϕ and f are expanded using Eqn. (4.4) and substituted into Eqn. (4.11). Upon integration, a system of linear equations similar to Eqn. (4.9) is obtained and the approximate solution may be found by inverting the new matrix A . Strang and Fix (1973) have shown that the Galerkin and Rayleigh-Ritz methods give the same system of equations if the problem has a variational form.

The right-hand-side of Eqn. (4.5) has, to this point, been treated as a simple function, $f(x,y)$. In fact, as in the model equation (2.26) for example, the right-hand-side may be a rather complicated expression. The nodal values of this expression must be evaluated (to give the nodal values of f) so that the solution may proceed via the Galerkin or Rayleigh-Ritz methods. Staniforth and Daley (1978) provide a detailed explanation of how this expression may be evaluated and the following brief explanation is based on their work.

The evaluation of complicated right-hand-sides of equations such as Eqn. (2.26) is most easily understood by breaking the procedure up into a number of smaller steps. In general, the right-hand-side involves the addition or subtraction of functions, the differentiation of functions and the products of functions. Each of these will be considered in turn.

The addition or subtraction of functions proceeds node by node. Thus, if the nodal values f_i and g_i of the functions $f(x,y)$ and $g(x,y)$ are known, the sum or difference, $b(x,y) = f(x,y) \pm g(x,y)$, is found using:

$$b_i = f_i \pm g_i \quad (4.12)$$

for all the nodes, $i = 1, 2, \dots, N$.

The differentiation of functions will be considered for a one-dimensional domain first. Let us consider the determination of a multiple of the first derivative of a function, $u(x)$, in the x -direction:

$$v = \alpha u_x \quad (4.13)$$

where $u_x \equiv du/dx$ and α is a scalar constant. It is assumed that u is known at the nodal points x_i , $i = 1, 2, \dots, N$ and that v is required at these points. The two functions, u and v , are expanded using the basis functions, $e^i(x)$, and substituted into Eqn. (4.13) to obtain:

$$v_i e^i(x) = u_i e_x^i(x) \quad (4.14)$$

where $e_x^i(x) = d e^i(x)/dx$ and where the Einstein summation convention for repeated indices is used. The Galerkin procedure is applied by multiplying by each of the basis functions successively and integrating over the domain, $[x_1, x_N]$, to obtain

$$v_i \int_{x_1}^{x_N} e^k(x) e^i(x) dx = \alpha u_i \int_{x_1}^{x_N} e^k(x) e_x^i(x) dx \quad (4.15)$$

for $k = 1, 2, \dots, N$. This may be written in matrix form as:

$$\underline{Pv} = \alpha \underline{P_x u} \quad (4.16)$$

where:

$$\underline{v} = (v_1, v_2, \dots, v_N)^T$$

$$\underline{u} = (u_1, u_2, \dots, u_N)^T$$

(4.17)

$$P^{ki} = \int_{x_1}^{x_N} e^k e^i \, dx$$

$$P_x^{ki} = \int_{x_1}^{x_N} e^k e_x^i \, dx$$

P^{ki} is the (k,i) element of an $N \times N$ matrix and $(\quad)^T$ is the transpose of that vector.

As the basis functions are nearly orthogonal, the matrices P and P_σ are tri-diagonal. Staniforth and Daley (1978) show that P is diagonally dominant and that the solution of this system of equations by Gaussian elimination is, therefore, stable with respect to round-off error. As a consequence of the separability of the basis functions, Staniforth and Daley (1978) find that a partial derivative in a multi-dimensional field may be obtained by taking the appropriate ordinary derivative in the direction defined by the partial derivative.

Consider now the evaluation of the product of two functions, $u(x)$ and $v(x)$, defined, again for the sake of simplicity, on a one-

dimensional domain:

$$w = \alpha uv \quad (4.18)$$

where α is a scalar constant. The three functions, u , v and w , are expanded using the basis functions and the Galerkin procedure is again applied to obtain:

$$w_i \int_{x_1}^{x_N} e^k(x) e^i(x) dx = \alpha u_i v_j \int_{x_1}^{x_N} e^k(x) e^i(x) e^j(x) dx \quad (4.19)$$

which is valid for $k = 1, 2, \dots, N$. The left-hand-side will be recognized as the matrix P defined by Eqn. (4.17). The right-hand-side of Eqn. (4.19) is more complicated than before and an efficient method for evaluating it is necessary. The left-hand-side is evaluated using numerical integration and, as the integrand is cubic in x , Simpson quadrature, a quadrature formula which is exact for cubics, is used. Over the range $[x_n, x_{n+1}]$, it is found that:

$$\int_{x_1}^{x_{n+1}} e^n(x) u(x) v(x) dx = \frac{(x_{n+1} - x_n)}{6} (u_n v_n + \frac{1}{2} \{u_n + u_{n+1}\} \{v_n + v_{n+1}\}) \quad (4.20)$$

and that:

$$\int_{x_n}^{x_{n+1}} e^{n+1}(x)u(x)v(x) dx = \frac{(x_{n+1}-x_n)}{6} (u_{n+1}v_{n+1} + \frac{1}{2}\{u_n+u_{n+1}\}\{v_n+v_{n+1}\}) \quad (4.21)$$

For a multi-dimensional domain, Staniforth and Daley (1978) show that product terms like Eqn. (4.18) may be reduced to weighted sums of products evaluated on a one-dimensional domain. That is, one numerically integrates in one direction and then integrates along the other direction.

Thus, the results which are necessary to evaluate the nodal values of f on the right-hand-side of Eqn. (4.9) have been established for bi-linear basis functions on a two-dimensional domain. In Section 4.2, we will consider the application of the FEM to the model equations, Eqns. (2.26), (2.29) and (2.30). The nodal values of the right-hand-sides of these equations were evaluated using the methods of this section.

4.2 The Application of the FEM to the Model Equations

The model equations, Eqns. (2.26), (2.29) and (2.30), are of the form:

$$\nabla^2 \phi - \lambda^2 \phi = f(x,y) \quad (4.22)$$

which is a Helmholtz equation. Staniforth and Mitchell (1977) have shown that the variational principle corresponding to Eqn. (4.22) is:

$$I(\phi) = \int_D \int \left(\{\nabla \phi\}^2 + \lambda^2 \phi^2 + 2f\phi \right) dx dy \quad (4.23)$$

given the boundary condition Eqn. (2.33). Here, $(\nabla \phi)^2 = (\partial \phi / \partial x)^2 + (\partial \phi / \partial y)^2$.

Staniforth and Mitchell (1977, 1978) have shown that with the basis functions defined by Eqn. (4.1), the approximate minimization of $I(\phi)$ in Eqn. (4.23), using the Rayleigh-Ritz method, yields a solution with second-order accuracy on any sub-domain with uniform grid spacing. They discuss an alternate approach using the Galerkin method which yields a fourth-order solution on any uniform sub-domain. This is the method used in this thesis.

Applying the Galerkin method to Eqn. (4.22) yields:

$$\int_D \int \nabla^2 \phi e^k(x) dx dy - \lambda^2 \int_D \int \phi e^k(x) dx dy = \int_D \int f e^k(x) dx dy \quad (4.24)$$

where λ is assumed to be a constant over the domain. If ϕ and f are expanded using the bi-linear basis functions, the second-order solution

obtainable from the Rayleigh-Ritz method is found. Eqn. (4.24) may be written:

$$\int_{x_{i-1}}^{x_{i+1}} \int_{y_{j-1}}^{y_{j+1}} (\phi_{xx} + \phi_{yy}) e^k dx dy - \lambda^2 \int_{x_{i-1}}^{x_{i+1}} \int_{y_{j-1}}^{y_{j+1}} \phi e^k dx dy = \int_{x_{i-1}}^{x_{i+1}} \int_{y_{j-1}}^{y_{j+1}} f e^k dx dy \quad (4.25)$$

as the basis function e^k for the node (i,j) is non-zero only over the four rectangles adjoining that node.

To demonstrate the technique for finding the fourth-order solution, consider the element which has the node (i,j) in its lower left hand corner, i.e. for which we have $x_i \leq x \leq x_{i+1}$ and $y_j \leq y \leq y_{j+1}$. For this element, the right-hand-side of Eqn. (4.25) may be written:

$$R = \int_{x_i}^{x_{i+1}} \int_{y_j}^{y_{j+1}} f e^k dx dy \quad (4.26)$$

Following Staniforth and Mitchell (1978), f is assumed to be symmetric about x_i , y_j and R is therefore evaluated as one-quarter of the doubly symmetric integral over four times the area, viz:

$$R = \frac{1}{4} \int_{x_i - H_i}^{x_i + H_i} \int_{y_j - K_j}^{y_j + K_j} f e^k dx dy \quad (4.27)$$

where $H_i = x_{i+1} - x_i$ and $K_j = y_{j+1} - y_j$. Note that in Eqn. (4.27), e^k must be re-defined for the new assumed grid and that f is expanded in terms of this new e^k . This interpolate for f is substituted into Eqn. (4.27) and the integration is performed yielding an expression for R in terms of the nodal values of f . Similar integrations of the right-hand-side of Eqn. (4.25) are performed over the three remaining rectangles adjoining (i,j) . A similar procedure is undertaken to obtain the left-hand-side, but an integration by parts is necessary first. For the node (i,j) , this yields Eqn. (4.28), which Staniforth and Mitchell (1978) show has fourth-order accuracy on any uniform sub-domain. (Eqn. (4.28) is on page 54.)

An equation similar to Eqn. (4.28) may be derived for all interior nodes. These equations form a system of linear equations in the nodal values of the functions ϕ and f . To complete this system of equations, it is necessary to impose the boundary conditions.

Denoting the north boundary with $j = 1$ and the south boundary with $j = NJ$, the first boundary condition, Eqn. (2.32), gives:

$$\phi_{1,1} = \phi_{2,1} = \dots = \phi_{NI,1} = \phi_{B,1} \quad (4.29)$$

and:

$$\phi_{1,NJ} = \phi_{2,NJ} = \dots = \phi_{NI,NJ} = \phi_{B,NJ} \quad (4.30)$$

$$\begin{aligned}
& \frac{1}{12} \left(\phi_{i+1,j+1} \left(\frac{k_j}{h_i} + \frac{h_i}{k_j} \right) + \phi_{i-1,j+1} \left(\frac{k_j}{h_{i-1}} + \frac{h_{i-1}}{k_j} \right) + \phi_{i-1,j-1} \left(\frac{k_{j-1}}{h_i} + \frac{h_i}{k_{j-1}} \right) + \phi_{i+1,j-1} \left(\frac{k_{j-1}}{h_i} + \frac{h_i}{k_{j-1}} \right) \right) \\
& + \phi_{i+1,j} \left(5 \frac{k_j}{h_i} - \frac{h_i}{k_j} + 5 \frac{k_{j-1}}{h_i} - \frac{h_i}{k_{j-1}} \right) + \phi_{i,j+1} \left(-\frac{k_j}{h_i} + 5 \frac{k_{j-1}}{k_j} - \frac{h_{i-1}}{k_j} + 5 \frac{h_{i-1}}{k_{j-1}} \right) + \phi_{i,j-1} \left(-\frac{k_{j-1}}{h_{i-1}} + 5 \frac{h_{i-1}}{k_{j-1}} - \frac{k_{j-1}}{h_i} \right) \\
& + 5 \frac{h_i}{k_{j-1}} \left(\phi_{i-1,j} + \phi_{i-1,j} \left(5 \frac{k_j}{h_{i-1}} - \frac{h_{i-1}}{k_j} + 5 \frac{k_{j-1}}{h_{i-1}} - \frac{h_{i-1}}{k_{j-1}} \right) - 5 \phi_{ij} \left(\frac{k_j}{h_i} + \frac{h_i}{k_j} + \frac{k_{j-1}}{h_i} + \frac{h_{i-1}}{k_{j-1}} + \frac{k_j}{h_{i-1}} + \frac{h_{i-1}}{k_j} \right) \right) \\
& - \frac{\lambda^2}{144} \left(h_i k_j \phi_{i+1,j+1} + h_{i-1}^k k_j \phi_{i-1,j+1} + h_{i-1}^k k_{j-1} \phi_{i-1,j-1} + h_i^k k_{j-1} \phi_{i+1,j-1} + 5 k_j (h_i + h_{i-1}) \phi_{i,j+1} \right. \\
& \left. + 5 h_i (k_j + k_{j+1}) \phi_{i+1,j} + 5 h_{i-1} (k_j + k_{j-1}) \phi_{i-1,j} + 5 k_{j-1} (h_i + h_{i-1}) \phi_{i,j-1} + 25 \{ h_i (k + k_{j-1}) + h_{i-1} (k + k_{j-1}) \} \phi_{ij} \right) \\
& = \frac{1}{144} \left(h_i k_j^f \phi_{i+1,j+1} + h_{i-1}^k k_j^f \phi_{i-1,j+1} + h_{i-1}^k k_{j-1}^f \phi_{i-1,j-1} + h_i^k k_{j-1}^f \phi_{i+1,j-1} + 5 k_j (h_i + h_{i-1})^f \phi_{i,j+1} \right. \\
& \left. + 5 h_i (k + k_{j-1})^f \phi_{i+1,j} + 5 h_{i-1} (k + k_{j-1})^f \phi_{i-1,j} + 5 k_{j-1} (h_i + h_{i-1})^f \phi_{i,j-1} + 25 (h_i + h_{i-1}) (k + k_{j-1})^f \phi_{ij} \right)
\end{aligned}$$

Equation (4.28)

In order to impose the second boundary condition, Eqn. (2.33), the integrand in Eqn. (2.33) is expanded using the basis functions. At the northern boundary, this gives:

$$\left. \frac{\partial \phi}{\partial y} \right|_{i,1} = \frac{\phi_{i,2} - \phi_{i,1}}{k_1} \quad (4.31)$$

and at the southern boundary, it yields:

$$\left. \frac{\partial \phi}{\partial y} \right|_{i,NJ} = \frac{\phi_{i,NJ} - \phi_{i,NJ-1}}{k_{NJ-1}} \quad (4.32)$$

for $i = 1, 2, \dots, NI$. In Eqns. (4.31) and (4.32), $\phi = \partial \psi / \partial t$.

Eqn. (2.33) is of the form:

$$\int_0^{L_x} g \, dx = 0 \quad (4.33)$$

Expanding g with the bi-linear basis functions and integrating Eqn. (4.33), one obtains:

$$(h_{NI-1} + h_1)g_{1,j} + (h_1 + h_2)g_{2,j} + \dots + (h_{NI-2} + h_{NI-1})g_{NI-1,j} = 0 \quad (4.34)$$

Substituting Eqns. (4.29) - (4.31) into Eqn. (4.34) and noting that in Eqn. (2.34), $g = \partial \phi / \partial y$, one obtains:

$$\phi_{B,1} = \frac{1}{2L_x} (\phi_{1,2}(h_{NI-1}+h_1) + \phi_{2,2}(h_1+h_2) + \dots + \phi_{NI-1,2}(h_{NI-2}+h_{NI-1})) \quad (4.35)$$

where $\phi_{B,1}$ is the value of ϕ at all nodal points along the northern boundary. Similarly, the expression for the nodal values on the southern boundary, $\phi_{B,NJ}$, is found to be given by:

$$\phi_{B,NJ} = \frac{1}{2L_x} (\phi_{1,NJ-1}(h_{NI-1}+h_1) + \phi_{2,NJ-1}(h_1+h_2) + \dots + \phi_{NI-1,NJ-1}(h_{NI-2}+h_{NI-1})) \quad (4.36)$$

If the vectors $\underline{\phi}$ and \underline{f} are defined to be vectors of all (both interior and boundary) nodal values of their respective functions, Eqns. (4.35) and (4.36) may be combined with all equations similar to Eqn. (4.28) to obtain a complete system of linear equations. This may be written in matrix form as:

$$A\underline{\phi} = B\underline{f} \quad (4.37)$$

An approximate solution to the Helmholtz equation, Eqn. (4.22), with the boundary conditions given by Eqns. (2.32) and (2.33),

may be found by solving the matrix equation, Eqn. (4.37). The procedure for doing this is described in Section 4.4.

4.3 The Energy and Potential Enstrophy Relations

The evaluation of the kinetic energy, the available potential energy, and the potential enstrophy for the finite-element solution is done by approximating Eqns. (2.36), (2.38) and (2.43) with equations which are numerically integrated. These three equations have the form:

$$Z = \alpha \int \int_D f(x,y) \, dx dy \quad (4.38)$$

where α is a constant and z represents the KE, PE or APE. The nodal values of the integrand, $f(x,y)$, are evaluated using the results of Section 4.2. This is done using Eqns. (2.36), (2.38) and (2.43) to define the integrand for the KE, APE and PE respectively. In order to perform the integration, Eqn. (4.38) is approximated by a simple finite-difference representation:

$$Z \approx \alpha \sum_i \sum_j f_{ij} \left(\frac{h_i + h_{i-1}}{2} \right) \left(\frac{k_j + k_{j-1}}{2} \right) \quad (4.39)$$

where the f_{ij} are the nodal values of the integrand, $f(x,y)$.

The limits of the summations in Eqn. (4.39) must be defined. In the definitions of KE, APE and PE in Section 2.3, we considered the integration to occur over the entire domain. Thus, the summation is with $i = 1, 2, \dots, NI - 1$ and $j = 1, 2, \dots, NJ$. Within these limits, the conservative properties of the solution by the FEM are checked and the results are presented in Chapter 5.

It is also of interest to calculate the KE, APE and PE in the region where the solution by the FEM has the greatest accuracy, i.e. the sub-domain with uniform resolution. This will enable a better judgement of the potential accuracy of the solution using the FEM. For the cases presented in Chapter 5, this is done for the area marked C in Fig. 4.1. These calculations are done for the spectral solution as well so the two solutions can be compared. To do this, it was necessary to calculate the nodal values of the spectral solution.

4.4 The Solution Algorithm

The five steps for the solution of the model equations (Eqns. (2.26), (2.29) and (2.30)) using the FEM are:

- 1) to determine the right-hand-sides of the model equations.
- 2) to solve the system of equations resulting from the discretization of Eqns. (2.26) and (2.29) for $\bar{\psi}/\partial t$ and $\hat{\psi}/\partial t$, respectively, at time t .
- 3) to extrapolate in time to estimate $\bar{\psi}$ and $\hat{\psi}$ at time $t + \Delta t$, where Δt is the time step.

- 4) to solve the system of equations resulting from the discretization of Eqn. (2.30) for $\tilde{\omega}_2$ at $t + \Delta t$.
- 5) to repeat steps one through four for the desired number of time steps.

The procedure used for performing the first step was described in Section 4.3. The solution of a system of linear equations, as required by steps 2 and 4 may be done using either direct or iterative methods. Tong and Rossettos (1977) provide a discussion of the advantages and disadvantages of each method when applied to finite-element problems. Their analysis indicates that, for the systems of equations in this thesis, a direct method would probably be most efficient in the use of computer resources. However, as the computer code was to be used a limited number of times, the efficient use of computer resources was judged to be less important than the efficient use of programming time. Hence, an iterative method is used.

Young (1971) describes many iterative methods and from these, successive over-relaxation (SOR) is chosen as it is relatively simple to program and is reasonably efficient in the use of computer resources. To apply SOR, the matrix A in Eqn. (4.37) is split into a lower triangular, L , and an upper triangular, U , matrix. Then, given an initial guess for the nodal values, $\underline{\phi}^0$, the approximate solution, $\underline{\phi}^n$, is found using:

$$L\Delta\underline{\phi} = \underline{Bf} - A\underline{\phi}^{n-1} \quad (4.40)$$

$$\underline{\phi}^n = \underline{\phi}^{n-1} + \omega \Delta \underline{\phi} \quad (4.41)$$

where ω is the over-relaxation factor. The solution proceeds by iterating with Eqn. (4.41) and (4.42) and the superscript n refers to the number of the iteration. One continues iterating until the difference between successive iterates, $\Delta \underline{\phi}$, is sufficiently small. In the present thesis, iterations are continued until the relative change from one iteration to the next at all nodes in the high-resolution sub-domain (Area C) is less than .1%.

The total number of iterations needed for a solution depends critically on the choice of the value for the over-relaxation factor. An attempt was made, using techniques described by Young (1971), to objectively determine the optimum over-relaxation factor, i.e. the one which resulted in the minimum number of iterations for a solution. This did not work well, however, and a factor of 1.795 was found, by trial and error, to be nearly optimum.

Step 3 was accomplished using the Euler and Adams-Bashforth methods described in Section 3.2. The length of the time step, Δt , for the solution by the FEM was chosen using different criteria than in Section 3.2 however. Haltiner (1971) shows that to prevent computational instability, the fastest travelling wave in the solution must move less than one grid interval, Δx , in one time step, Δt . This leads to the following condition:

$$\frac{C\Delta t}{\Delta x} \leq \frac{1}{\sqrt{2}} \quad (4.42)$$

for a two-dimensional domain where C is the speed of the fastest travelling wave. This criterion was derived using finite-difference methods rather than finite-element methods for the discretization of the space coordinates. However, the physical interpretation of this criterion presented by Haltiner (1971) suggests that it could at least be used as a guideline for the present work. For synoptic-scale systems, in quasi-geostrophic models where sound and gravity waves are not permitted, Holton (1972) estimates 50 m/s to be the maximum value of C . With this value and a time step of 1/2 hour, Eqn. (4.42) indicates that computational instability should be prevented in the present model for all portions of the grid with grid spacing greater than 127 km. Thus, only in a very narrow band, near the north and south boundaries of the present grid, is the possibility of computational instability indicated. A time step of 1/2 hour is used as it will mean that the time truncation errors in both methods of solution are the same and because the criterion have suggested that computational instability is unlikely to be a major problem.

CHAPTER 5

THE RESULTS

5.1 Introduction

The model equations derived in Chapter 2 are solved using both the spectral and the finite-element methods, as discussed in Chapters 3 and 4. In this chapter, some examples of the solutions obtained using these methods, with various initial conditions, will be presented. In the accompanying discussion, some of the strengths and weaknesses of the finite-element solution will be demonstrated through a comparison with the spectral solution. Prior to presenting these results, a discussion of some of the conventions used and the numerical values of certain constants will be given.

A total of four sets of initial conditions will be presented with both methods of solution used for all four. Cases I, II and III are model atmospheres which, initially, favour the conversion of APE to KE. This conversion leads to a development (amplification) of the wave in the mean height field. The cases will be referred to as weakly, moderately, and strongly developing cases. Case IV is a model atmosphere in which a conversion of KE to APE takes place. The wave in the mean height field decays and this is called the decaying case.

As shown in Chapter 3, the spectral solution allows only waves of wavenumber n in the x -direction to exist. This wavenumber must be chosen and its value determines the scale of motion to be consi-

dered. As mid-latitude synoptic scale motions are to be studied, a wavenumber of 3 to 9 is appropriate. For the chosen channel length, these wavenumbers lead to waves with wavelengths of the order of thousands of kilometers. Holton (1972) suggests that wavenumber 7 is close to the average wavenumber of mid-latitude synoptic systems. He also shows that this wave is near the wavelength of maximum baroclinic instability, i.e. the wavelength which becomes baroclinically unstable with the lowest thermal contrast. A wavenumber of 7 is used in all the cases in this chapter.

The static stability parameter, σ , is defined by Eqn. (2.16). It is calculated by writing Eqn. (2.16) in finite-difference form using a central difference about 500 mb. Tabulated values of θ and p for the ICAO standard atmosphere were used to find $\sigma = 2.8 \times 10^{-6} \text{ m}^4 \text{ s}^2 \text{ kg}^{-2}$. This is within 25% of the value assumed by Holton (1972) for mid-latitude synoptic scale systems.

Teweles and Wobus (1954) developed the S1 score for comparing two fields with their values determined at grid points. This score is used later in this chapter as an aid in comparing the solutions by the two methods. The S1 score compares the gradients of the fields rather than their magnitudes, for example. Such a comparison is significant meteorologically because the gradient in the stream function determines the wind speed and direction, through the geostrophic wind relation, Eqn. (2.7). The wind and the vertical wind shear are more important weather elements than the actual value of the stream function

at a point, because they determine the potential for baroclinic and barotropic growth of weather systems.

To calculate the S1 score, one first calculates the gradients in the two fields which are being compared. In the present work, the components of the gradients of the two fields in the x and y directions are calculated for all points in the verifying area (area C of Fig. 4.1) including the boundaries of that area. If e_G is defined to be the difference between the gradients of the two fields at a point and G_L is defined to be the larger of the two gradients, the S1 score is calculated using the formula:

$$S1 = 100 \frac{\sum |e_G|}{\sum |G_L|} \quad (5.1)$$

where the summation occurs over all points in the verifying area and both components of the gradients are compared at each point. The S1 score varies from 0 to 100 with lower values indicating a greater agreement of the two fields.

The mean difference (MD) and mean absolute difference (MAD) between the finite-element and spectral solutions are also calculated over the verification area. In both the MD and the MAD, the finite-element solution is subtracted from the spectral solution at all grid points within area C, including the boundaries. The mean used is the arithmetic mean.

During initial test runs, it was found that waves of short wavelength were being generated along the northern and southern boundaries when the finite-element solution was computed. These waves grew in amplitude and moved away from the boundary, gradually contaminating the solution by the FEM even in the high-resolution sub-domain. Investigations indicated that a major cause of these waves was the poor evaluation of the vorticity near the boundary. In particular, the evaluation of the second derivative of the stream function in the y-direction was subject to a large error near the boundary. A simple 3-point smoothing algorithm was introduced to damp these waves. Thus, if the function $f(x)$ is known at the grid points (x_1, x_2, \dots, x_N) , the smoothed function, $g(x)$, at the point x_i is given by:

$$g(x_i) = (1 - S(x_i))f(x_i) + \frac{S(x_i)}{2} (f(x_{i+1}) + f(x_{i-1})) \quad (5.2)$$

where $S(x_i)$ is the smoothing parameter at x_i . This smoother was applied to the calculated vorticity field in the vicinity of the northern and southern boundaries. It was applied to the vorticity field each time it was evaluated, first in the y-direction and then in the x-direction. Although most smoothing was necessary in the y-direction, some smoothing in the x-direction was found to be helpful in controlling the spurious waves near the boundary. For the six grid points nearest the boundary, the smoothing parameters were .30, .50, .30, .15, .08, and .02. The largest values were near the northern and southern boundaries. The smoothing parameter was zero for all other interior grid points. The smoothing parameters for the first six interior lines of

grid points in the x-direction, adjacent to and including the northern and southern boundaries, were .10, .15, .10, .05, .01 and .005.

The values for the smoothing parameters were determined by trial and error. This set of parameters seemed to have little effect on the long wavelengths of interest but controlled the spurious short wavelengths reasonably well. It was found that using this smoothing also reduced the amount of computation time by 10 - 15% as it allowed more rapid convergence to occur in the solution of the boundary value problems by relaxation.

In displaying the height and thickness fields later in this chapter, the stream functions for the mean height and thickness fields have been converted to a height in meters. This was easily done as, in this model, the gravitational constant, g , is assumed to be independent of height. Following Holton (1972), the height in meters, z , is obtained using:

$$z = \psi \frac{f_o}{g} \quad (5.3)$$

where $g = 9.806 \text{ m s}^{-2}$. The stream function, ψ , is either $\bar{\psi}$ or $\hat{\psi}$, depending on which field is being converted.

The omega fields displayed have the units $\mu\text{bar/sec}$. This is the unit which is commonly used as it gives values of the order of 1. It should be noted, however, that negative omega values imply ascent.

In the display of these fields, roughly two-thirds of the total domain in Fig. 4.1 is shown. Referring to Fig. 5.1, for example, the lower panel is area A of Fig. 4.1 while area B of Fig. 4.1 is the upper panel. With this display, both the high and low resolution portions of the two solutions may be easily compared.

As previously stated, area C of Fig. 4.1 is the verification area. This is the sub-domain over which the various energy quantities, the PE, the SI score, the MD and the MAD are calculated, as shown, for example, in Figs. 5.10 - 5.12. The KE, APE and PE are also calculated over the whole domain for some of the cases.

5.2 Case I - Weak Development

Case I is an example of a weakly developing situation. Figs. 5.1 and 5.2 show that, initially, the trough in the thickness field lags (i.e. is west of) the trough in the mean height field by 90° . When the trough in the thickness field lags the trough in the height field by between 0° and 180° , the perturbations (i.e. the waves) in the mean height field will amplify. In this process, the flow becomes more meridional and the amplitude of the wave increases. Holton (1972) shows that this amplification will occur most rapidly when the lag is 90° . Holton (1972) did a linear analysis of an atmospheric model similar to the one under consideration. He found that if the gradients in the thickness field were less than a critical value, the waves would not develop, i.e. they were stable waves. With the values used in the present study, his analysis suggests the wave should be

stable for this case. Due to model differences and non-linear interactions, the wave does, in fact, develop.

As was discussed in Section 2.1, this growth is due to the baroclinic conversion of energy from APE to KE through the lowering of the centre of mass of the atmosphere. This will be demonstrated more clearly when, later in this section, Fig. 5.10a is discussed.

The spectral solution (SP) after twenty-four hours of integration is given in Figs. 5.4 - 5.6. The wave in the height field has undergone the expected development and retrogressed 15° . The thickness field has developed slightly and progressed eastward 10° . Thus, the phase difference between the two fields has been reduced by 25° . The omega field has amplified somewhat during the period and has retrogressed 22° .

Figs. 5.1 - 5.2 are also the initial conditions used for the solution by the FEM. Figs. 5.7 - 5.9 are the resulting twenty-four hour forecasts for the height, thickness and omega fields, respectively. Only a very detailed comparison of the forecasts produced by the two methods reveals any differences between them over the high-resolution portion of the grid. Near the boundaries, it is apparent that the troughs and ridges in the height field of the finite-element solution (FE) have retrogressed slightly faster than those of the SP.

The diagnostic omega fields for the two solutions, Figs. 5.6 and 5.9, are considerably different, however. The maxima in the omega field are larger in the FE and are closer to the southern boundary in the FE than in the SP. The minima of the FE behave in an analogous manner, i.e. they are more negative and are closer to the northern boundary than are those of the SP. Whereas the SP has a value very near 0.0 at the boundary, the boundary value of the FE is somewhat larger than 0.0, in absolute magnitude. Investigations revealed that this remarkable difference in the two omega fields was present even at the initial time. Thus, this difference is not due to the development and growth of errors during the integration but rather due to the evaluation of the diagnostic omega field of the FE. Using the boundary condition, Eqn. (2.32), one may readily show that the right-hand-side of Eqn. (2.30) should be zero on the boundary, for the chosen initial conditions. However, since the estimate of the vorticity, \bar{Q} , is very poor on the boundary, the term $J(\hat{\psi}, \bar{Q})$ is subject to a large error on the boundary. When the boundary value of the omega field is in error, the interior values of the omega field are also in error. The poor estimation of the vorticity near the boundary also causes an error in the evaluation of the right-hand-sides of Eqns. (2.26) and (2.29). However, this error is not as easily seen in these cases as one is dealing with an error in the forecast of a relatively small change (i.e. $\partial \bar{\psi} / \partial t$) in a large quantity (i.e. $\bar{\psi}$).

The evaluation of the boundary condition, Eqn. (2.33), also makes a contribution to the errors near the boundaries. This boundary condition requires an integration along the boundary. Thus, values

from both the high and low resolution portions of the grid are used to find the boundary condition for the entire boundary. This effectively increases the truncation error in the x-direction, along the entire boundary, to that of the portion of the boundary with the lowest resolution.

The resolution in the y-direction is very high near the boundary. This band with high resolution along the boundaries was used to try to reduce the errors along the boundaries. With such resolution, a better estimate of the vorticity on and near the boundary can be obtained. In addition, the evaluation of the boundary condition, Eqn. (2.33), has greater accuracy as the truncation error present in the discretization of $\partial\psi/\partial y$ is reduced. Because of this area of high resolution, however, a shorter time-step must be used in the integration to prevent computational instability, as was discussed in Chapter 4.

Cullen (1976), Grammelvedt (1969) and others have used different initial conditions when performing numerical experiments with this model on this domain. Their initial conditions were characterized by fields with very little variation in the stream functions near the boundary. In effect, then, there was a buffer of 600 - 800 km. between the boundary and the region of interest. This approach was not used in the present work, however, because the highly accurate spectral solution would not have been available for comparison.

Fig. 5.10a shows the evolution of the TE, KE and APE in the verification area (area C of Fig. 4.1) for the two solutions during the

forecast period. The TE of the SP increases by approximately 30% during the period of integration. This is not, however, in violation of the conservation of TE, which was demonstrated in Chapter 3, as this TE has been calculated only in the verification area and not over the whole domain. As noted in Chapter 3, the conservation of the TE of the SP over the whole domain was verified for a wide variety of initial conditions. The increase of the TE of the SP in area C is due to the motion of the fields and the development of the fields during the forecast period.

Referring to Fig. 5.10a, we see that the gain in TE of the SP is due to a gain in KE while the APE has remained nearly constant. Thus, during the development of the waves, APE has been converted to KE but also, in the verification area, there has been energy imported. It is not possible to say what proportion of the gain in KE is due to importation and what proportion is due to development.

Fig. 5.10a confirms that, with respect to the three energy quantities, the FE has performed well. The TE of the FE has decreased relative to the SP in the verification area. The FE has lost approximately 2% of the TE in the verification area. This small error is seen to be due to a loss in the APE of the FE relative to the SP, while it is noted that the KE of the FE is extremely close to that of the SP. The growth rate of the error seems to be nearly linear, with the error first being noticeable after approximately ten hours of integration.

Fig. 5.10b indicates that the FE has also lost a small amount of PE relative to the SP. The total loss is only .4%. Careful scrutiny of Fig. 5.10b reveals that the FE actually gained PE relative to the SP for the first twelve hours and then lost PE for the last 12 hours. The small variation of the PE of the FE about that of the SP indicates that the FE behaves quite well in this case.

In Fig. 5.11a, the KE of the two solutions is presented in the zonal and eddy forms, while Fig. 5.11b gives the corresponding values for the APE. As noted earlier, the KE of the two solutions is essentially the same. However, the FE has lost ZKE relative to the SP and gained an equal and opposite amount of EKE. There is roughly a 1% difference in the ZKE between the two solutions. The loss in the APE of the FE relative to the SP is seen to be due to a loss of ZAPE while the EAPE of the two solutions is essentially the same.

The reasons for the differences between the two solutions, as shown in Figs. 5.10 and 5.11, are rather hard to determine precisely. As mentioned previously, there are errors in the FE due to space and time discretization. Also, it has been shown previously that, as a result of space discretization, significant errors can occur near the boundary due to the imperfect implementation of the boundary condition. As the time discretization scheme is the same for both models, it is expected that this, by itself, makes only a very small contribution to the differences between the two solutions. It seems probable, however, that the loss in ZAPE of the FE relative to the SP is due to the poor evaluation of the boundary condition since the boundary value plays a

relatively larger role in the determination of the zonal quantities than of the eddy quantities. The error in the ZKE of the FE is thought to be smaller because the boundary condition on the height field is satisfied better than the one for the thickness field. In this model, the boundary value of the height field should not change with time but that of the thickness field does. It was found that the forecast change in the thickness field on the boundary in the FE was subject to a large error but that the FE maintained the boundary value of the height field fairly well. It is to be noted that the thickness field determines the APE whereas the KE is determined by both the thickness and height fields. The small errors in the eddy forms of the energy quantities are thought to be due to a combination of the various errors. It is probable that, during the initial few hours, errors due to space discretization are most important. As the integration proceeds, errors from the boundary will gradually infiltrate the high-resolution portion of the domain and become more important.

Fig. 5.12a gives the S1 scores for the FE relative to the SP. There is nearly a linear growth of the S1 score for both the height and thickness fields with the S1 score for the height field growing more rapidly. In the development of this score, Teweles and Wobus (1954) indicated that, in the forecasting of the real atmosphere, an S1 score of under 30 could be considered a perfect forecast while an S1 score of over 70 would indicate a useless forecast. For the present integration, the S1 score has only grown to near 5 which indicates, according to Teweles and Wobus's subjective criteria, an exceptional forecast.

However, care must be exercised here. In the present study, a highly simplified model atmosphere is used and only a very limited number of types of waves are allowed, in contrast to many types of wave motion possible in the real atmosphere. Also, in the present study, the initial conditions are very well specified and this is not true of forecasts for the real atmosphere where the initial conditions must be specified by some objective analysis of actual reported data. Thus, it seems reasonable to assume that the S1 scores for the present study should be significantly lower than what might be expected of forecasts of the real atmosphere. It is difficult, however, to objectively determine how small the score should be to conclude that a "good" forecast has been generated. The differences which were seen between the two forecasts were very small and this, at least, gives one the impression that the FE was very close to the SP and that, therefore, S1 scores of about 5 or less indicate a good forecast. With the presentation of the remaining cases, an approximate upper limit for the S1 score representing a good forecast may be found.

Fig. 5.12b shows that, in the mean over the verification area, the height field of the SP is 0.24 m higher than that of the FE and that the thickness of the SP is 0.08 m lower than that of the FE. Considering that the fields under consideration have magnitudes of nearly 100 m and that some points in the height field underwent changes of nearly 30 m during the forecast period, these MD errors are quite respectable. The MD for the height field is larger than that for the thickness field because there were larger changes in the height field during the forecast period.

Fig. 5.12b also gives the MAD between the two solutions in the verification area. The MAD for both fields grows to nearly 1.10 m during the period of integration. The MAD for the thickness field shows nearly linear growth while that for the height field appears to be closer to exponential growth.

It is interesting to note that while the MD for the thickness field was smaller, in absolute value, than that for the height field, the MAD for the thickness field is larger than that for the height field. Thus, the errors in the thickness field are smaller than those of the height field but there is a greater bias in the errors of the thickness field than in the height field. In effect, these results suggest that although the error at individual points is, in general, smaller in the thickness field than in the height field, the mean of the thickness field is not being kept constant as well as is that of the height field. It is not possible to suggest the precise cause of this at this time.

An interesting comparison may be made between Figs. 5.12 and 5.11. Fig. 5.11 shows that the KE of the FE is closer to that of the SP than is the APE. Fig. 5.12a shows that the gradient of the thickness field is forecast well while in Fig. 5.12b, we have seen that the forecast magnitude of the thickness field has a bias in it. The KE depends on the gradient of the thickness field while the APE depends on the magnitude of the thickness field. Given that the ZAPE is in error by the largest amount, it is suggested that the errors in the forecast magnitudes of the thickness field are produced largely by a poor implementation of the boundary condition.

5.3 Case II - Moderate Development

Case II is a situation favourable for greater development than Case I. Figs. 5.13 and 5.14 show that, initially, the trough in the thickness field lags the trough in the mean height field by 90° . The thickness field for this case is very similar to that of Case I, but the wave in the height field is characterized by a much larger amplitude for Case II than for Case I. As a result, the amplitude of the omega field is larger for this case, as shown in Fig. 5.15.

With the larger meridional gradient in the height field in this case, a large mean zonal wind exists in the domain. This means that the wave in the mean height field is expected to move faster than in Case I. Also, the larger meridional variation in the mean height field implies greater thermal advection in this case. This increases the rate of baroclinic development. Thus, both the rate of propagation and rate of development of the wave in the mean height field are expected to have been changed in Case II.

Figs. 5.16 - 5.18 give the mean height, thickness and omega fields, respectively, after 48 hours of integration with the spectral method. (The length of integration has been increased as this will provide a more stringent test on the conservative properties and of the accuracy of the FE. This will also allow a better investigation of the possibility of computational instability in the FE.) The wave in the mean height field has developed and progressed eastward approximately 15° . The wave in the thickness field has decayed and progressed

eastward approximately 45° . Thus, the phase difference between the waves in the two fields has been reduced to 60° . The wave in the omega field has decayed slightly and progressed approximately 20° .

Figs. 5.19 - 5.21 show the mean height, thickness and omega fields, respectively, for a 48 hour forecast using the FEM. The differences between the two solutions are much more apparent in this case than in Case I. Referring to Figs. 5.16 and 5.19, it is apparent that the FE has not maintained the north-south anti-symmetry present in the SP. The troughs in the FE have undergone greater development than those of the SP while the ridges of the FE have not been built as high as those of the SP. This seems to be due to problems with the boundary values of the FE. The SP has maintained a constant value for the mean height field along the boundaries while the boundary value of the FE has changed. Although any of the errors previously discussed could be the cause of this, it is probable the main cause is the poor evaluation of the vorticity near the boundary. This, as discussed in Sec. 5.2, was found to be a major source of error near the boundary.

Related to the above problem is the curvature in the trough and ridge axes of the FE. The SP has trough and ridge axes which are north-south while those of the FE are slightly curved. It appears that the speed of these waves in the FE is slower than those of the SP near the boundaries but is faster than those of the SP in the central portion of the domain. These errors are due, either directly or indirectly, to the problems near the boundaries. The overdevelopment of the troughs of the FE has led to larger zonal wind speeds in the trough in the

middle of the domain. This leads to a more rapid progression of the trough axis in the central portion of the domain. Similarly, larger gradients in the height field near the ridge axes in the centre of the domain has caused a more rapid progression of the ridge axes. Near the boundaries, just the opposite has happened. Gradients have been reduced and the progress of the waves has been retarded.

Finally, the presence of a few small amplitude short wavelength waves in the height field of the FE near the boundaries should be noted. These are the waves which the smoother, discussed in Sec. 5.1, was designed to control. Also of note is the absence of these waves in the region of the domain where the grid of the FE has very poor resolution. The short wavelength waves could not be resolved in this portion of the domain.

Comparing Figs. 5.17 and 5.20, it is apparent that there are vast differences between the two solutions for the thickness field. The FE has very nearly maintained the correct value for the maximum value of the highs and the minimum value of the lows. However, the maxima and minima of the FE are much closer to the boundaries than are those of the SP. This has been caused by the problems with the boundary value of the FE. For this case, the boundary value of the SP has changed very little during the integration period while the boundary value of the FE has definitely changed during this period.

In the central portion of the grid, the FE is within a few percent of the SP however. The zero thickness isopleth is forecast very well by the FE near the centre of the grid even in those regions with relatively poor resolution in the x-direction. The phase speeds in the central portion are nearly identical as is the meridional amplitude of the zero isopleth.

The short wavelength waves are more noticeable in the thickness field of the FE than they are in the height field. Their amplitudes are small however. The waves are not distinct enough in either field to judge whether these short wavelength waves are in a position favourable for development or not.

The omega fields of the two solutions, shown in Figs. 5.18 and 5.21, are drastically different. The trough and ridge axes of the FE have a distinct curvature to them whereas those of the SP are straight lines oriented north-south. This curvature is simply a reflection of the curvature of the axes in the height and thickness fields of the FE found previously. The maxima/minima of the FE are more positive/negative than are those of the SP. This is indicative of the greater development which the FE has undergone. Also, the maxima and minima of the FE are located nearer to the boundaries than are those of the SP.

There are numerous spurious short wavelength waves present in the omega field of the FE. Although they are most common and have their largest amplitudes near the boundaries, they are also present well

into the domain. These waves in the interior of the domain were not visible in the height and thickness fields but they must have been there or they would not be present in the omega field. Thus, some small amplitude short wavelength waves have propagated well into the grid.

Fig. 5.22a shows the TE, APE, and KE in the verification area for this case. The TE of the SP rises during the integration period for the reasons discussed with Case I. However, the TE of the FE rises faster than that of the SP. After 48 hours, the TE of the FE is approximately 6% larger than that of the SP. This increase in TE is due largely to the increased gradients in the mean height field of the FE which were seen earlier. The gain in the TE confirms that the wave in the mean height field of the FE has developed more than the wave in the SP. Also, it should be noted that the divergence in the TE of the two solutions does not begin until after twenty-four hours.

Further study of Fig. 5.22a reveals that the gain in the TE of the FE relative to the SP is due to a large gain in the KE of the FE relative to the SP which is only partially compensated for by a loss in the APE of the FE relative to the SP. The divergence of the KE and APE of the two solutions begins after only twelve hours of integration. During the period from twelve to twenty-four hours, the KE and APE of the FE, although in error with respect to the SP, adjust themselves so that the TE of the two solutions is the same. Thus, during the first twelve hours, the two solutions behave similarly. During the twelve to

twenty-four hour period, the FE converts APE to KE at a faster rate than does the SP. During the final 24 hours, the FE either imports energy into the verification area or it has created some spurious source of energy. It seems probable that there has been importation of energy. In the FE relatively large amounts of KE have been concentrated in the middle portion of the domain while the KE near the boundaries is lower in the FE than the SP. This can be seen by comparing the gradients in the mean height field of the two solutions. A calculation of the meridional momentum transport might have given further insight here but this was not thought of until after the computation was finished.

Fig. 5.22b shows that the PE of the FE is less than that of the SP during the forecast period. The PE in the verification area changes significantly during the period and the FE under-forecasts the change in PE by nearly 20%. This result is a little surprising as, with the greater development of the FE, one would expect larger vorticities in the FE than in the SP. However, the troughs and ridges are sharper in the SP than in the FE and, thus, the vorticity of the FE is lower. Also, the thickness field of the FE has smaller values over the verification area than does the SP. These effects both act to give the FE a lower PE than the SP.

In Fig. 5.23a, it is seen that the FE gains a significant amount of ZKE relative to the SP while losing a small amount of EKE relative to the SP during the forecast period in the verification area.

The difference appears in the ZKE after fifteen hours and in the EKE after twenty-four hours. The aforementioned meridional transport of east-west momentum towards the centre of the domain in the FE has presumably led to this increase in the ZKE in the verification area. It seems that EKE in the FE has been converted to ZKE as well. This is reflected in the relatively broad troughs in the FE. This, I believe, is a natural result of the process of baroclinic development. During development, ZAPE is converted to EAPE; EAPE is converted to EKE; and, finally, EKE is converted to ZKE, as discussed by Holton (1972). The wave in the FE is at a more advanced state of development than the wave in the SP. Physically, the wave in the FE has presumably begun to slow down its rate of growth and soon, it is expected the wave will undergo the process of occlusion and begin to decay.

Fig. 5.23b shows that, in fact, the APE of the two solutions is not the same during the initial twelve hours as suggested when Fig. 5.22a was discussed. The resolution of that figure was not high enough to see that the FE continuously loses APE relative to the SP. It is important to note, however, that in this case, the APE is much smaller than the KE. Thus, for this case, a 10% error in the APE has the same effect on the TE as approximately a 1% error in the KE, i.e. the error in the APE is a less important measure of the performance of the FE than is the KE.

The loss of APE in the FE relative to the SP is seen, in Fig. 5.23b, to be due to a loss of both ZAPE and EAPE. During development, there is a conversion of ZAPE to EAPE. The FE loses ZAPE slightly

faster initially than the SP does but the difference in the ZAPE between the two solutions remains relatively small. However, after twelve hours, the FE begins to lose EAPE rapidly with respect to the SP. It seems that energy in the FE, during the last thirty-six hours of development, is being rapidly transferred from ZAPE through the EAPE and EKE forms to ZKE, i.e. the FE has an accelerated transfer of energy during the baroclinic development. The result is a loss of ZAPE, EAPE and EKE in the FE relative to the SP with only the ZKE of the FE larger than that of the SP. The error in the ZKE begins to grow most rapidly a few hours after the error in the EAPE begins its rapid growth. This suggests that there is a causal relationship between the two errors. The relationship is somewhat clouded by the possibility of either net energy importation to or exportation from the verification area.

Fig. 5.24a shows that the S1 score for the thickness field has grown more rapidly than for the height field. The growth of the S1 score for both fields is approximately twice as fast as for Case I. The faster growth of the S1 scores in Case II may be due either to the faster growth or to the greater motion in Case II. In both cases, the field which underwent the greatest change had the larger S1 score. The final S1 score for the height field is near 18. In the preceding discussion of the various energy quantities, it was found that errors began to grow rapidly after twenty-four hours. After thirty hours, large differences between the two solutions were common in the energies. This suggests that an S1 score of less than ten or twelve indicates a "good" forecast. This is a subjective criterion and further discussion will be

presented with the final cases.

Fig. 5.24b gives the MD and MAD curves for the two models for this case. Most notable is the large growth of the MAD between the mean height fields. This is indicative of the greater development present in the FE. The MD between the mean height fields grows to approximately 7 meters. This indicates that the areas of low height have developed faster than the highs have built, in agreement with what was observed earlier. In contrast, the MD between the thickness fields stays relatively small and negative. The MAD between the thickness fields rises to almost 7 m. These results indicate that although there are large errors in the thickness field of the FE, there is very little bias towards either raising or lowering the mean of the field in the verification area.

In Fig. 5.25a, the KE, APE and TE of both solutions over the entire domain are presented. The TE of the SP changes by approximately 7% during the forecast period. The TE of the SP should be conserved and this error is due to the method used for calculating the TE here. For this figure, the values of the SP at the grid points were found and a numerical integration was performed, as described in Section 4.3. Thus, this error is due to the discretization of the SP and the truncation errors in the numerical integration. In fact, as discussed in Section 3.3, it was confirmed that the SP conserved TE to better than 1 in 10^4 . During the forecast period, the TE of the FE changes by less than 3%. Considering that the SP was subject to a 7% error due to the discretization process, the FE has conserved TE quite well. The maximum

difference between the TE of the two solutions is only 3%. The KE of the two solutions is essentially identical but the FE does gain APE relative to the SP. This is probably due to the problems with the boundary value of the thickness field. During the integration period, this has caused a gain of 3% in the TE for the entire domain.

Fig. 5.25b gives the PE of the two solutions for the entire domain. Once again, there is a small error in the PE of the SP due to the numerical integration process. More important, however, is the striking gain in the PE of the FE after twenty-four hours. This large amount of PE is found in the spurious short wavelength waves generated near the boundaries. These waves were found to have vorticities which were an order of magnitude larger than vorticities elsewhere in the domain. It is apparent that the smoother which was used was not adequate to control the development and growth of these waves. It is probably pure coincidence that the error in the PE of the FE begin to grow most rapidly at approximately the same time the forecast in the verification area was found to degrade rapidly. The PE of the FE was found to be smaller than that of the SP right to the end of the integration. This indicates there was little or no contamination of the forecast in the verification area by the short wavelength waves. However, the exponential growth of the PE indicates the potential for future problems if the integration had proceeded further in time.

5.4 Case III - Strong Development

Case III is a situation favorable for strong development. Figs. 5.26 - 5.28 give the initial height, thickness and omega fields,

respectively. The height and thickness fields are characterized by large gradients and the troughs in the thickness field lag those of the height field by 90° . The thermal advection, which is essential for baroclinic development, is considerably greater than in the previous cases. This will promote strong development. The initial omega field has relatively large values for synoptic-scale vertical motion, indicating the vigour of the situation under consideration.

It was intended to integrate this case for forty-eight hours but computational instability occurred in the FE and the integration had to be terminated after thirty-two hours. Fig. 5.29 shows the mean height field of the SP after twenty-four hours. The wave has developed and progressed eastward approximately 30° . Fig. 5.30 shows the thickness field for the SP after twenty-four hours. The wave has progressed approximately 50° . The meridional amplitude of the thickness field has increased which implies the conversion of ZAPE to EAPE during the forecast period. On Fig. 5.31, we see that the amplitude of the omega field has decreased by approximately one-fourth and that the wave has progressed approximately 40° .

Figs. 5.32 and 5.33 show the mean height and thickness fields of the FE after twenty-four hours. The differences in the mean height fields of the two solutions are similar to those present in Case II. The trough and ridge axes are tilted. There has been preferential development in the northern half of the domain and spurious short-wavelength waves are present near the boundaries again. The thickness

fields differ in many of the ways found in Case II. Once again, the high and low centres of the FE have migrated towards the boundaries.

Figs. 5.34 and 5.35 show the mean height and thickness fields of the FE after thirty hours. It is apparent that some of the short-wavelength waves, which were present along the boundaries, have grown explosively and overwhelmed the wave of interest, thereby ruining the forecast. This displays the effect of numerical computational instability. By thirty-three hours, the amplitude of these waves had grown so large that an overflow condition was obtained on the computer.

In Fig. 5.36a, we see that the FE has gained TE relative to the SP, as in Case II. As with Case II, the error is quite small after twenty-four hours of integration. The FE has, once again, gained KE relative to the SP but lost APE. After twenty-four hours, the FE in both Case II and Case III has nearly maintained the TE in the verification area through equal but opposite errors in the EKE and APE. However, the fields in Case III have undergone greater development and faster motion than those in Case II. Thus, it seems the growth of errors is not very dependent on either the rate of development or speed of the fields.

In Fig. 5.36b the PE of the two solutions is presented. The error in the PE of the FE is approximately 5%, which is nearly identical to the error in the PE in Case II after twenty-four hours.

In Fig. 5.37, it will be seen that the zonal and eddy forms of the KE and APE of the FE behave in a similar manner to those in Case II. However, the errors are larger than in Case II. For example, the error in the ZKE of the FE after twenty-four hours is only 7% in Case II but is over 30% in Case III. Thus, although the error in the TE of the two cases had similar errors, the individual elements making up the TE are subject to greater errors in this case. In Case II, it was suggested that the FE underwent an accelerated development relative to the SP. The results of this section indicate that this has occurred in this case also, and, that the rate of accelerated development depends, at least somewhat, on the rate of development.

Fig. 5.38 shows the S1 scores and the MD and MAD curves for this case. Using the subjective criteria suggested in Section 5.3, the S1 scores suggest the forecast was "good" for approximately sixteen hours. The MD and MAD curves exhibit similar events to those of Case II. The errors have grown faster in this case.

5.5 Case IV - Moderate Decay

In Case IV, the troughs in the thickness field lead the troughs in the mean height field by 90° , as the initial conditions given in Figs. 5.39 - 5.41 show. The initial amplitudes of the waves are the same as in Case II. With such a phase relationship between mean height and thickness, the wave in the mean height field decays and there is a conversion of KE to APE. This is the opposite of what happened in the previous cases.

Figs. 5.42 - 5.44 give the mean height, thickness and omega fields, respectively, for the SP after forty-eight hours of integration. The wave in the mean height field has decayed in amplitude and has retrogressed approximately 12° . The wave in the thickness field has also decayed in amplitude and has progressed 70° . At this time, the waves are almost 180° out of phase. They appear to be in the process of adjusting themselves so that their phase becomes favourable for development. The omega field appears to have changed very little in either amplitude or position during the forecast period.

Figs. 5.45 - 5.47 give the three fields for the FE after forty-eight hours. The mean height field of the FE and the SP are very similar over the high-resolution portion of the grid. In fact, near the centre line of the domain, the FE is barely distinguishable from the SP until the grid spacing in the x-direction approaches 800 km. However, short-wavelength waves are present near the boundaries and those near the north boundary have not decayed as rapidly as those near the south boundary. Comparing this with the results of previous cases, it indicates there is a predilection for development/decay near the north/south boundary. This is probably due to problems with either the calculation of various quantities (e.g. vorticities) near the boundary or due to the implementation of the boundary condition. It seems probable that this bias is related to the sign of the height field at the boundary, i.e. the positive/negative boundary wants to become more positive/negative. The thickness field exhibits a large amount of curvature in the trough and ridge axes in this case. The precise cause of this is

unknown. However, the major cause of change in the thickness field in this case was motion as opposed to decay. Thus, it seems reasonable to expect the source of the problem to be connected to the motion of the field.

The omega field of the FE has a larger amplitude than does that of the SP. This is partially due to the causes discussed in Section 5.1. A secondary cause is the increased thermal advection in the FE over the SP due to the errors in the height and thickness fields of the FE.

In Fig. 5.48a, it can be observed that the FE has gained TE relative to the SP during the forecast period in the verification area. This also occurred in Case II and is likely indicative of the bias with respect to development at the boundaries. The slow decrease in the TE of the SP simply indicates there is a net export of energy from the verification area, whereas a net import was found for the developing cases. The APE of the solutions is very nearly the same during the period but the FE has gained KE relative to the SP. This was the cause of the gain in the TE of the FE in the developing cases also. The APE changes very little during the period suggesting there is very little decay in the verification area, or that an importation of APE to the verification area has taken place.

Fig. 5.48b shows that the FE loses PE relative to the SP during the first thirty-six hours of integration but gains it back during

the last twelve hours. The maximum difference between the PE of the two solutions is less than 4%.

Fig. 5.49a shows that the gain of KE by the FE is due to a gain of ZKE. The EKE of the two solutions is very similar throughout the forecast period. The error in the ZKE of the FE occurs early in the forecast period and grows slowly during the period to a maximum of 14%.

According to Fig. 5.49b, the FE loses a small amount of APE relative to the SP during the forecast period. The error is very small during the first thirty hours of integration as there are roughly equal and opposite errors in the ZAPE and EAPE. However, the errors in the ZAPE and EAPE stay relatively small throughout the forecast period.

Fig. 5.50 gives the S1 scores and the MD and MAD curves for this case. The S1 score rises rapidly for the thickness field. This is further confirmation of the errors which evolved in the FE during the rapid movement of this field. The S1 score for the height field rises at a much slower rate. Near the end of the period, the S1 score for the height field begins to rise more rapidly. It seems possible that this is a reflection of the poor thickness forecast as the two fields are connected through the model equations.

The MAD curves for the height and thickness fields are nearly parallel as they rise. They rise to a maximum of less than 5 m, which

is quite respectable compared with the previous cases. Once again, the MD curves suggest the height field of the FE is subject to some effect which causes the mean height field of the FE to become progressively lower than that of the SP. As discussed in Sections 5.2 and 5.3, it is felt that this is due to problems at the boundary. This effect occurs early in the period and seems to have less importance in the thickness field forecast.

Fig. 5.51a gives the TE, KE and APE of both solutions over the entire domain. It is apparent that the FE has conserved TE very well. However, the FE does not convert as much KE to APE as does the SP. This is due to the problems with development and decay near the boundaries which was discussed earlier. Fig. 5.51b shows the PE for the two solutions over the entire domain. Once again, the PE of the FE rises rapidly during the forecast period, as a result of the short-wavelength waves generated near the boundaries.

CHAPTER 6

CONCLUSIONS

The finite-element method and the spectral method have been used to numerically integrate the equations describing a two-level quasi-geostrophic model of atmospheric flow on a β -plane. The solutions using the two numerical methods for a group of cases have been compared. The spectral solution, being believed to be highly accurate, was used to evaluate the performance of the finite-element solution.

It was found that the largest errors in the finite-element solution arose near the boundaries of the grid. In the interior of the grid, the finite-element solution compared very well with the spectral solution. Over that portion of the grid with high-resolution, only minor phase or amplitude differences were found between the two solutions. As the time of integration increased, the errors in the finite-element solution generated along the boundaries slowly infiltrated the interior of the grid.

Two major sources of the boundary errors were found. It was suggested that the major error source was the evaluation of derivatives perpendicular to the boundaries. In particular, higher-order derivatives were found to be subject to the greatest errors. Thus, vorticities were not well calculated by the finite-element method along the boundary. In an attempt to minimize this problem, a portion of the grid near the boundaries was chosen to have high accuracy in the direction perpendicu-

lar to the boundaries. Further steps could possibly be taken to minimize this source of error. For example, the vorticity on the edge could be "tied" to the first interior grid point, i.e. the vorticity on the edge could be set equal to the vorticity at the first interior grid point. The success or failure of this technique would depend on the system of equations which would result from this approximation. It was not attempted as it would have required a major rewriting of computer routines and there was no strong evidence that it would dramatically improve the results. It is, however, a technique deserving of future consideration.

The second source of error at the boundary was that due to the implementation of the boundary condition, Eqn. (2.33). This required an integration along the boundary and was subject to the truncation error in the numerical integration. The magnitude of this truncation error was determined by that portion of the boundary with the largest grid spacing in the x-direction. The effect of this error on the solution could be most effectively minimized by ensuring that the boundaries are well away from the area of interest and by, if possible, using a high-order numerical integration scheme. In this work, the first technique could not be used. The second technique was not fully studied as it seemed the poor evaluation of perpendicular derivatives at the boundary was of much greater importance.

The techniques for avoiding or preventing the numerical instability found in Case III are reasonably well known. The small grid

lengths in the y-direction combined with the relatively long time-step have led to this instability. Increasing the grid length and/or decreasing the time-step would prevent this. A different time integration scheme could also be used. In the problem under consideration, the elimination of the small grid spacings near the boundary would be most economical in terms of computer time. This could be implemented when, as discussed earlier, an improved method of evaluating the derivatives perpendicular to the boundaries was found.

The two methods of solution were used for only a limited number of cases. There are other atmospheric flows which would be interesting to study. For example, it would be interesting to compare the two solutions when a barotropic atmosphere (i.e. one in which the thickness field was zero everywhere) was considered.¹ In this case, the height field would simply translate, i.e. there could be no development. By studying cases with varying speeds of translation, one could study how well the finite-element solution deals with translation and seek to determine a relation between the growth of errors and the speed of translation. A companion set of cases to these would be those in which the height field is stationary (or nearly so) but developing. Then, one could seek a relation between the growth of errors and the rate of development. These two sets of cases would help to answer some of the questions raised in Chapter 5 where the effects of translation and development could not be separated.

¹. Note that the perturbation thickness field is implied here.

Although some difficulties have been experienced in implementing the FEM, it has demonstrated that it is a technique worthy of study when one is considering atmospheric flow problems. In the future, the author hopes to use the FEM for the solution of a more complicated atmospheric model. This model is presently being used for operational weather forecasting at the Alberta Weather Centre using finite-difference techniques.

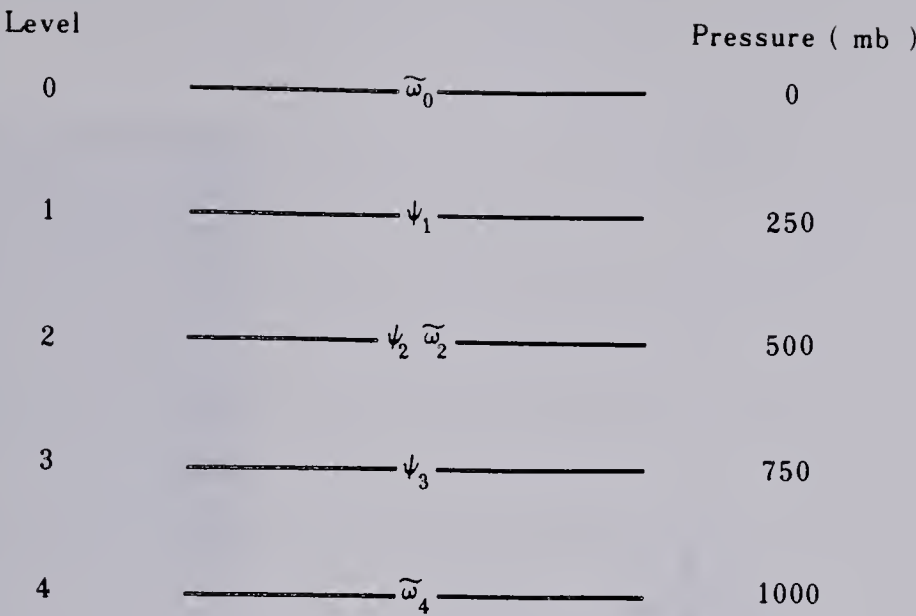


Fig. 2.1 Vertical discretization of the model showing the five levels used, with pressure as the vertical coordinate. ψ_i is the stream function at level i and $\widetilde{\omega}_i$ is the vertical velocity at level i .

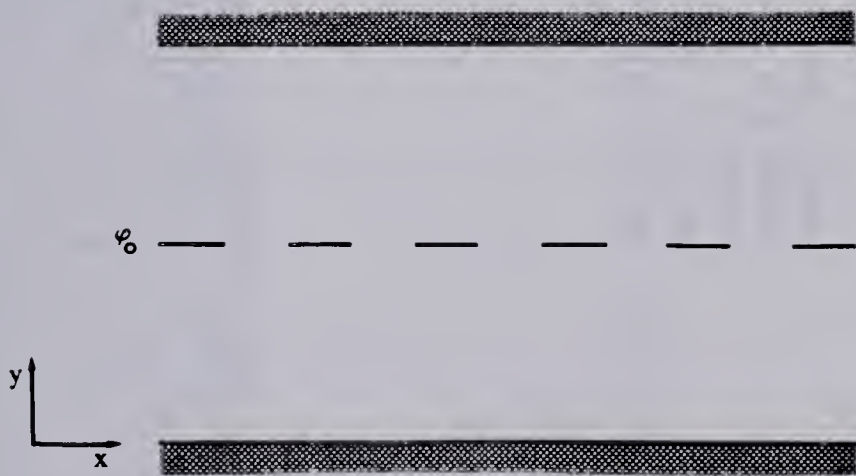


Fig. 2.2 The domain in the horizontal plane ϕ_i is the latitude at which the plane is tangent to the earth; the x-direction is east; the y-direction is north.

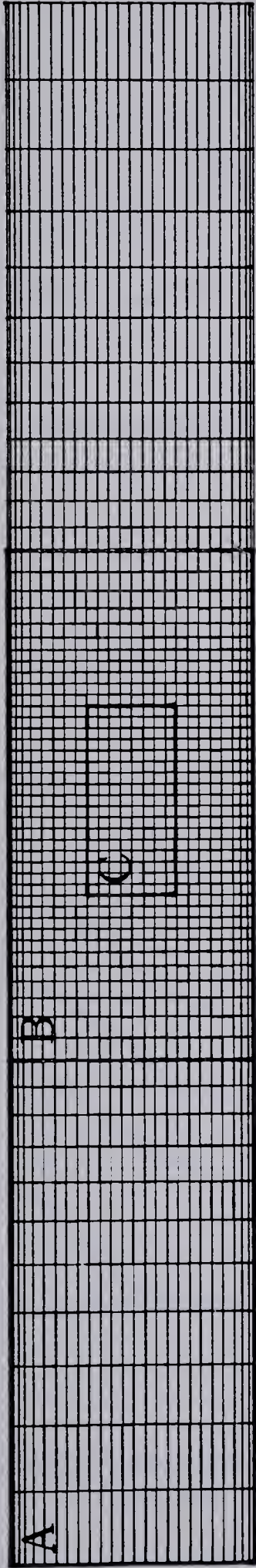
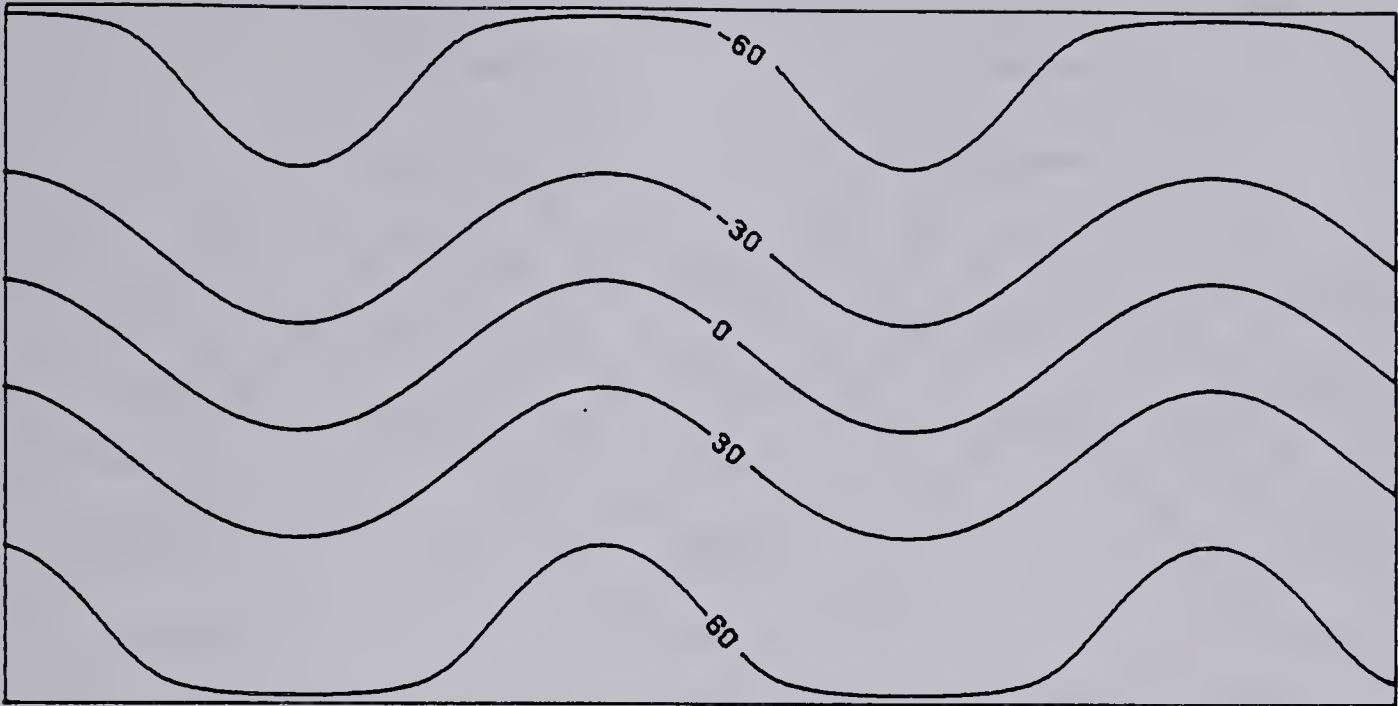
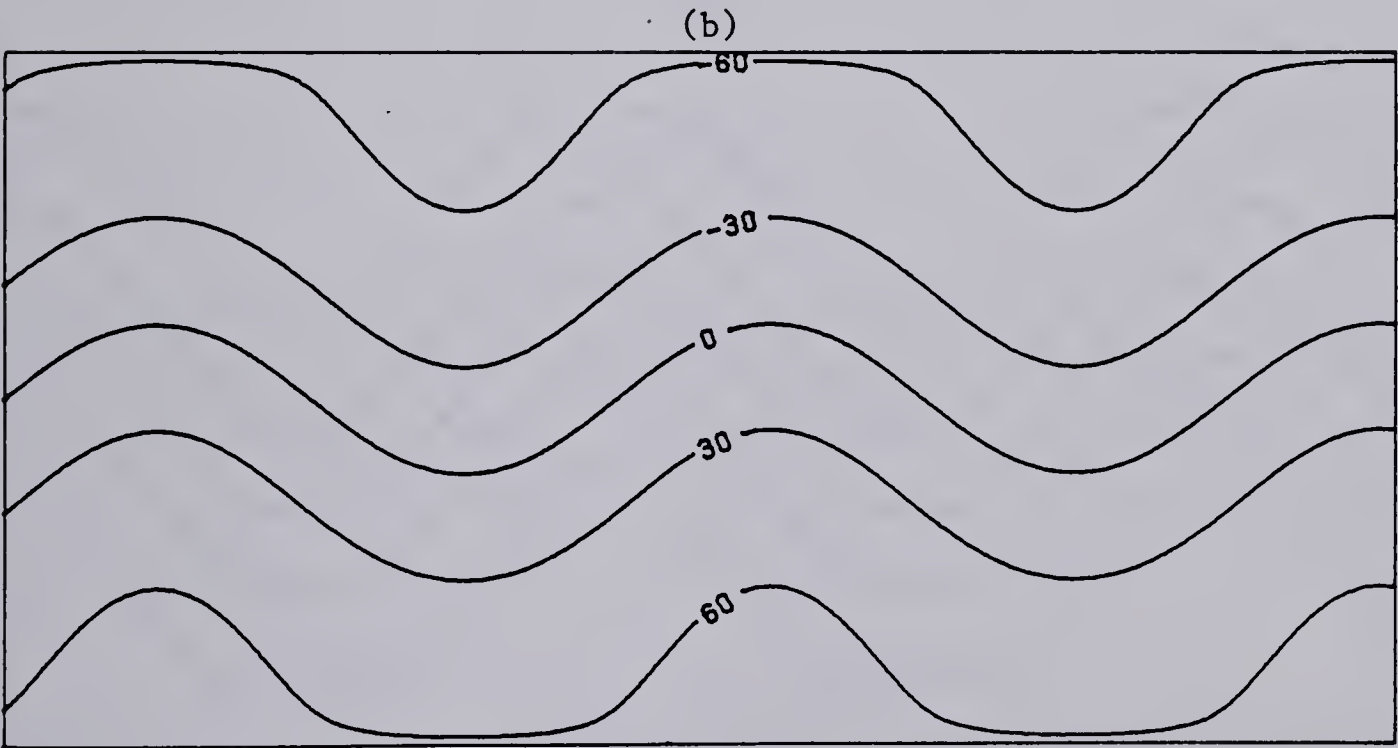


Fig. 4.1 A 62 x 27 variable grid-length mesh having a 27 x 19 sub-domain of uniform high resolution. Areas A and B are the portions of the domain used for displaying the fields in Chapter 5 while area C is the verification area.

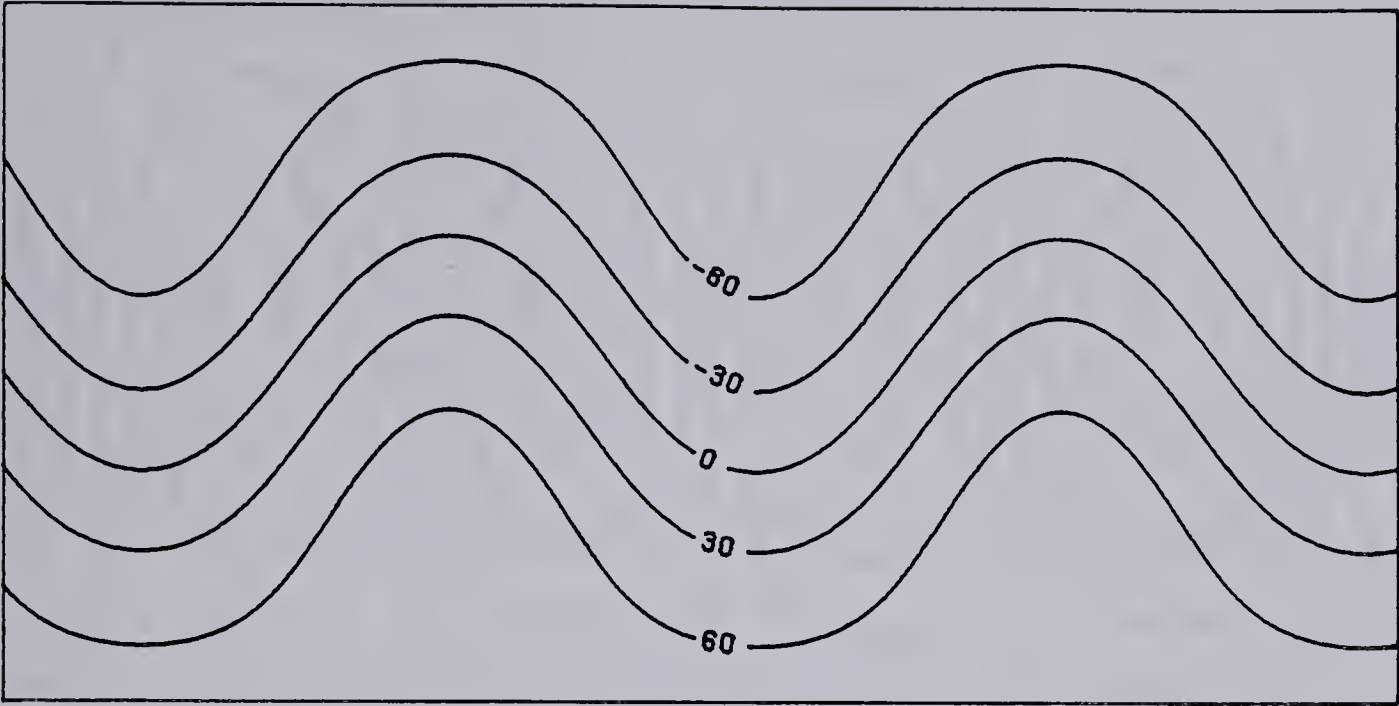


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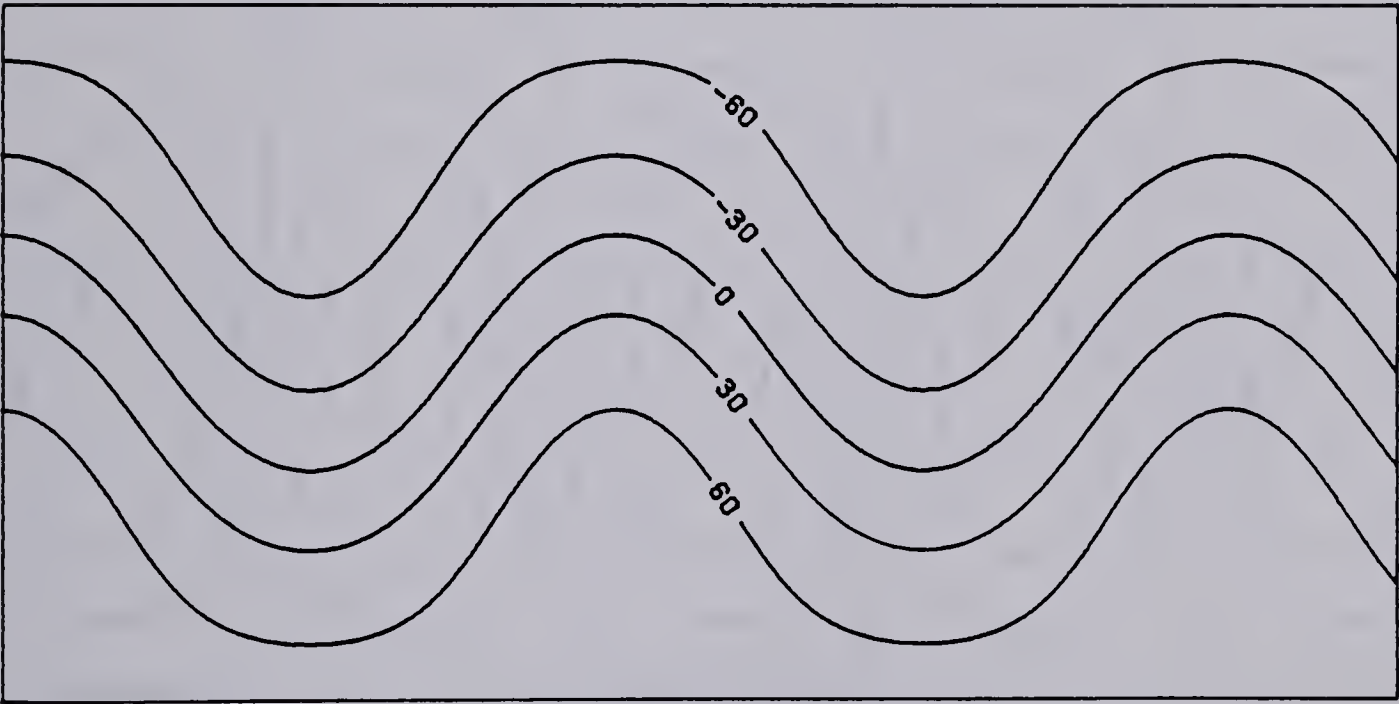


(b)

Fig. 5.1 Initial mean height field in metres for Case I in (a) area B and (b) area A.

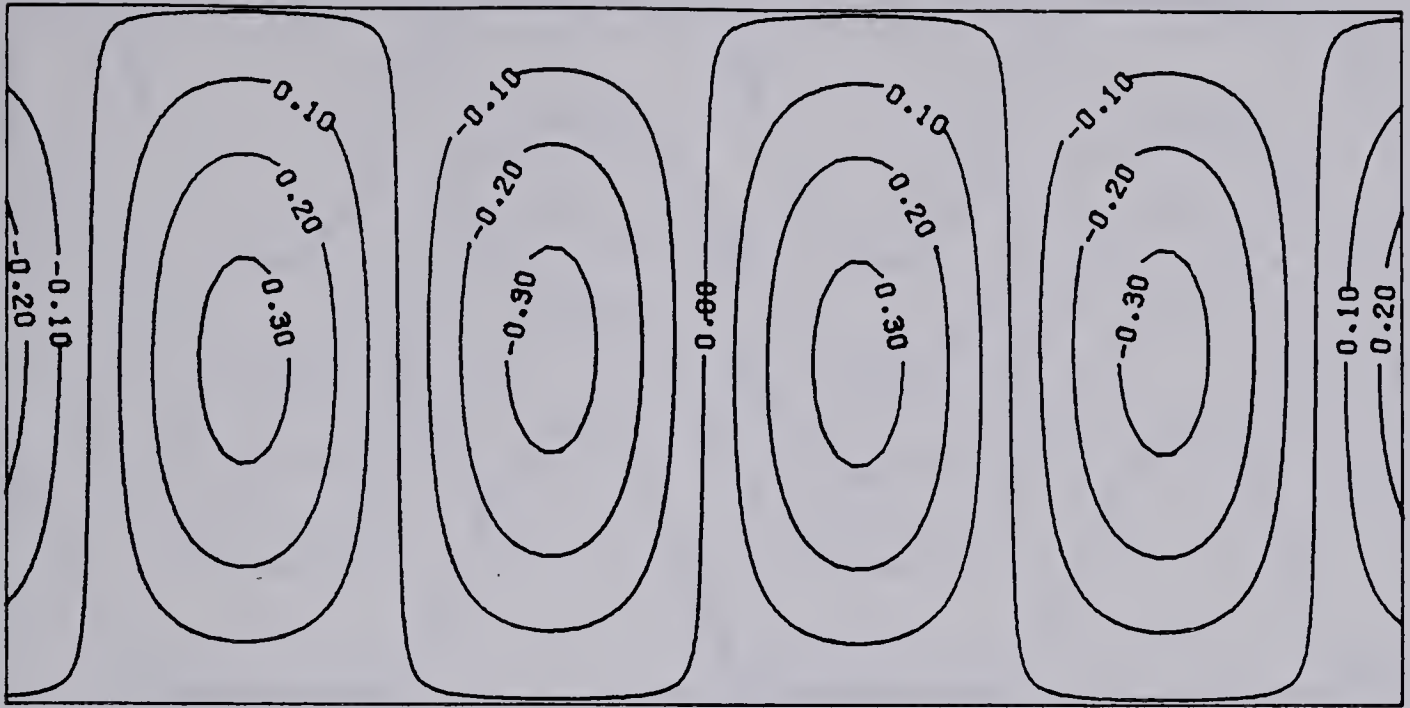


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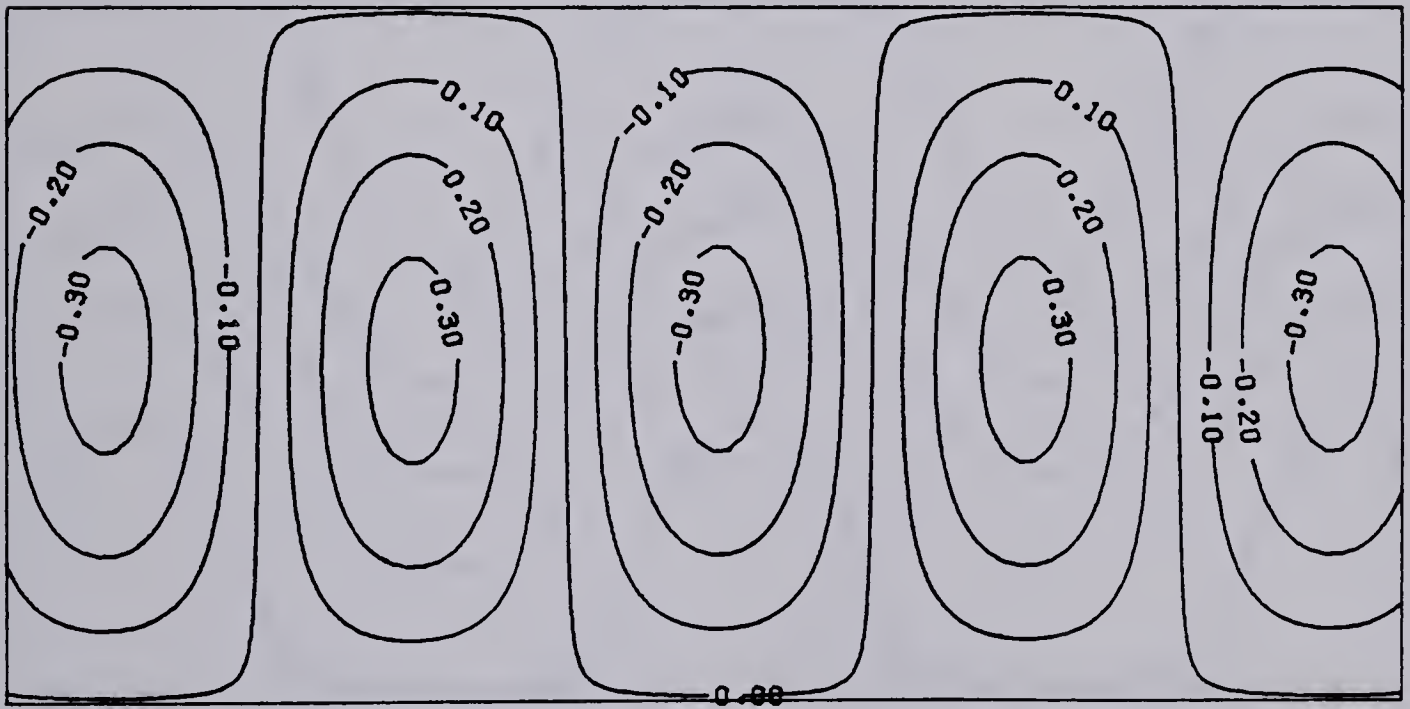


(b)

Fig. 5.2 Initial thickness field in metres for Case I in (a) area B and (b) area A.

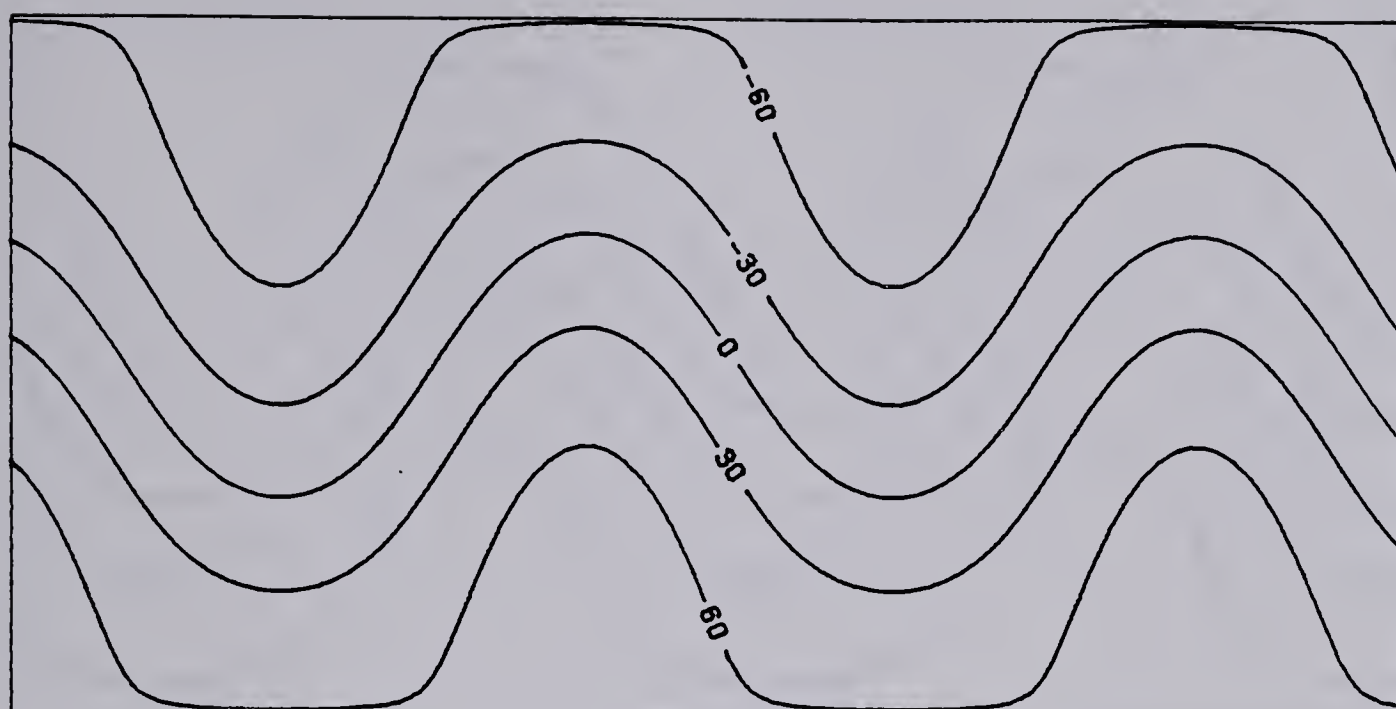


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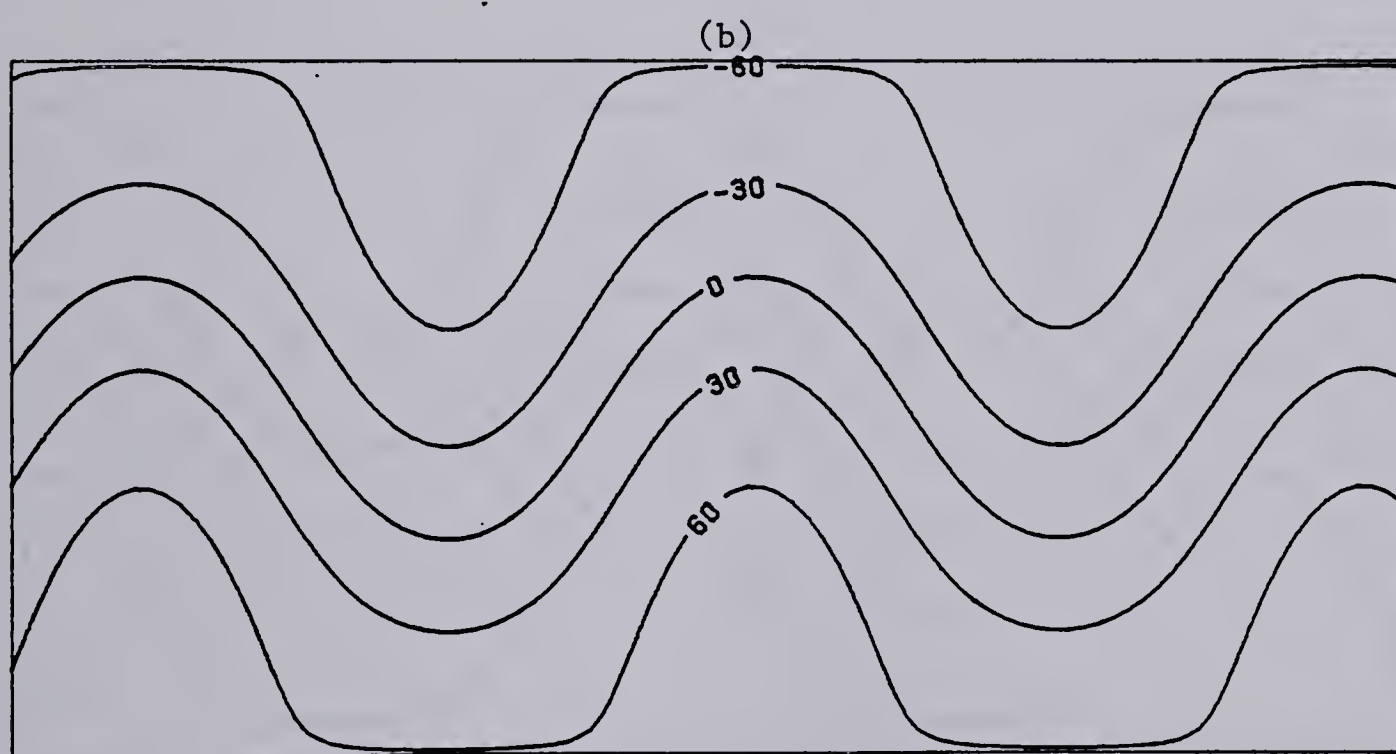


(b)

Fig. 5.3 Initial omega field in $\mu\text{bar}/\text{sec}$ for Case I in (a) area B and (b) area A.

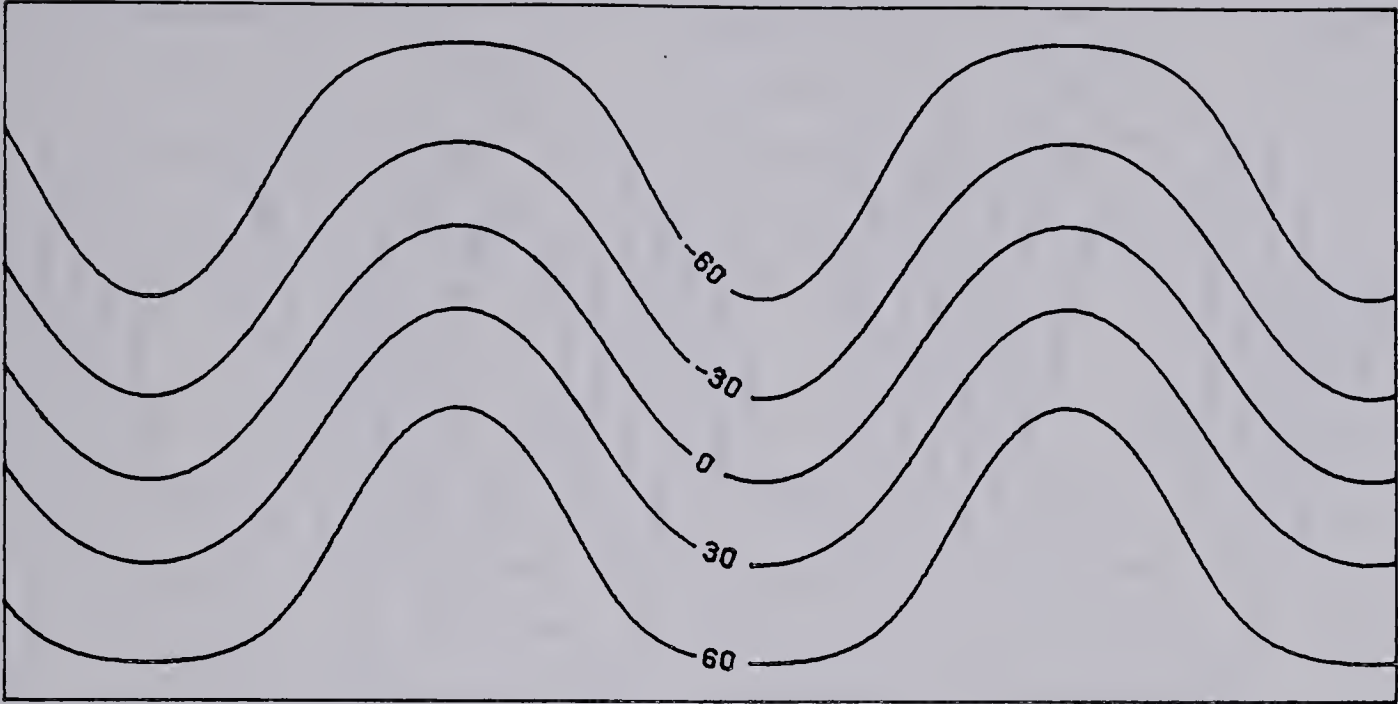


(a)



(b)

Fig. 5.4 The 24-hour spectral solution for the mean height field in metres for Case I in (a) area B and (b) area A.



(a)

(b)

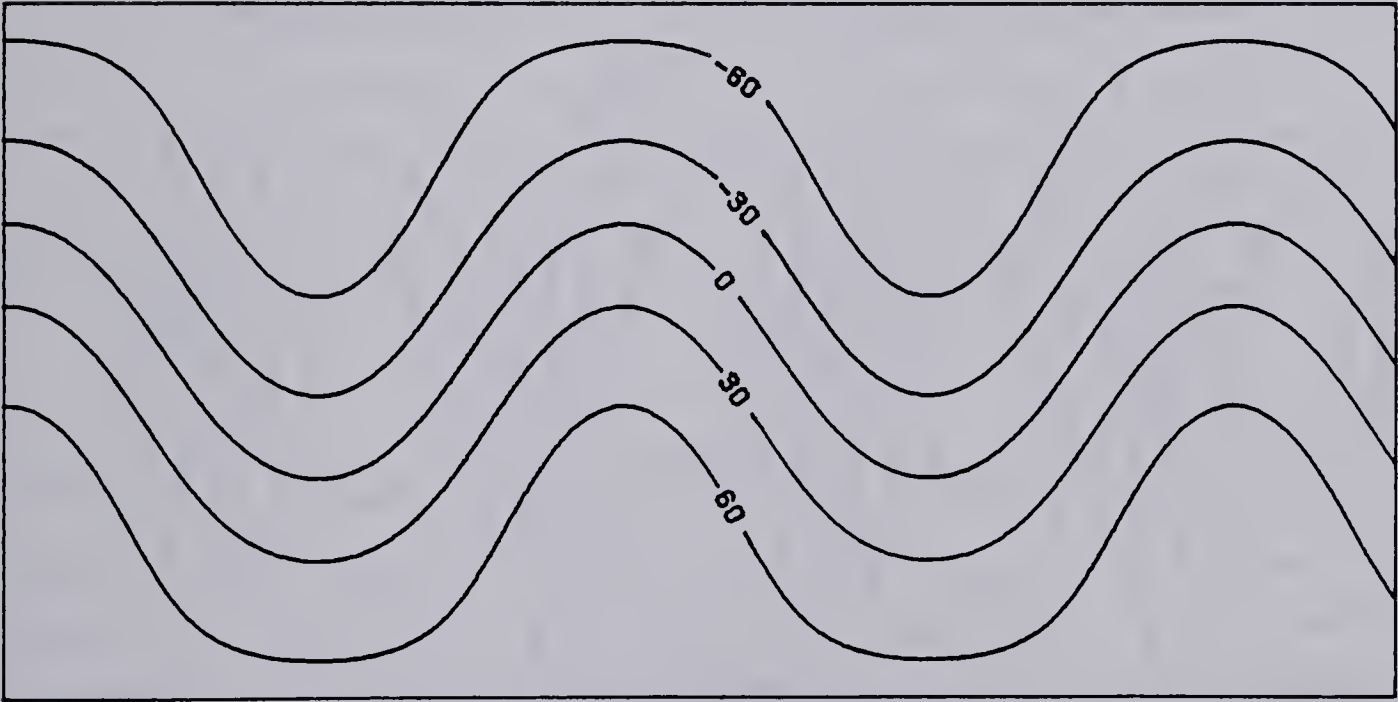
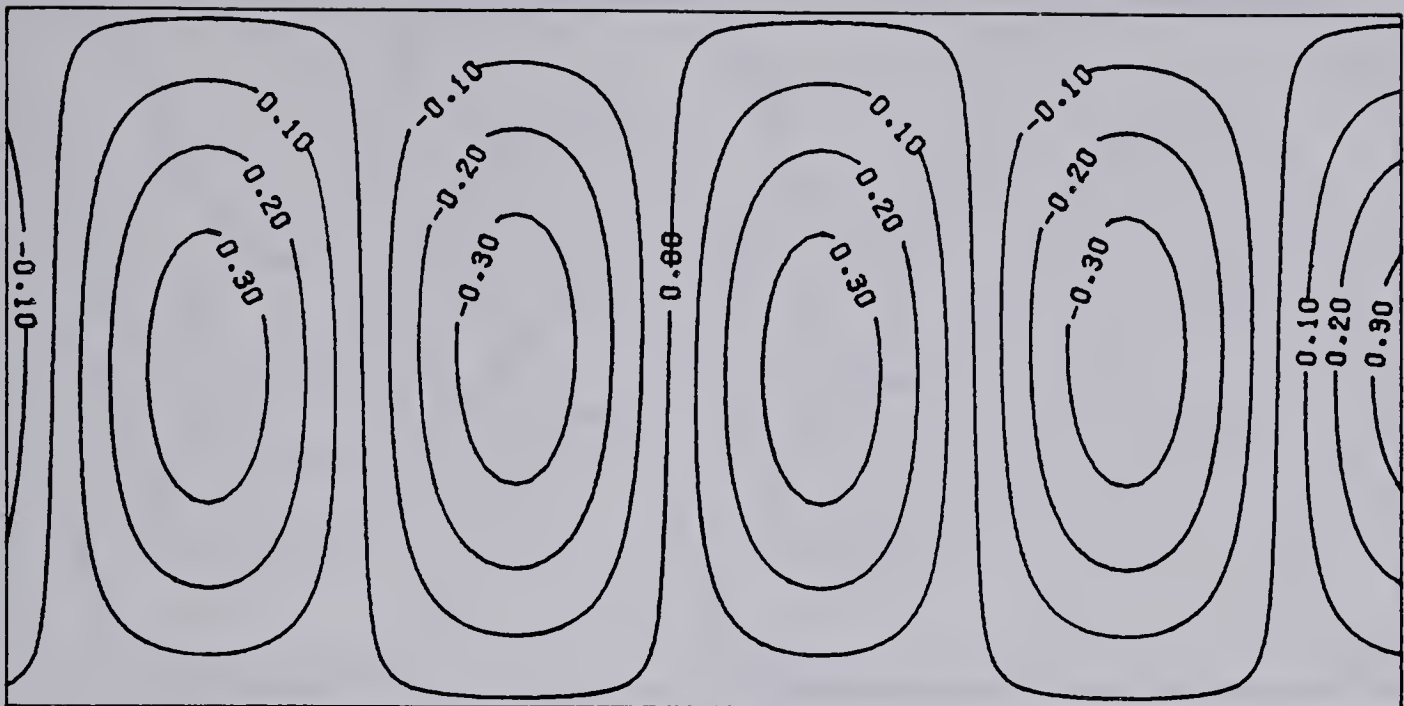
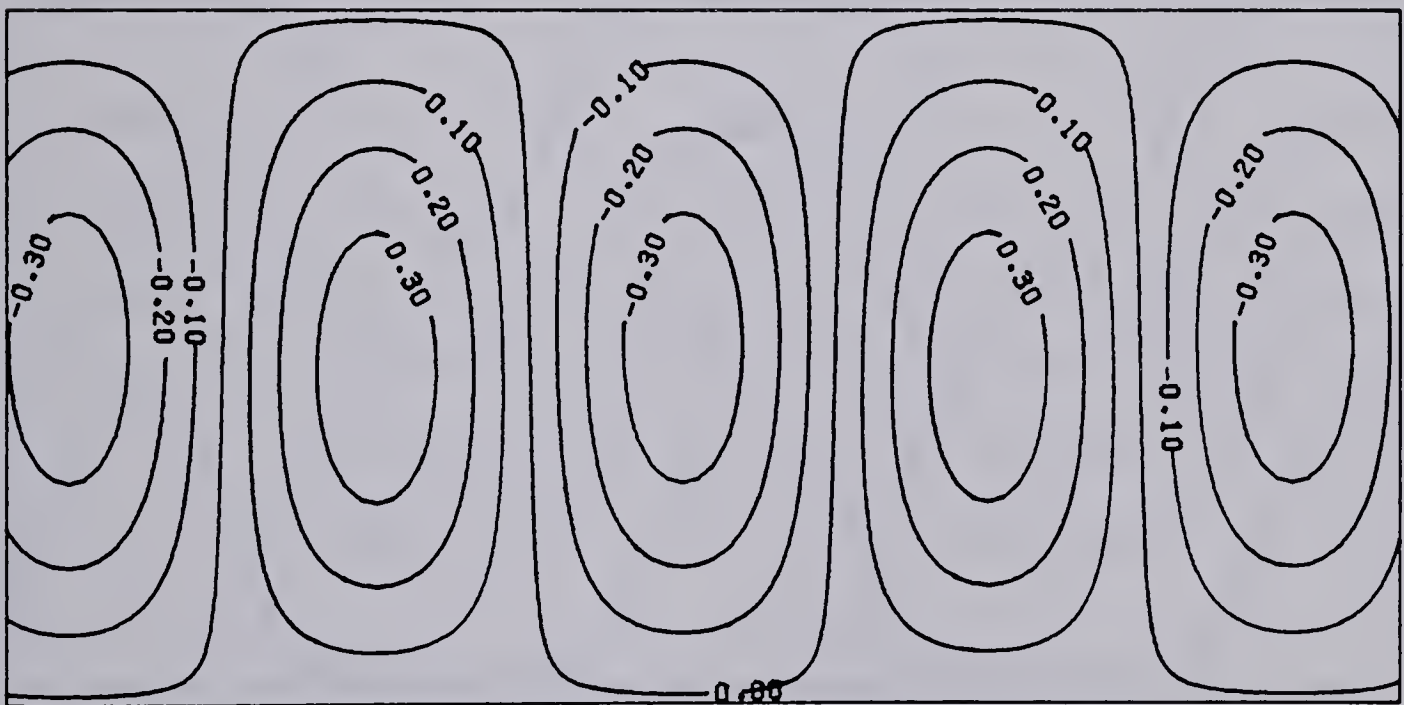


Fig. 5.5 The 24-hour spectral solution for the thickness field in metres for Case I in (a) area B and (b) area A.

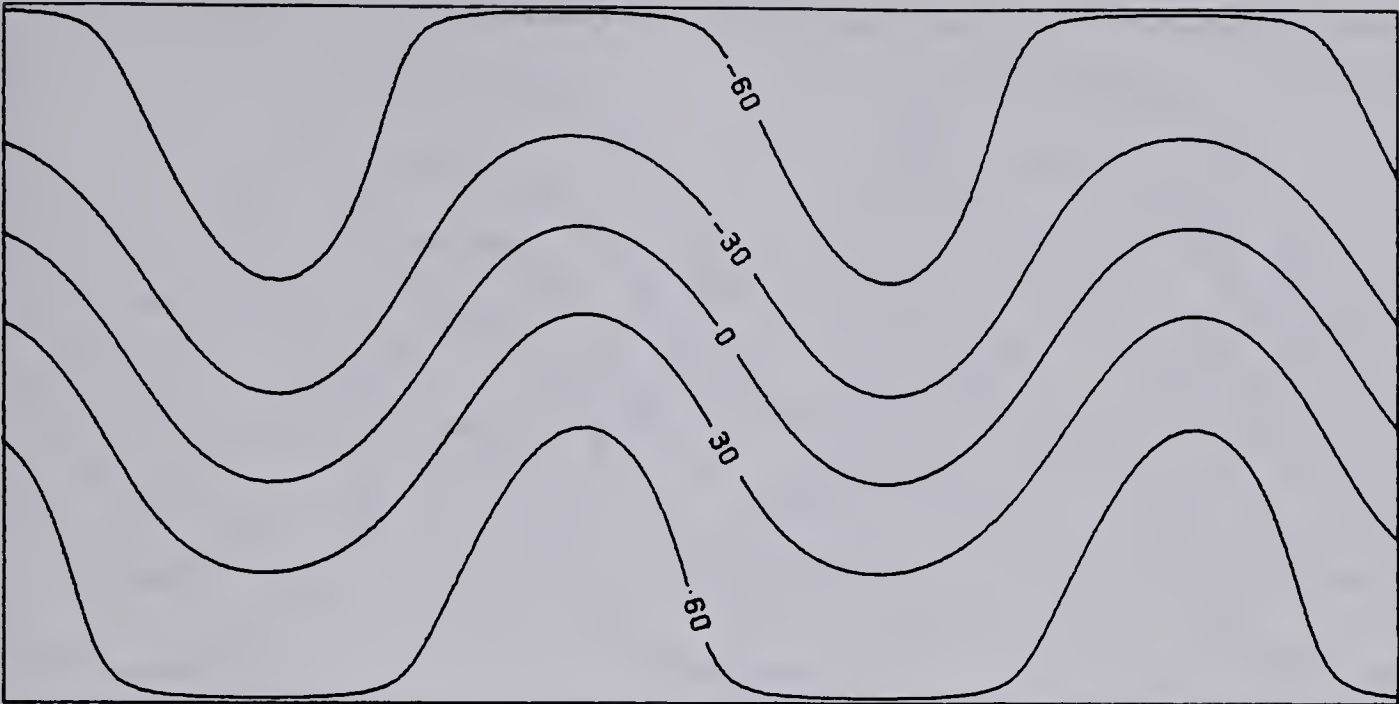


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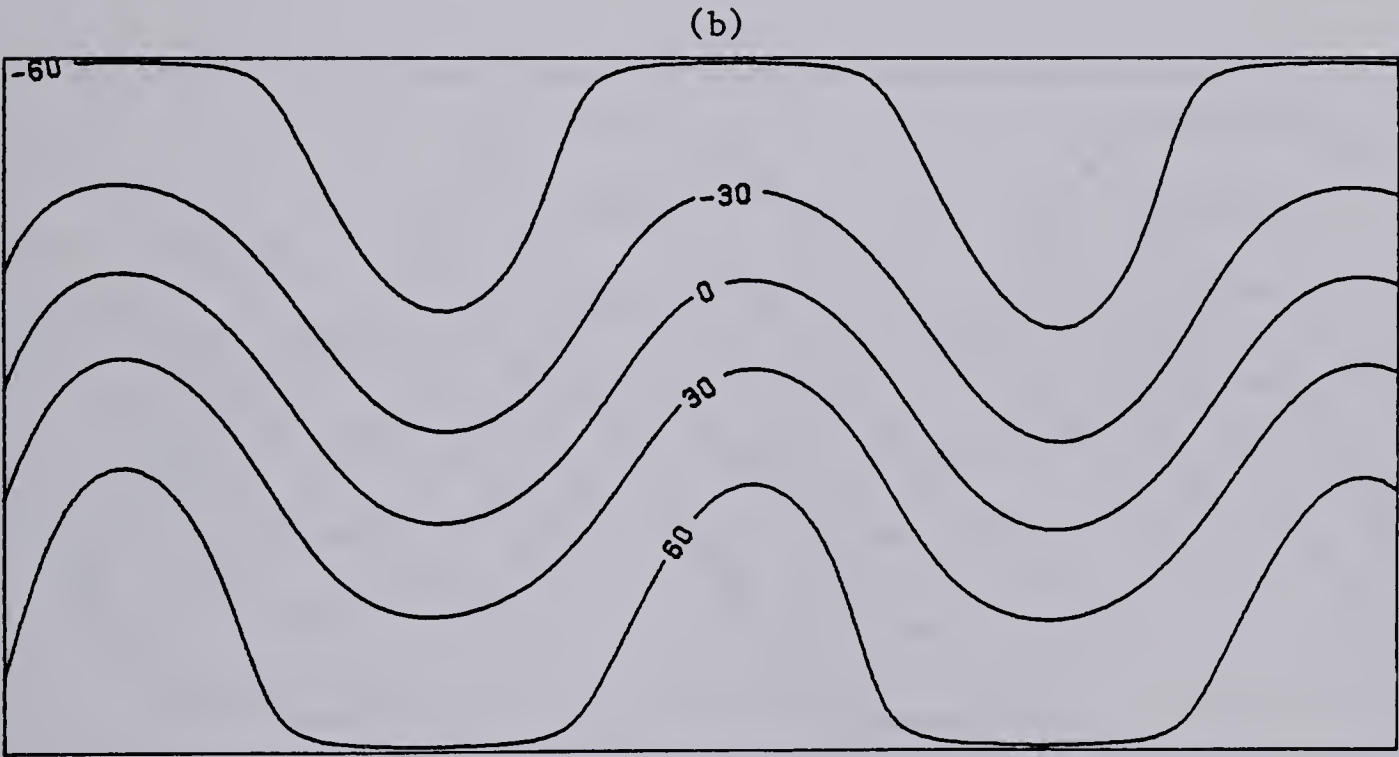


(b)

Fig. 5.6 The 24-hour spectral solution for the omega field in $\mu\text{bar/sec}$ for Case I in (a) area B and (b) area A.

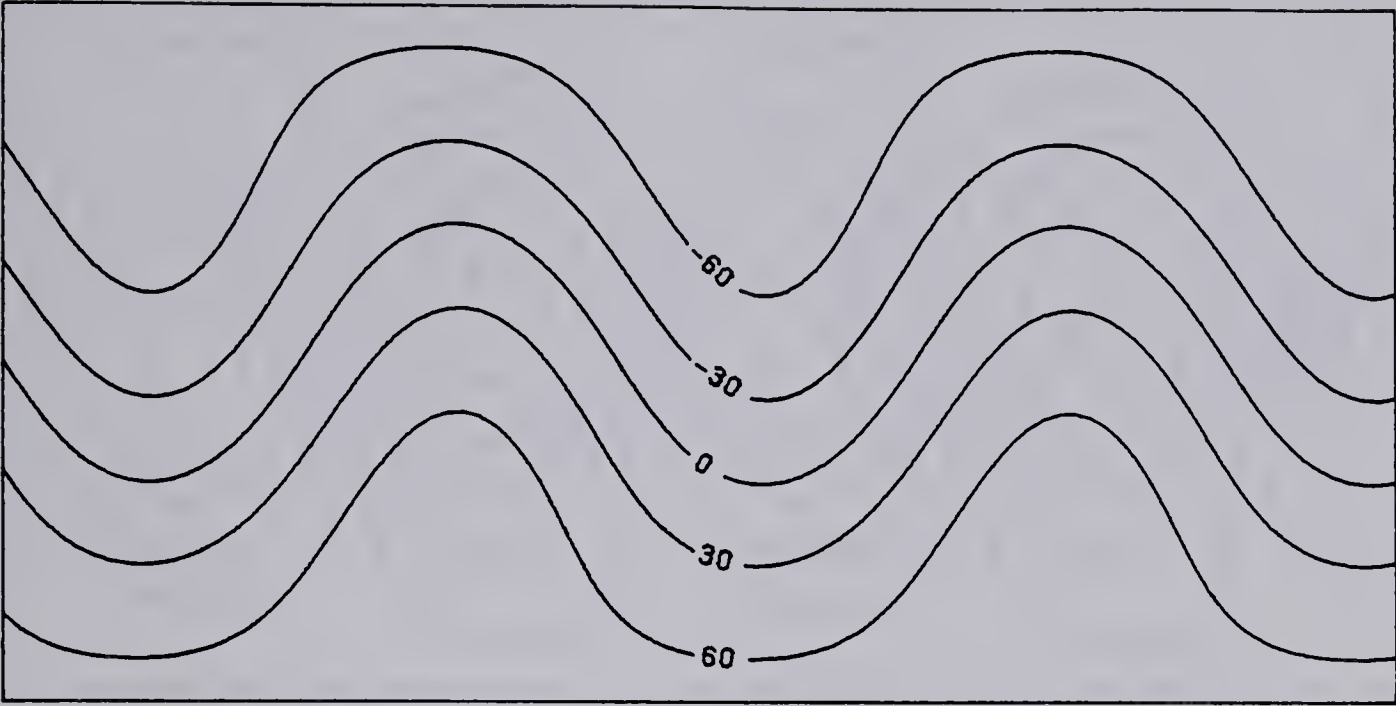


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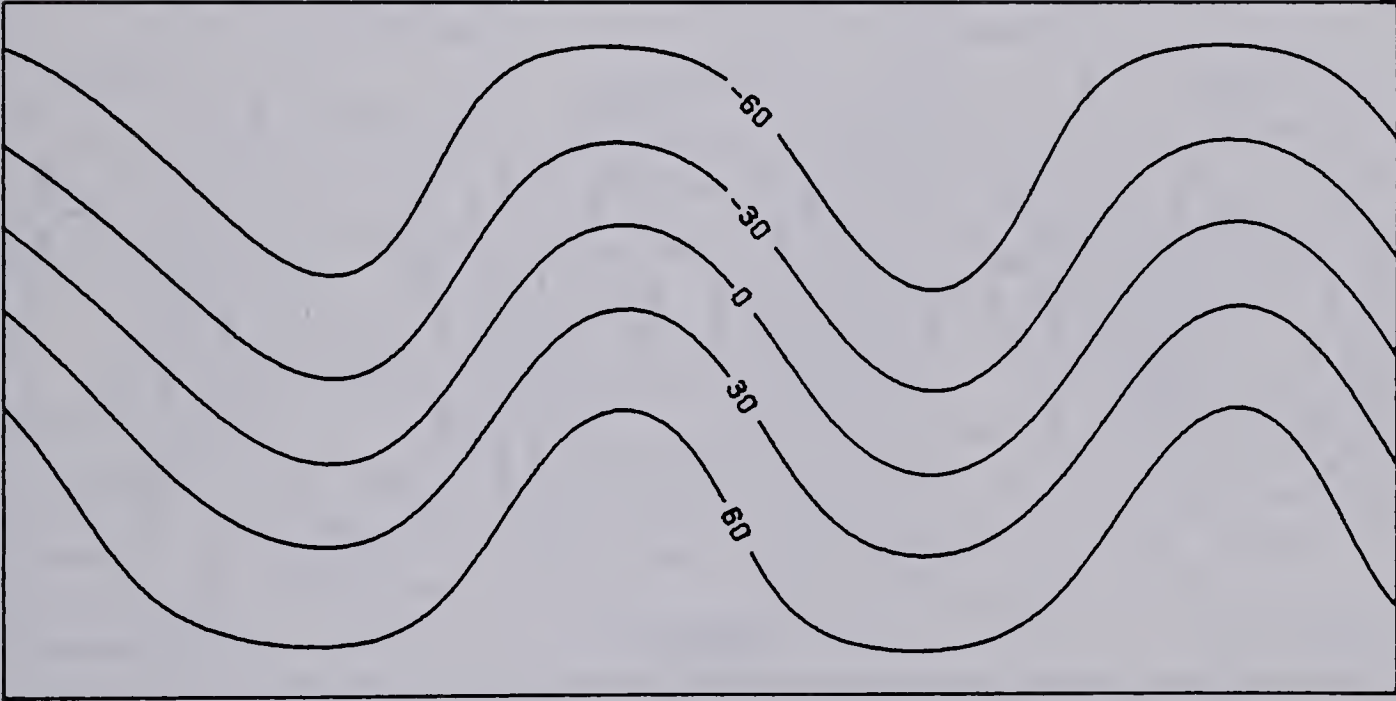


(b)

Fig. 5.7 The 24-hour finite-element solution for the mean height field in metres for Case I in (a) area B and (b) area A.

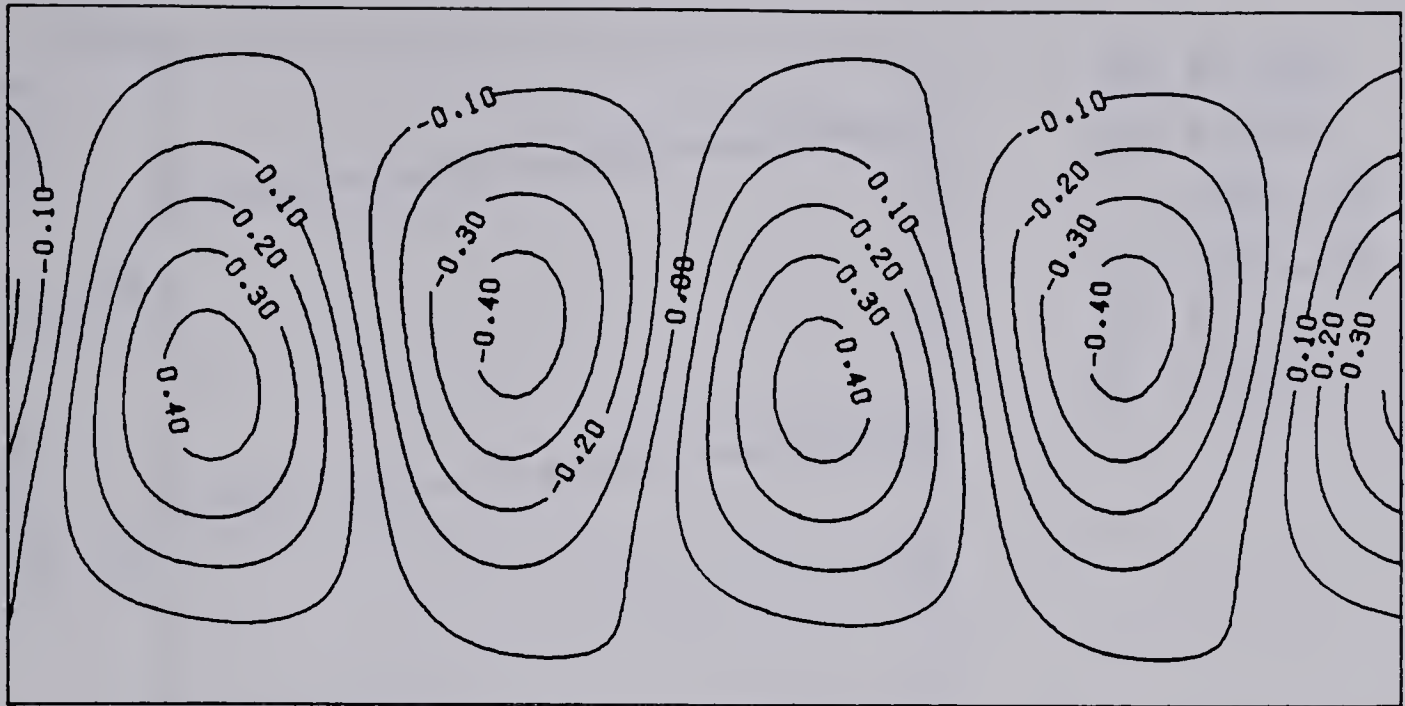


(a)

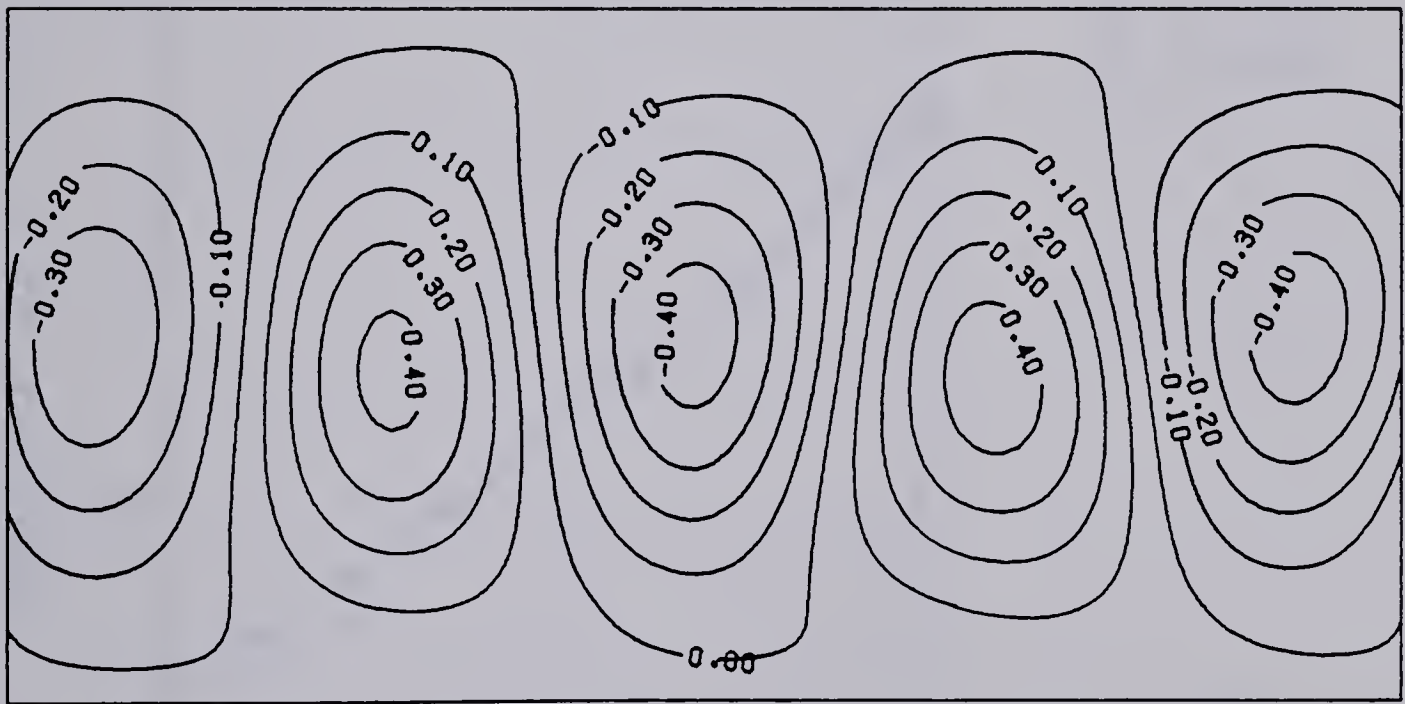


(b)

Fig. 5.8 The 24-hour finite-element solution for the thickness field in metres for Case I in (a) area B and (b) area A.

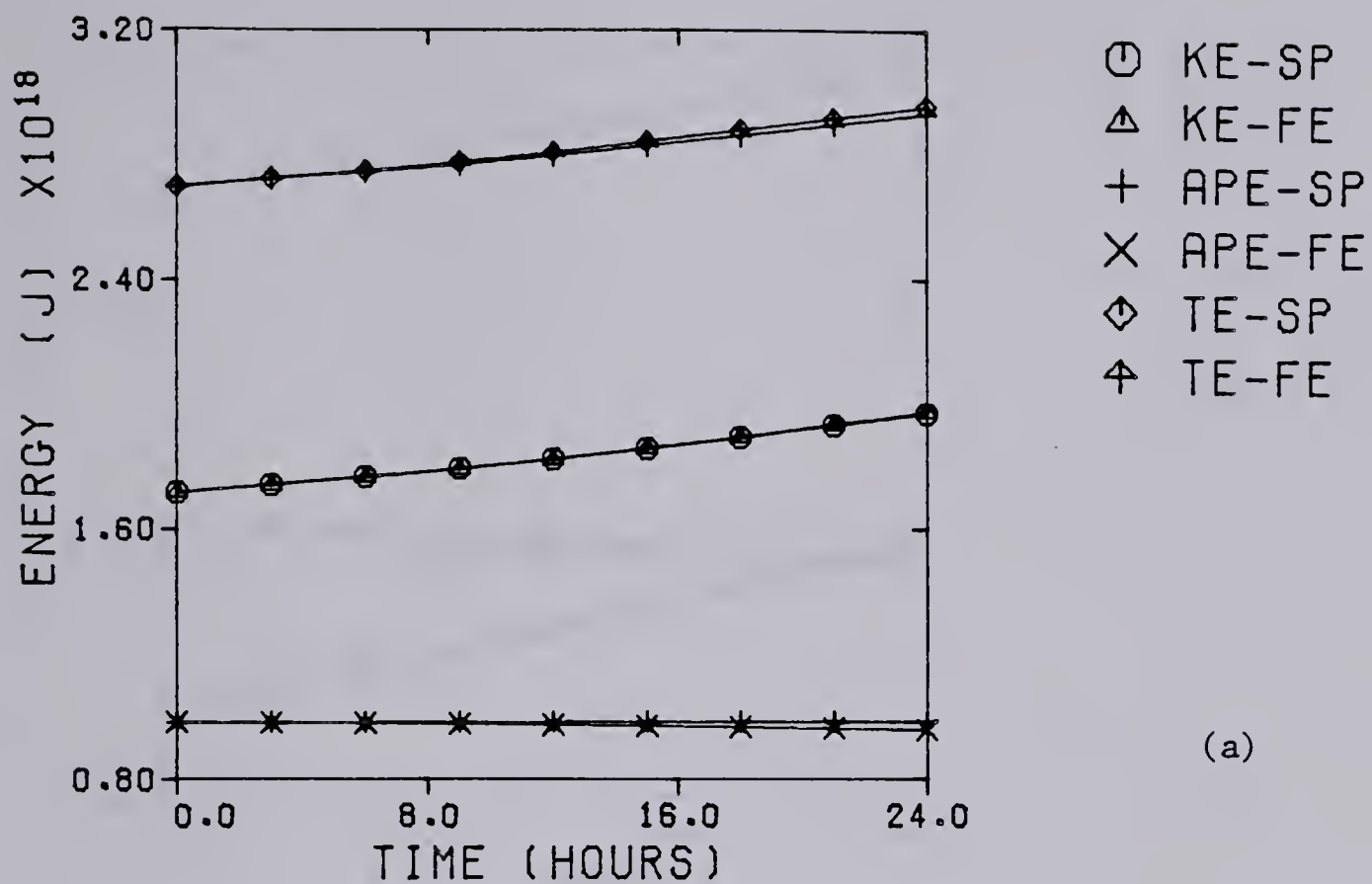


(a)

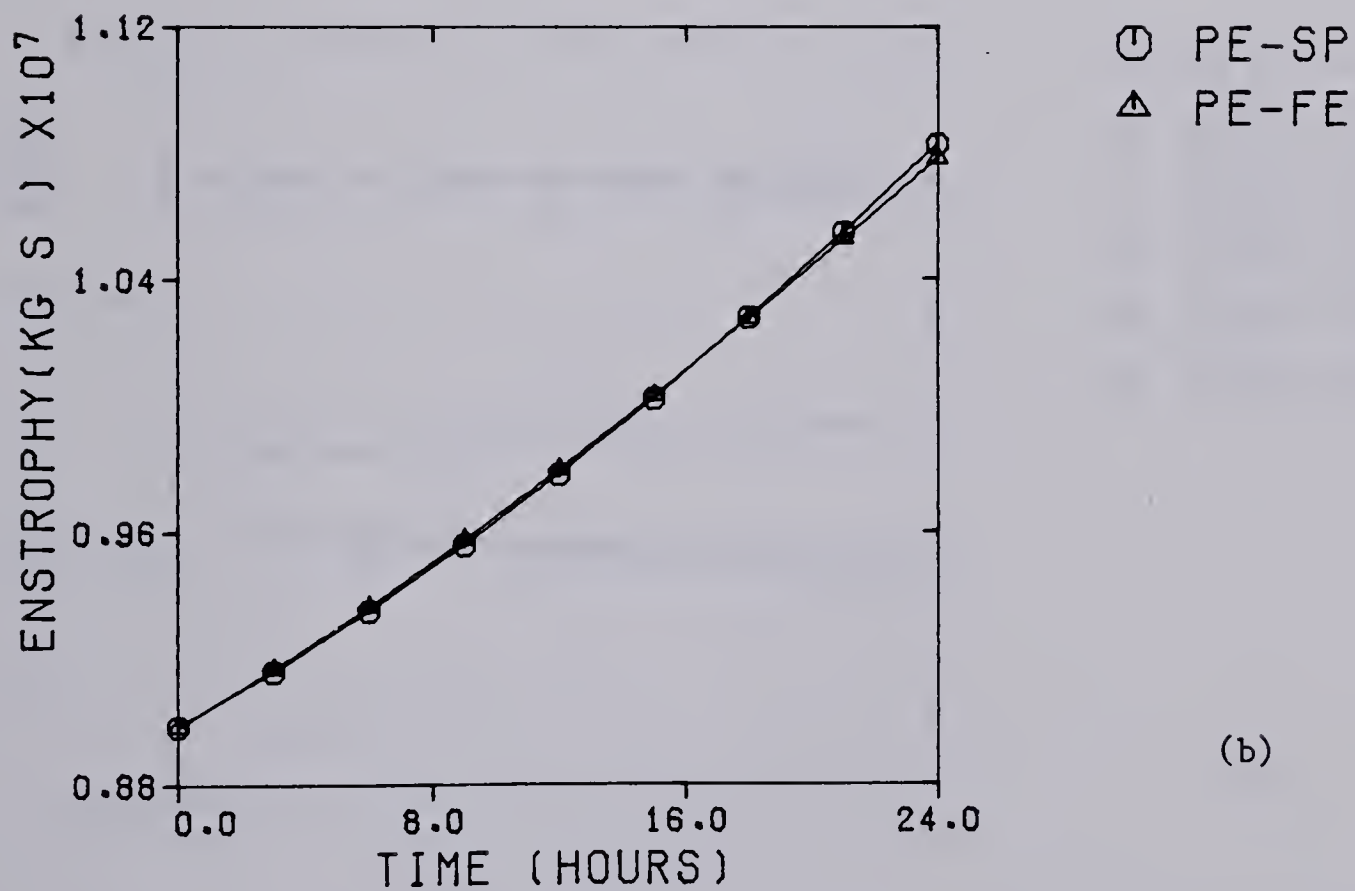


(b)

Fig. 5.9 The 24-hour finite-element solution for the omega field in $\mu\text{bar/sec}$ for Case I in (a) area B and (b) area A.

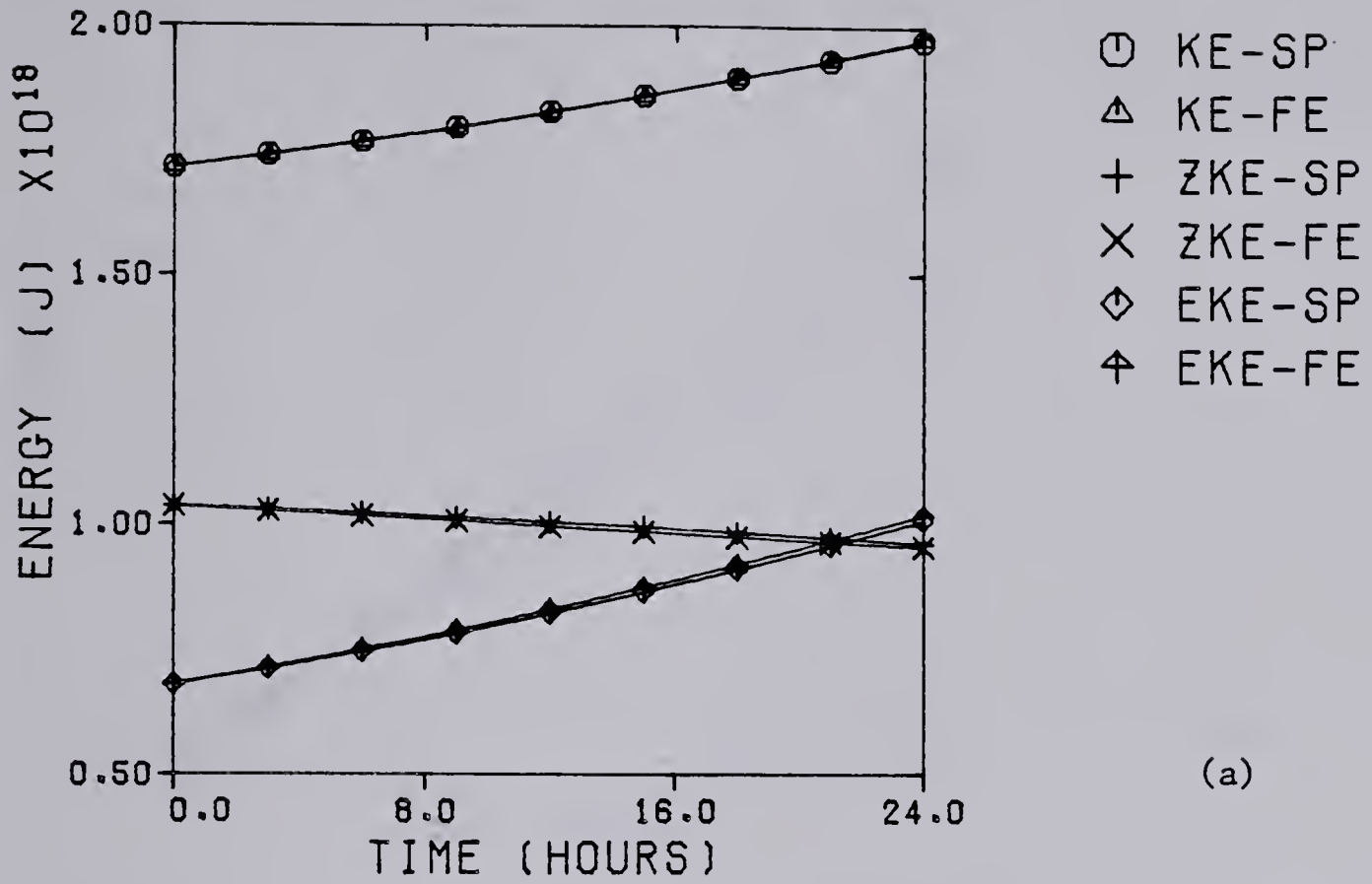


(a)

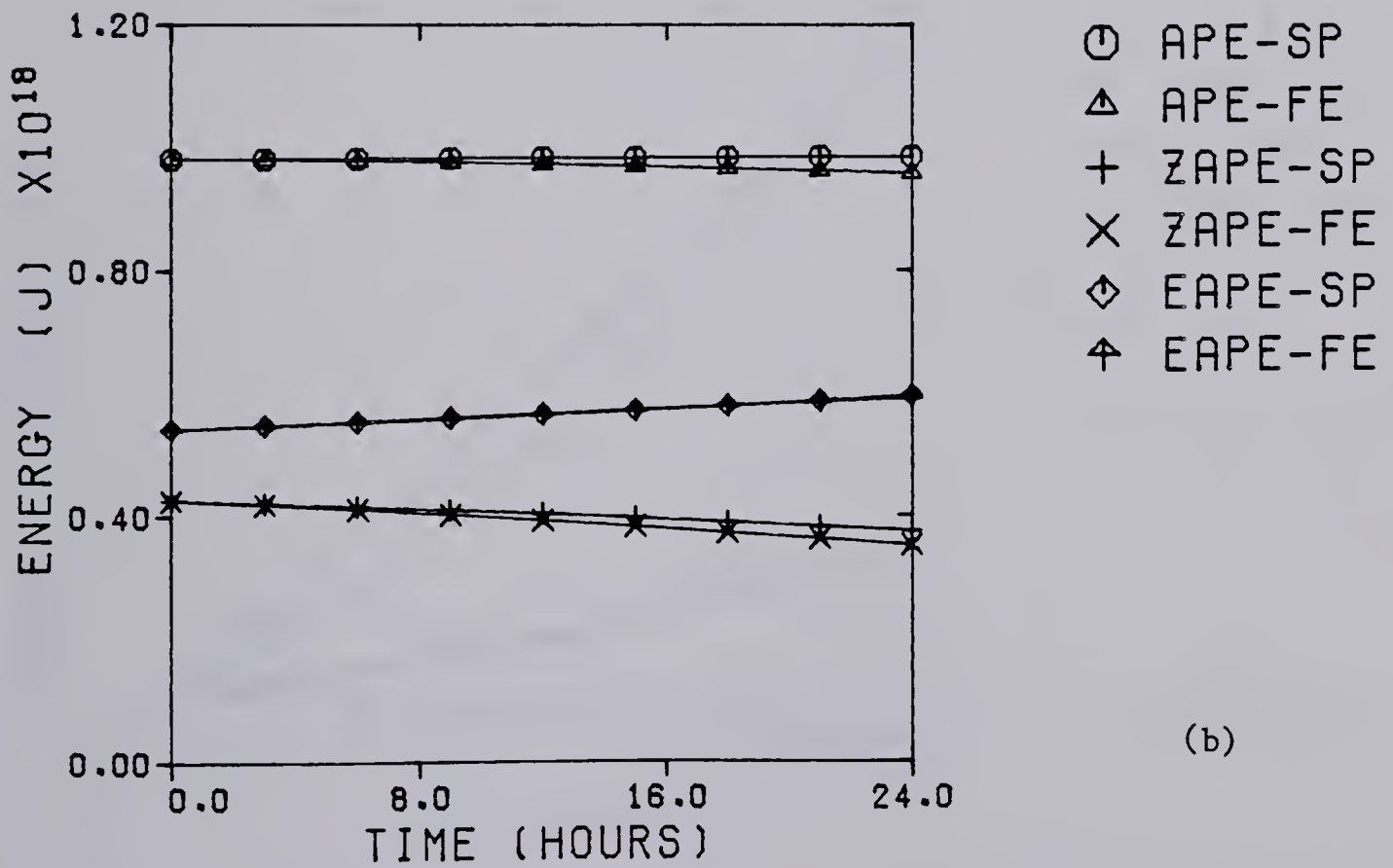


(b)

Fig. 5.10 The (a) energies and (b) potential enstrophy of Case I for the two solutions in the verification area.

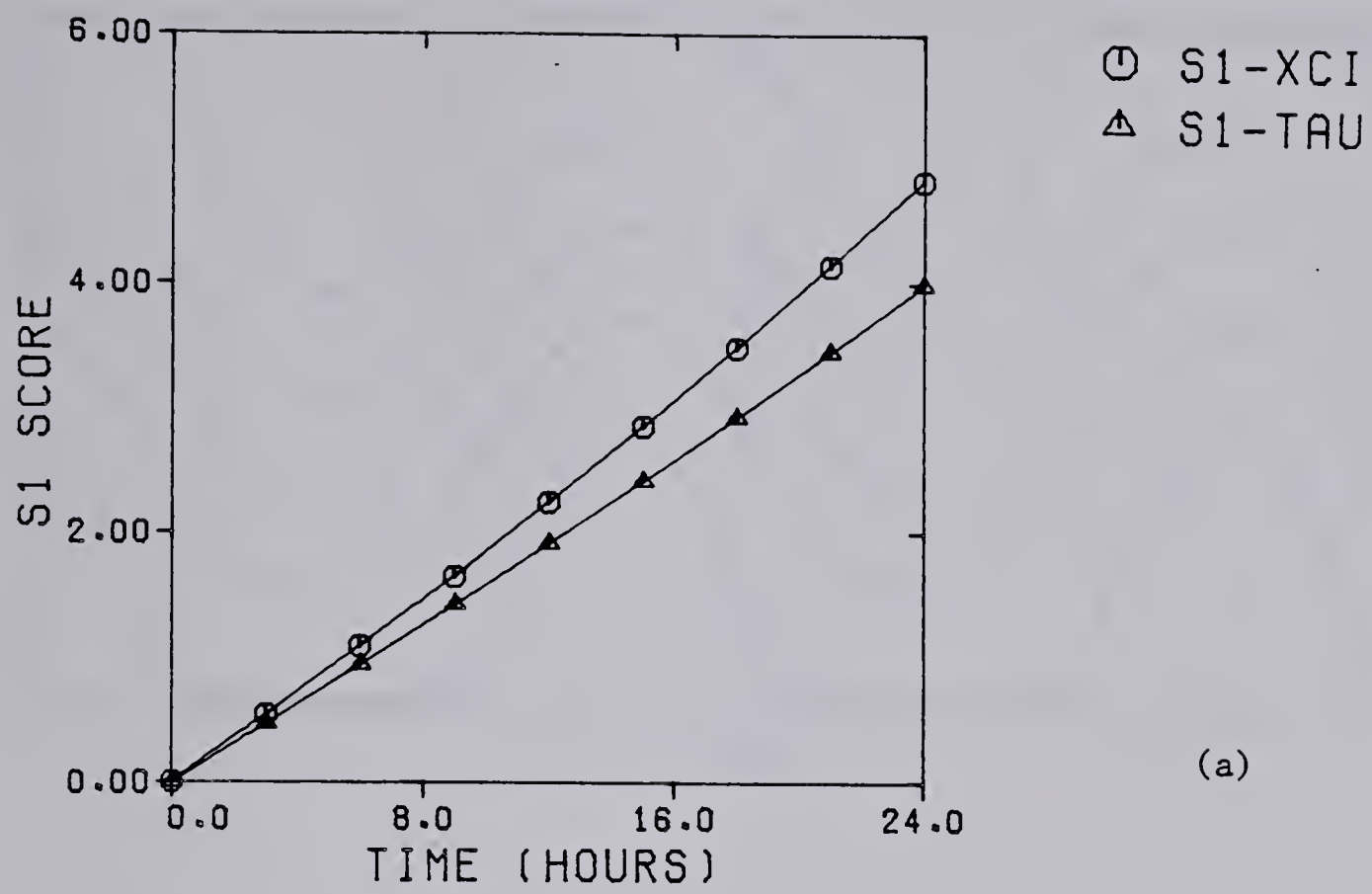


(a)

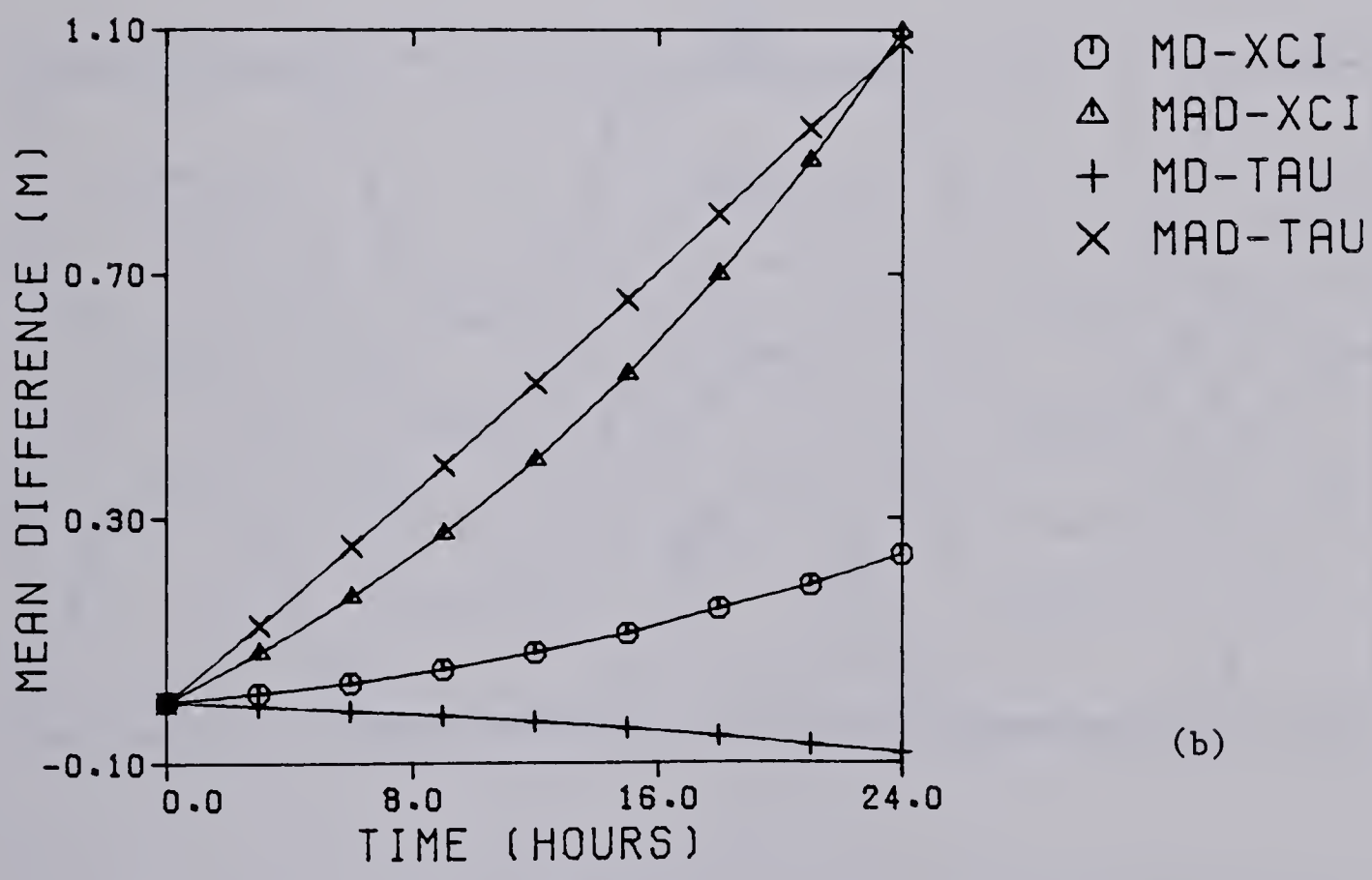


(b)

Fig. 5.11 The (a) kinetic energies and (b) potential energies of Case I for the two solutions in the verification area.

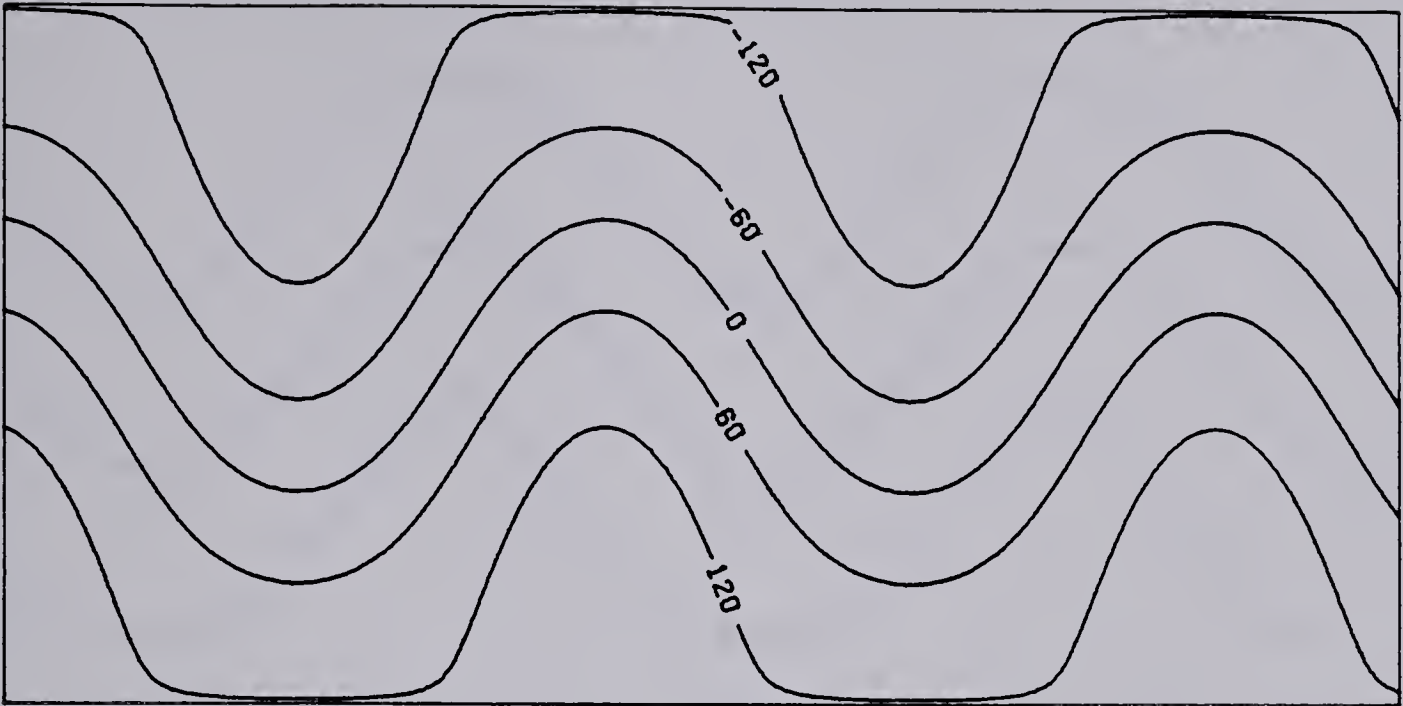


(a)

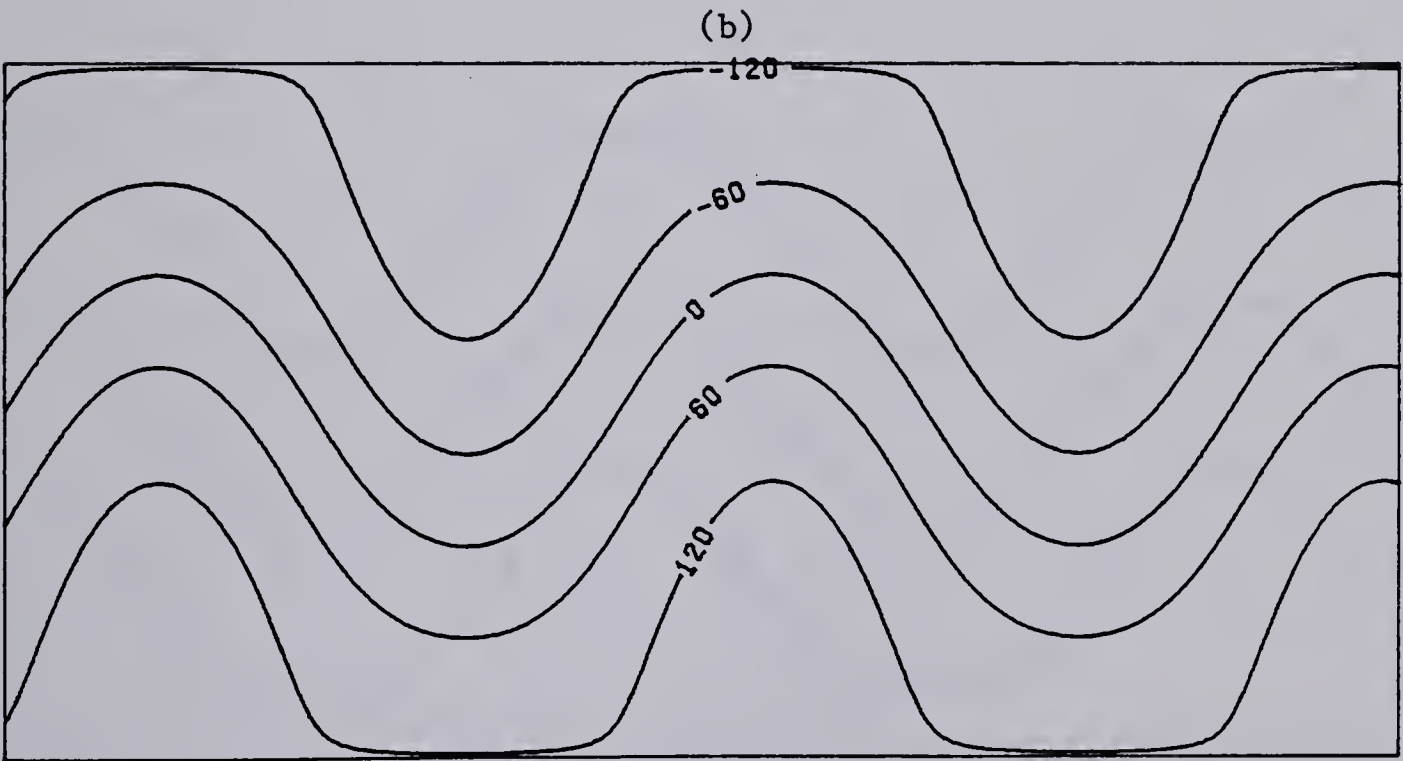


(b)

Fig. 5.12 The (a) S1 scores and (b) the MD and MAD curves for Case I for the two solutions in the verification area.

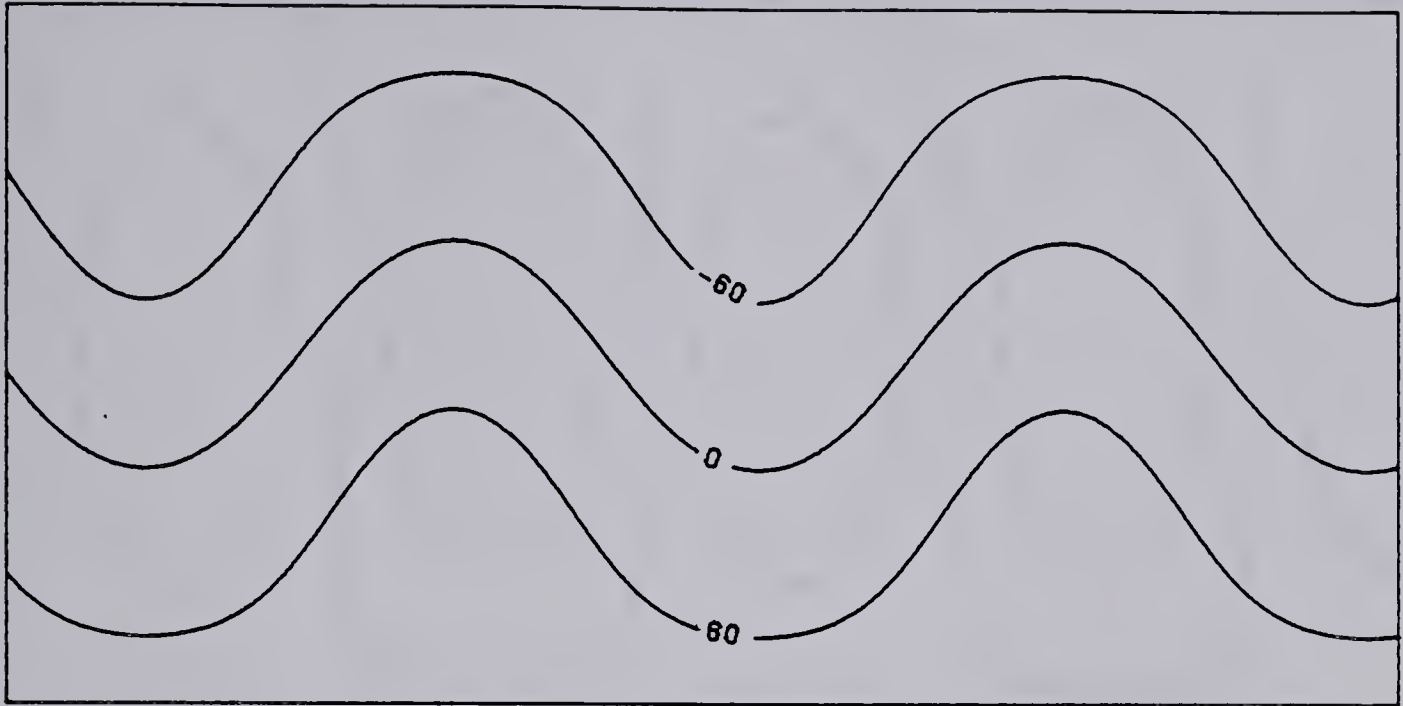


(a)

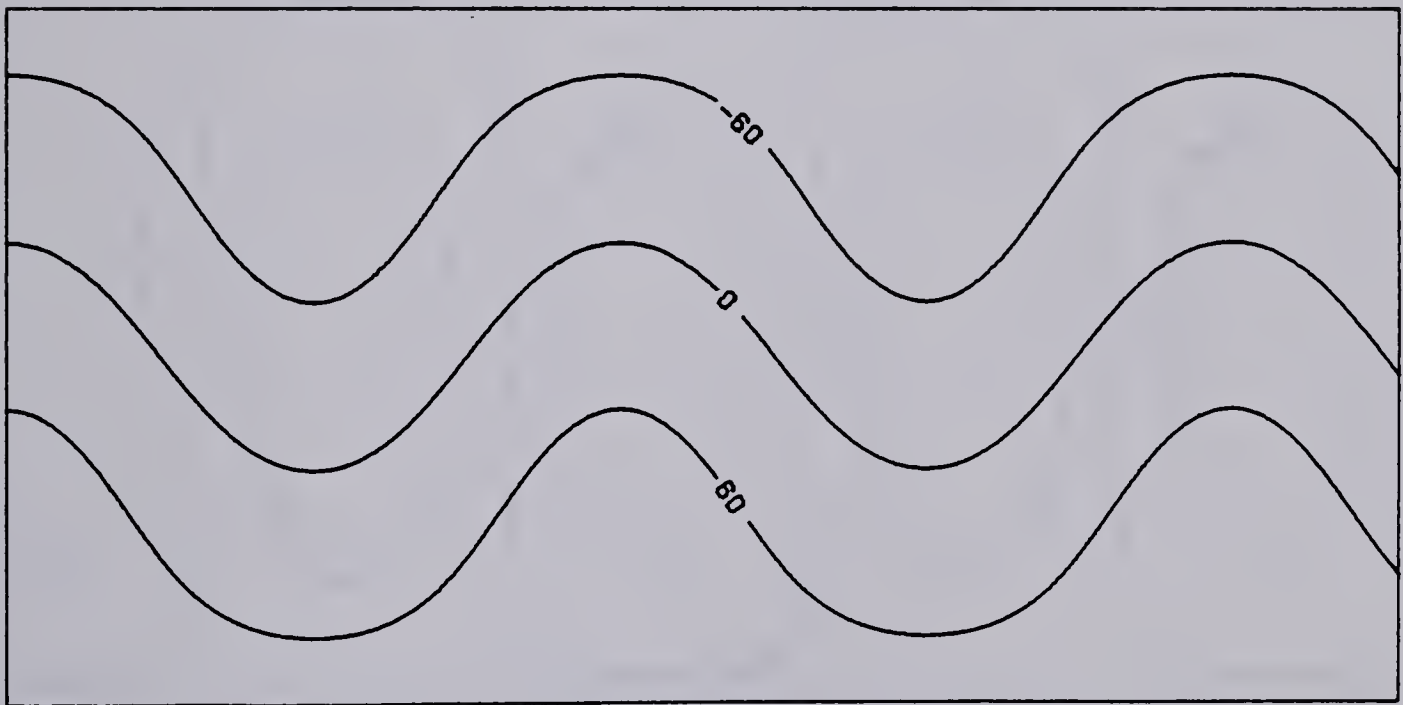


(b)

Fig. 5.13 Initial mean height field in metres for Case II in (a) area B and (b) area A.

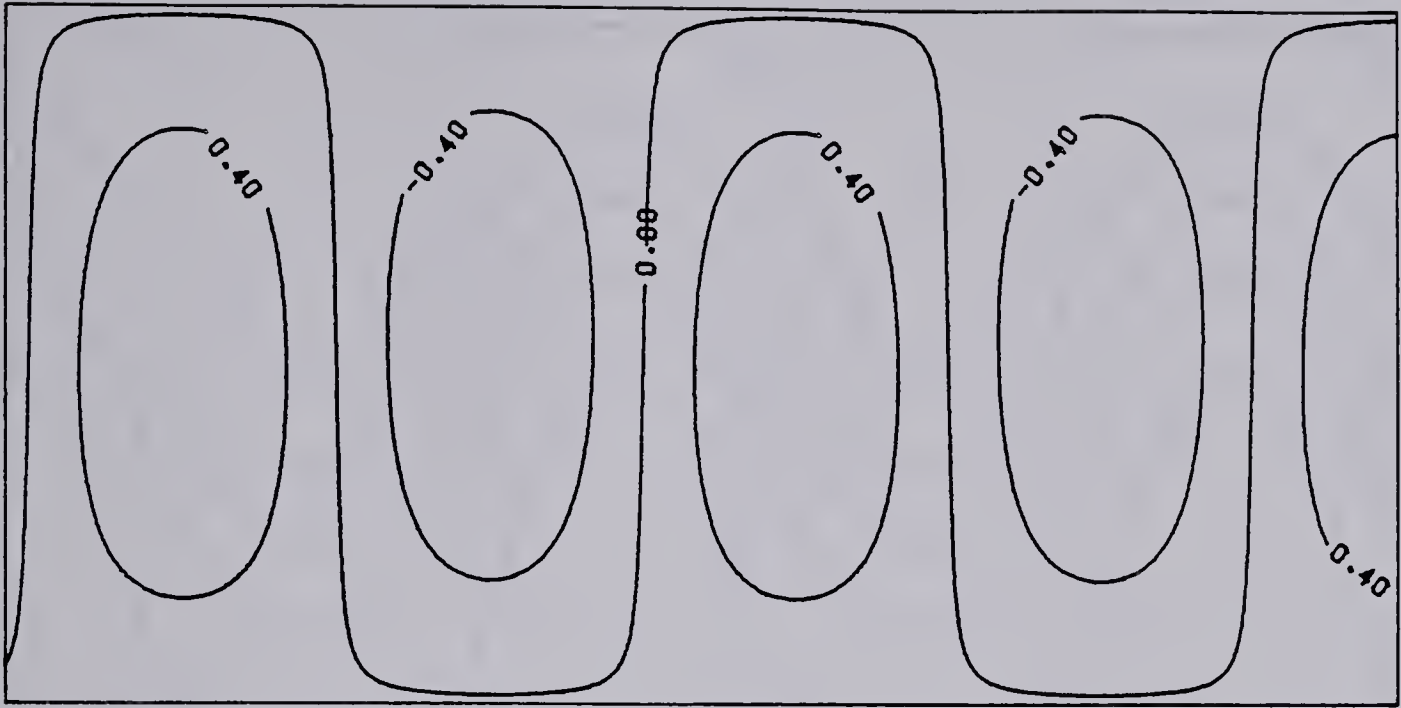


(a)

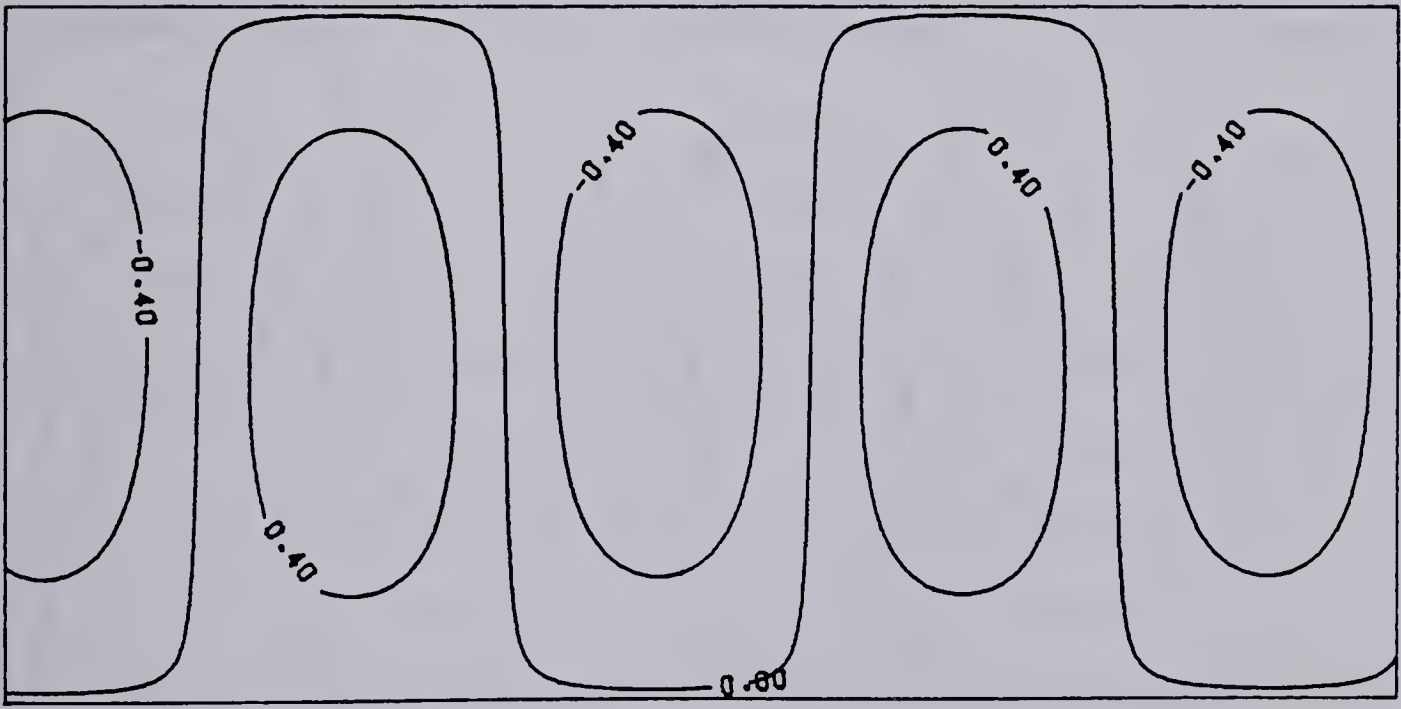


(b)

Fig. 5.14 Initial thickness field in metres for Case II in (a) area B and (b) area A.

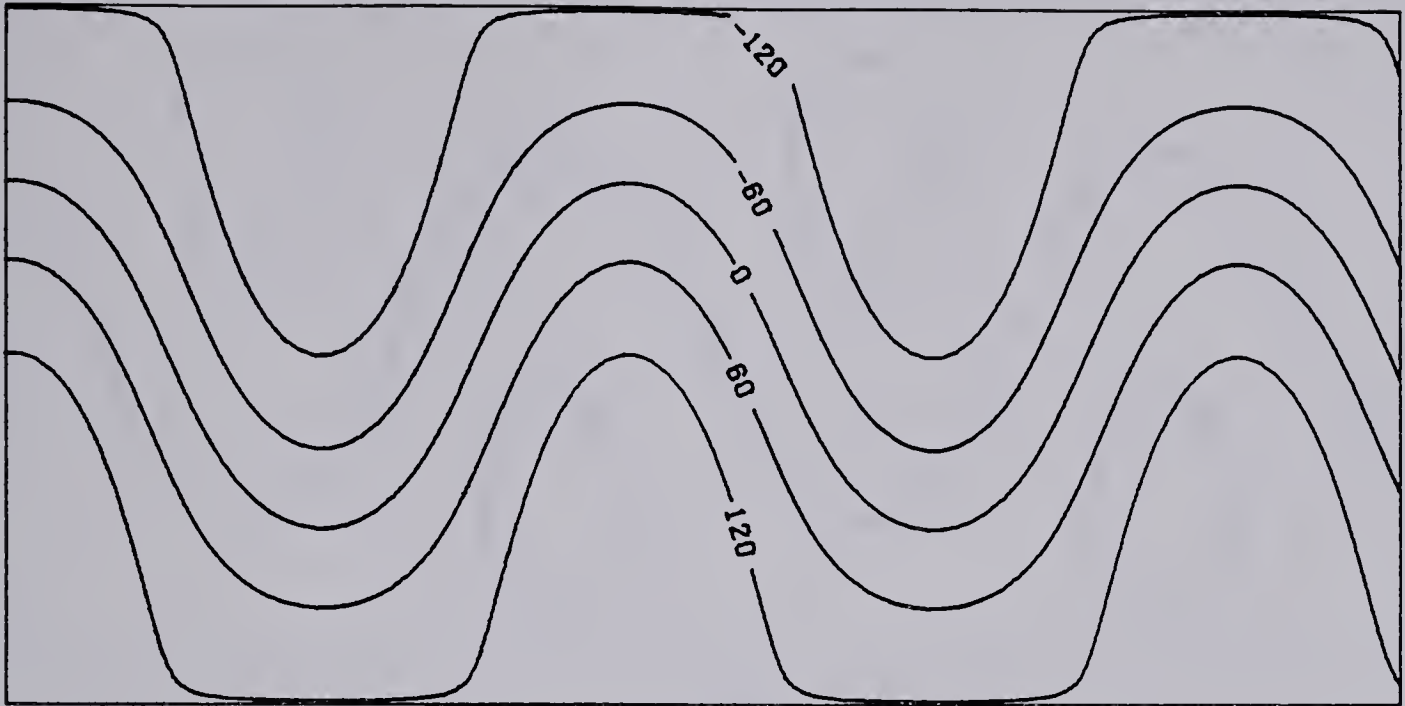


(a)

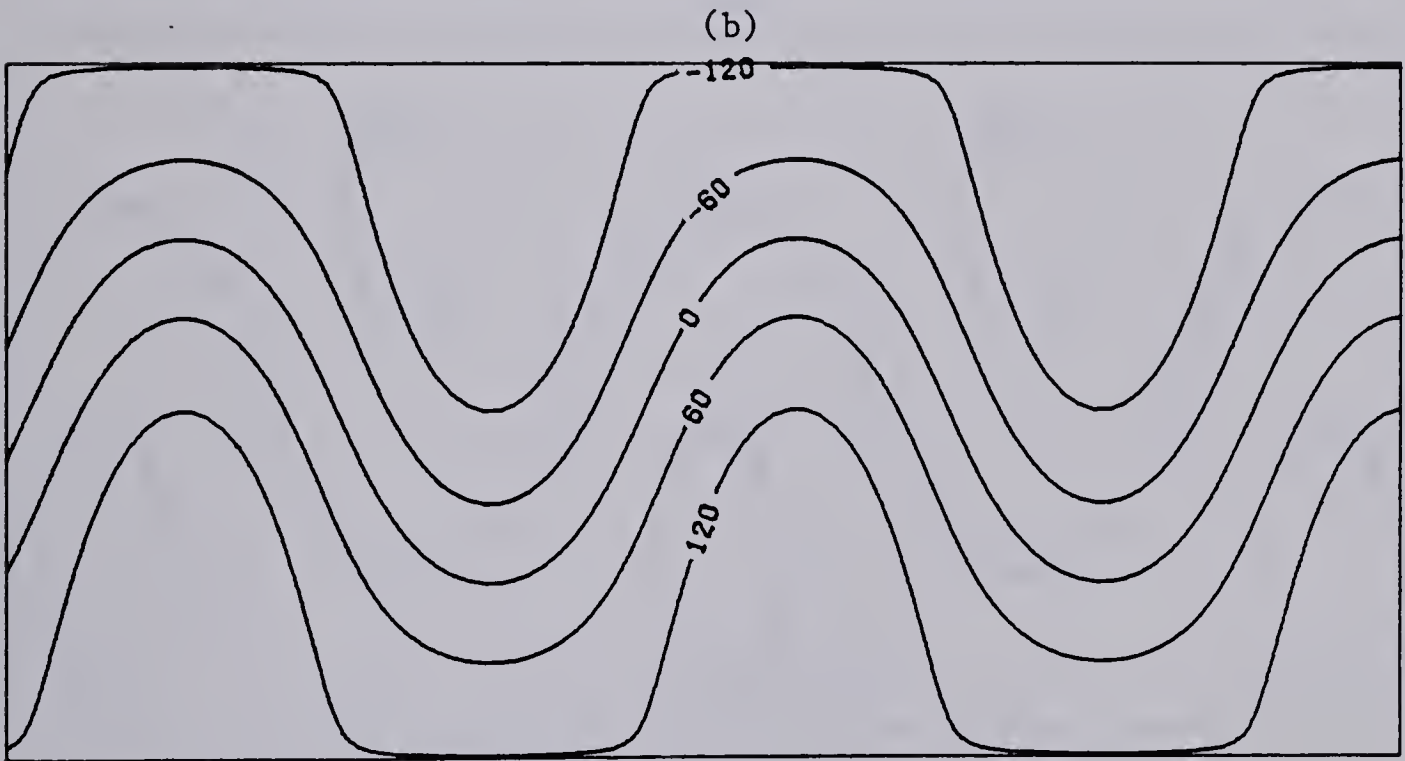


(b)

Fig. 5.15 Initial omega field in $\mu\text{bar}/\text{sec}$ for Case II in (a) area B and (b) area A.

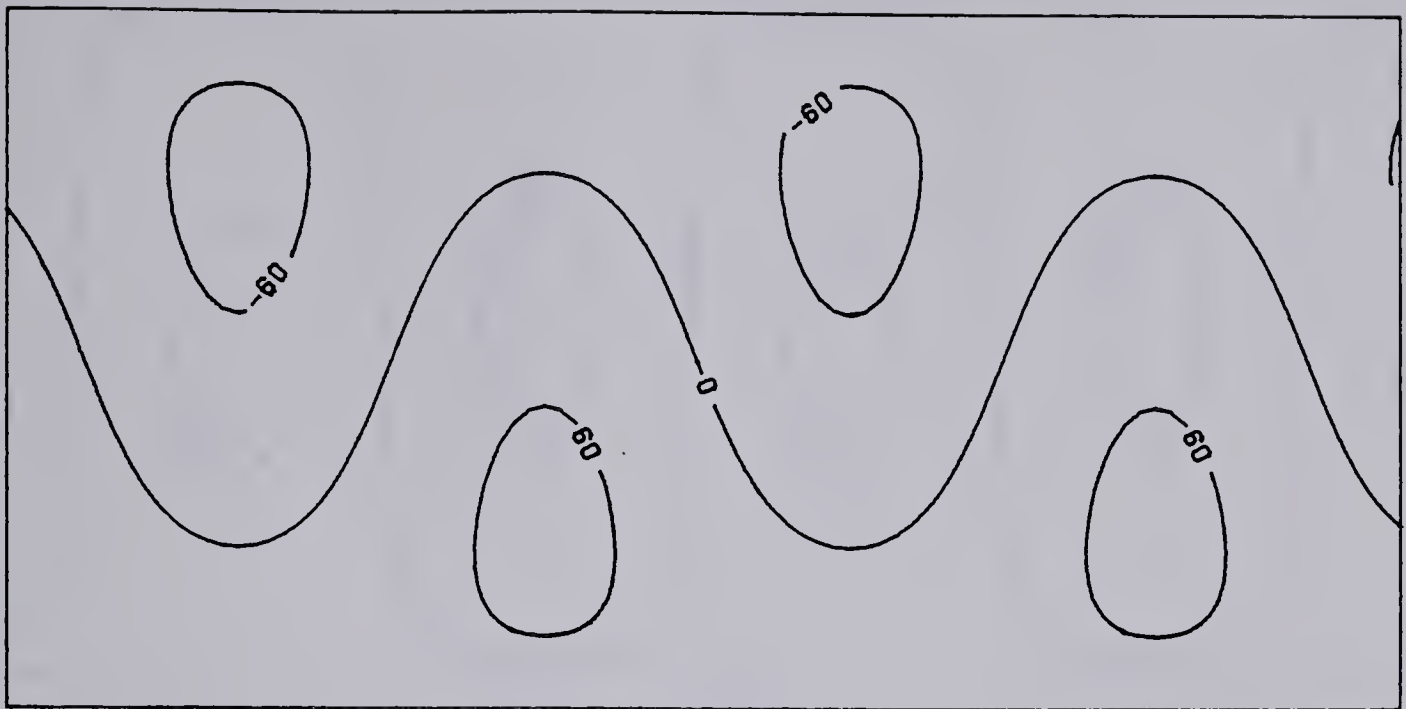


(a)



(b)

Fig. 5.16 The 48-hour spectral solution for the mean height field in metres for Case II in (a) area B and (b) area A.



(a)

(b)

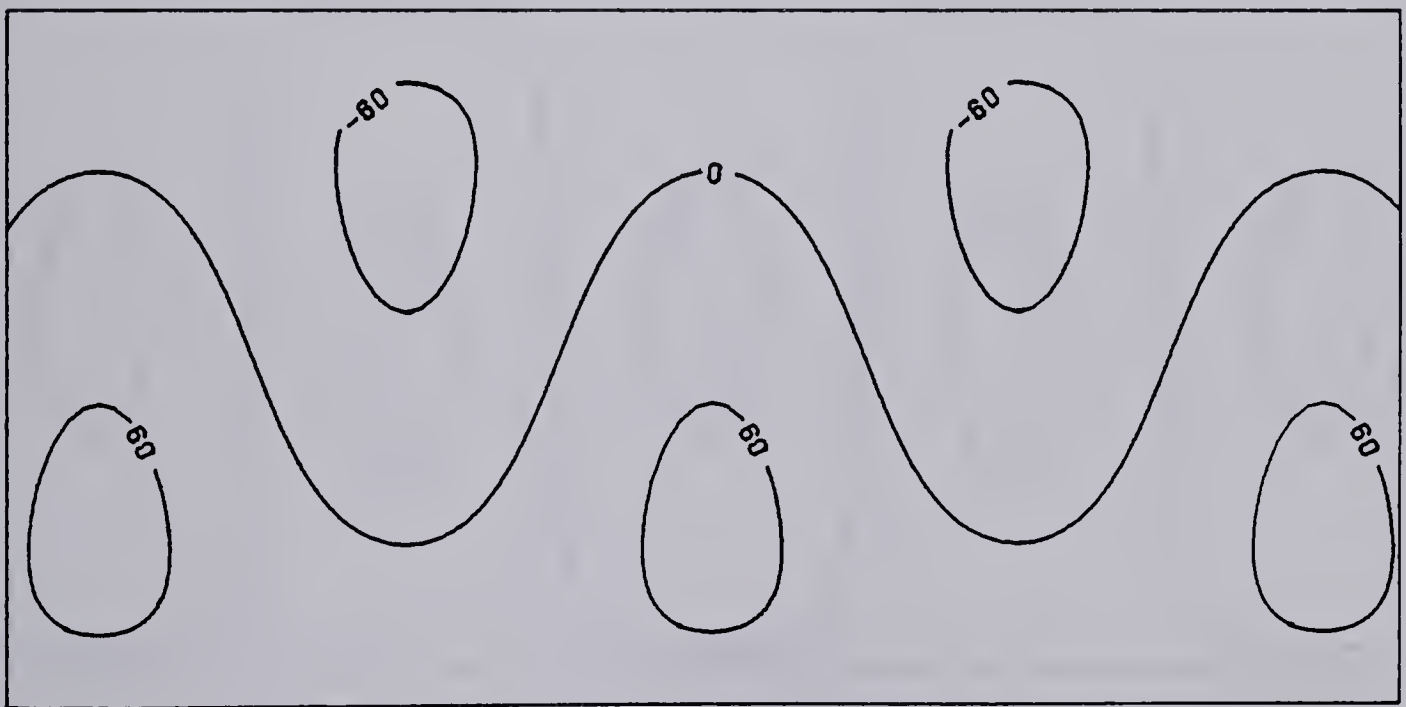
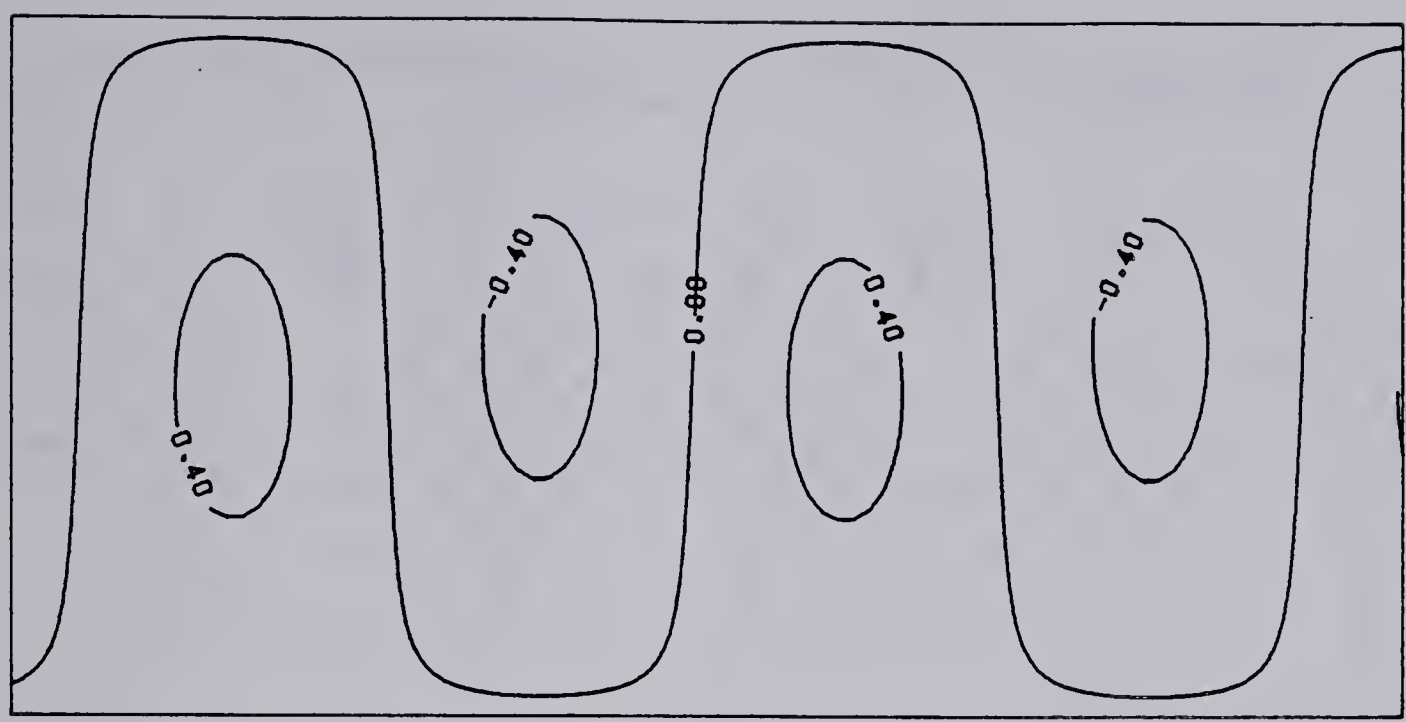
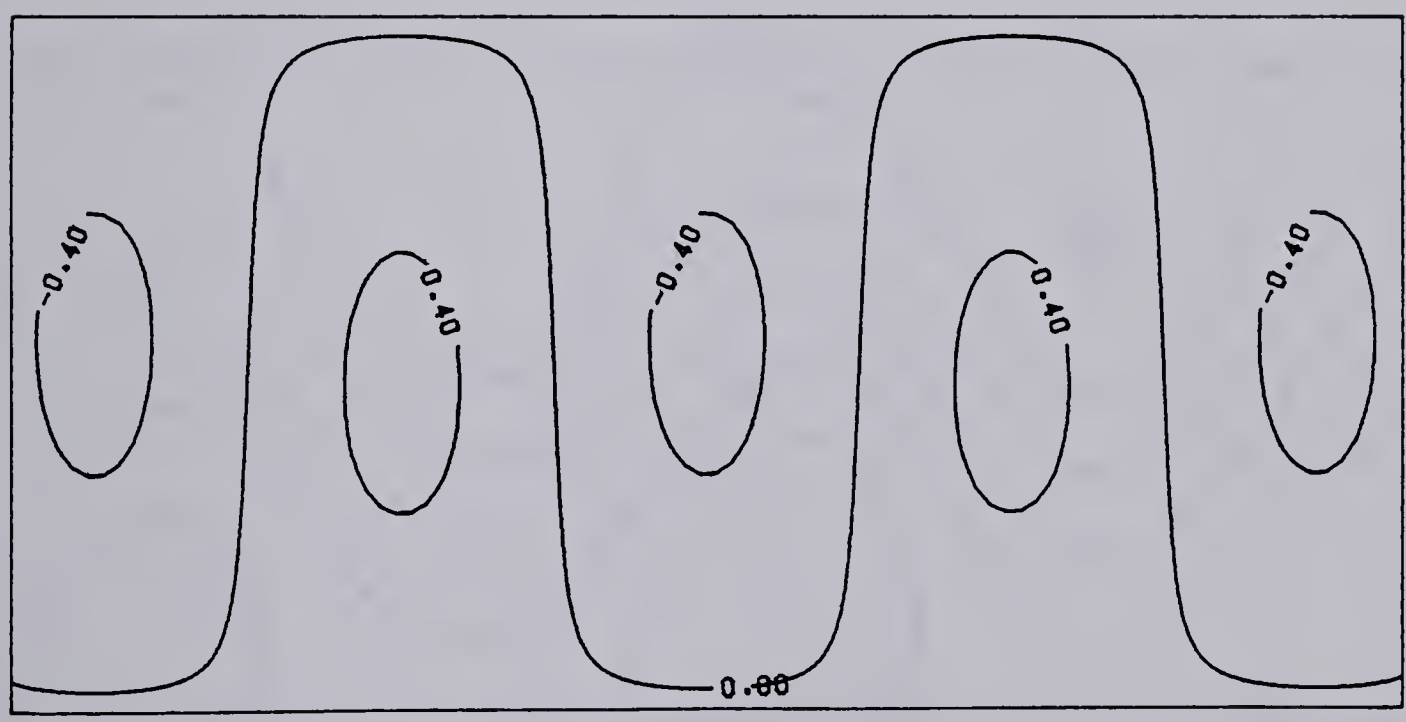


Fig. 5.17 The 48-hour spectral solution for the thickness field in metres for Case II in (a) area B and (b) area A.

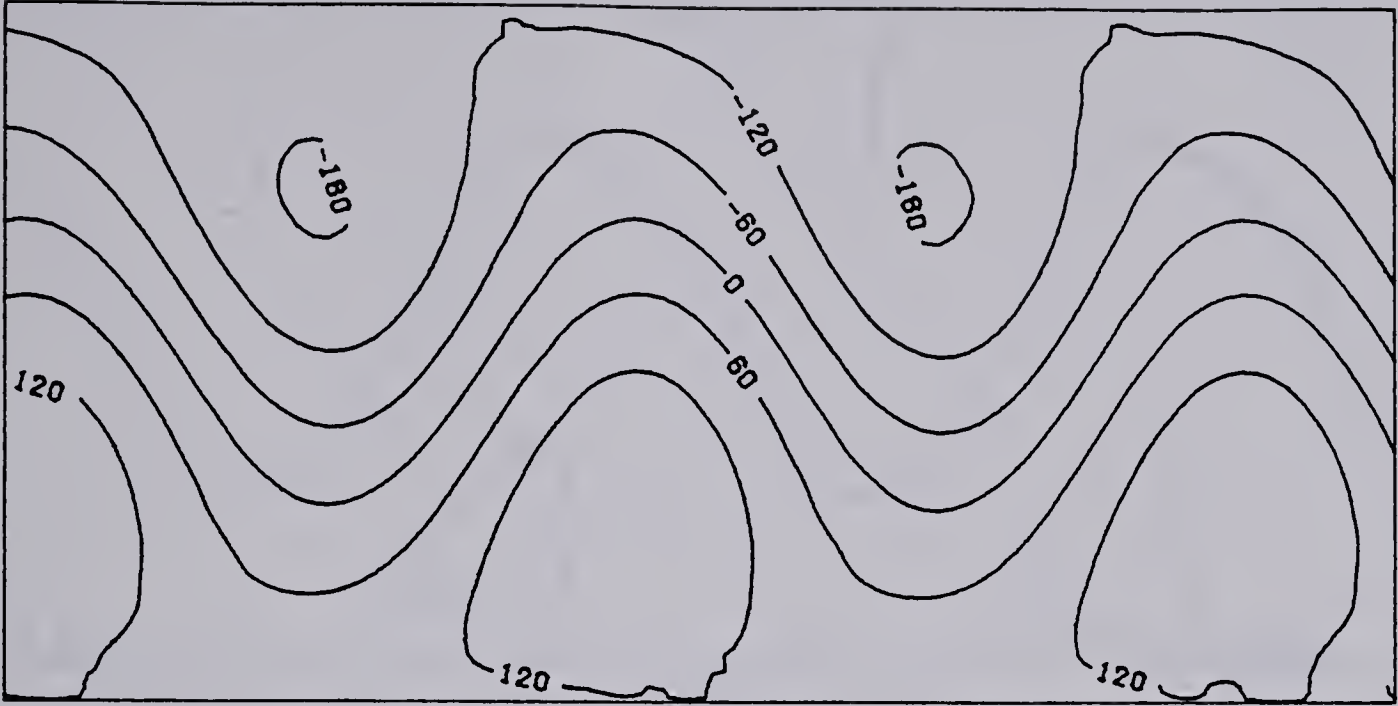


(a)

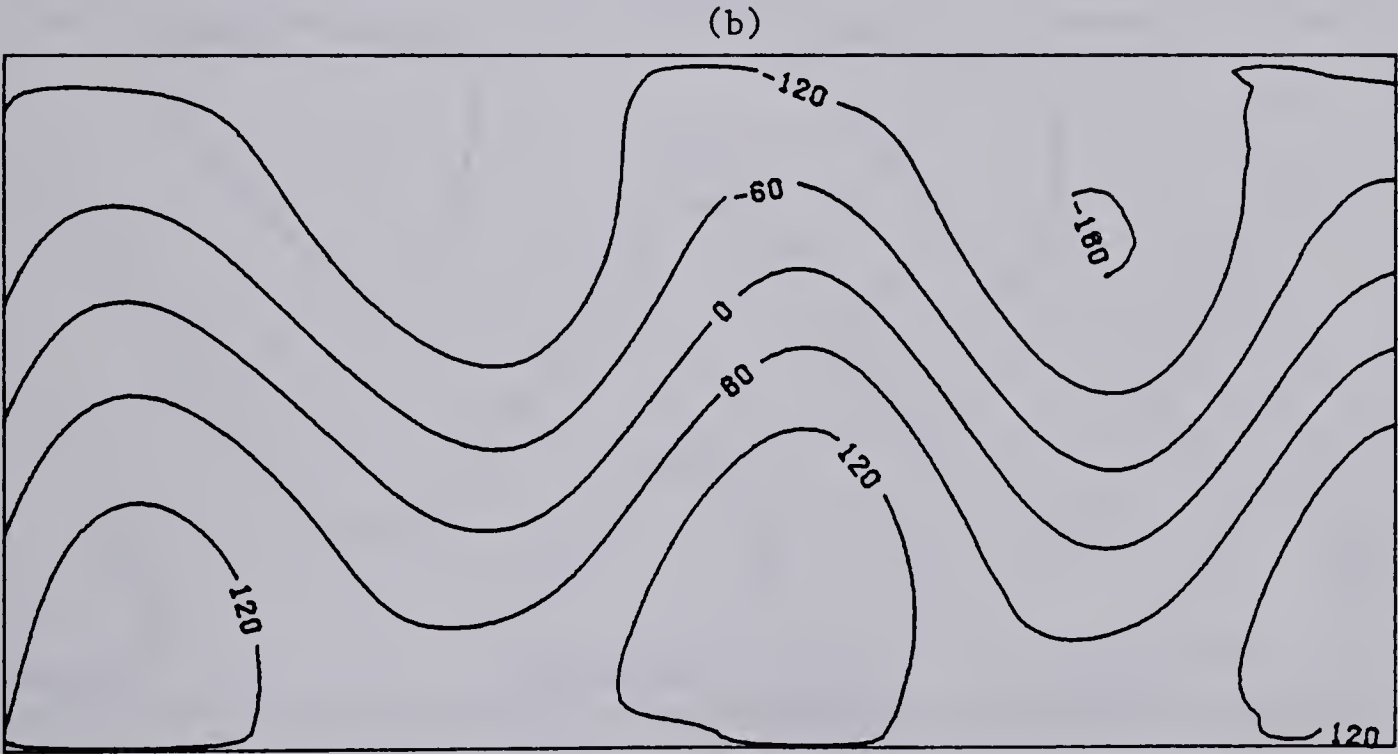


(b)

Fig. 5.18 The 48-hour spectral solution for the omega field in $\mu\text{bar/sec}$ for Case II in (a) area B and (b) area A.

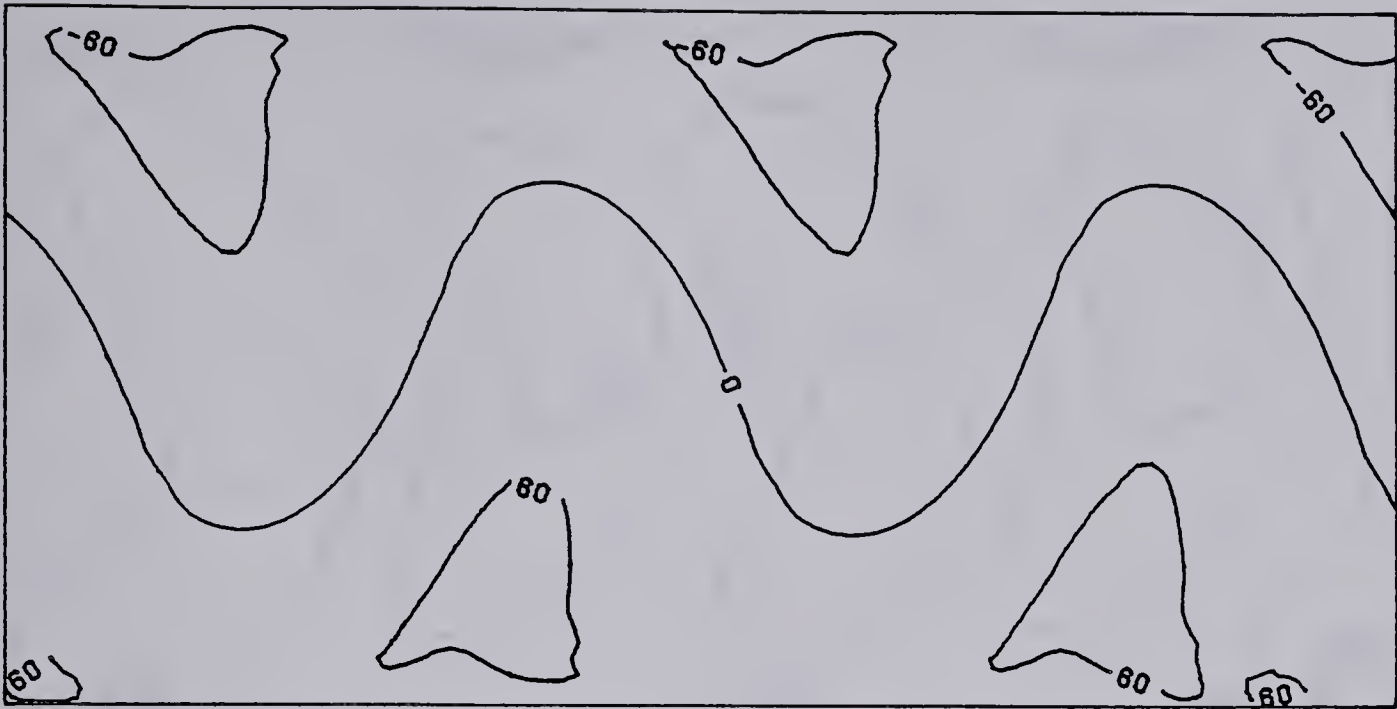


(a)

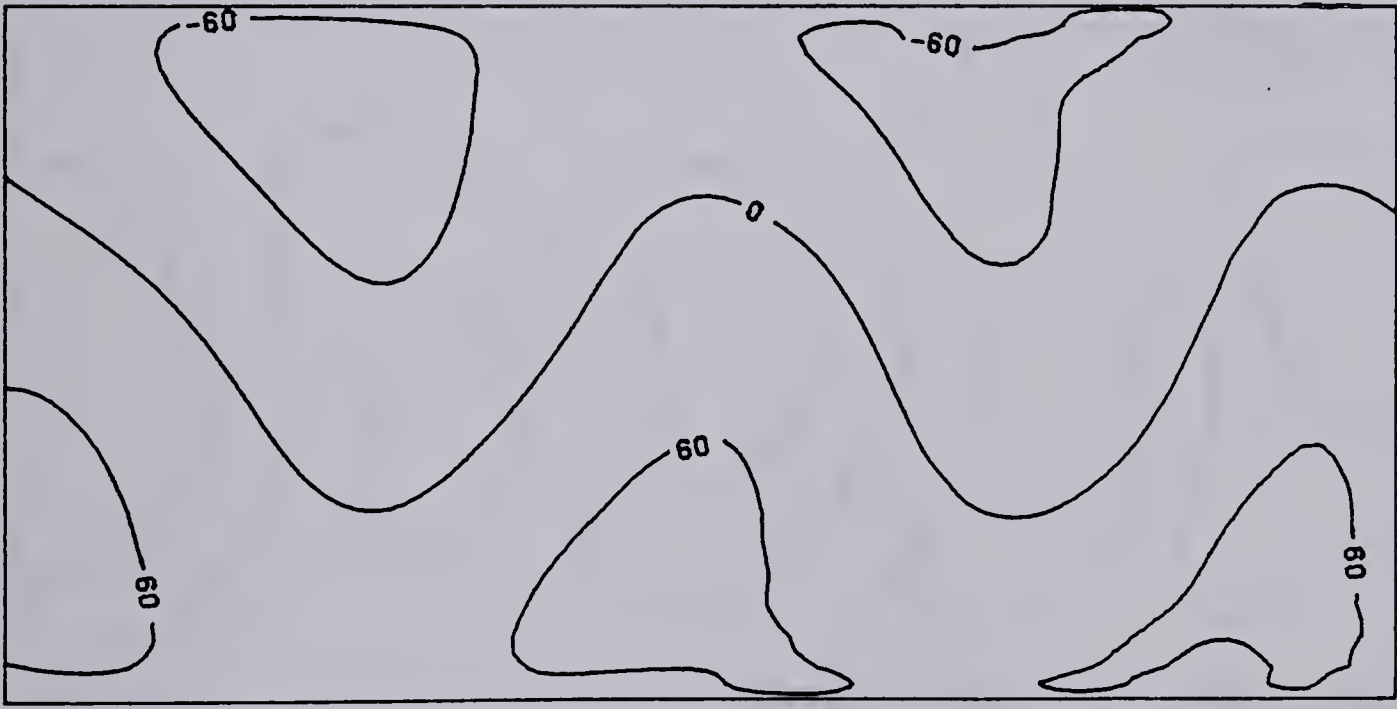


(b)

Fig. 5.19 The 48-hour finite-element solution for the mean height field in metres for Case II in (a) area B and (b) area A.

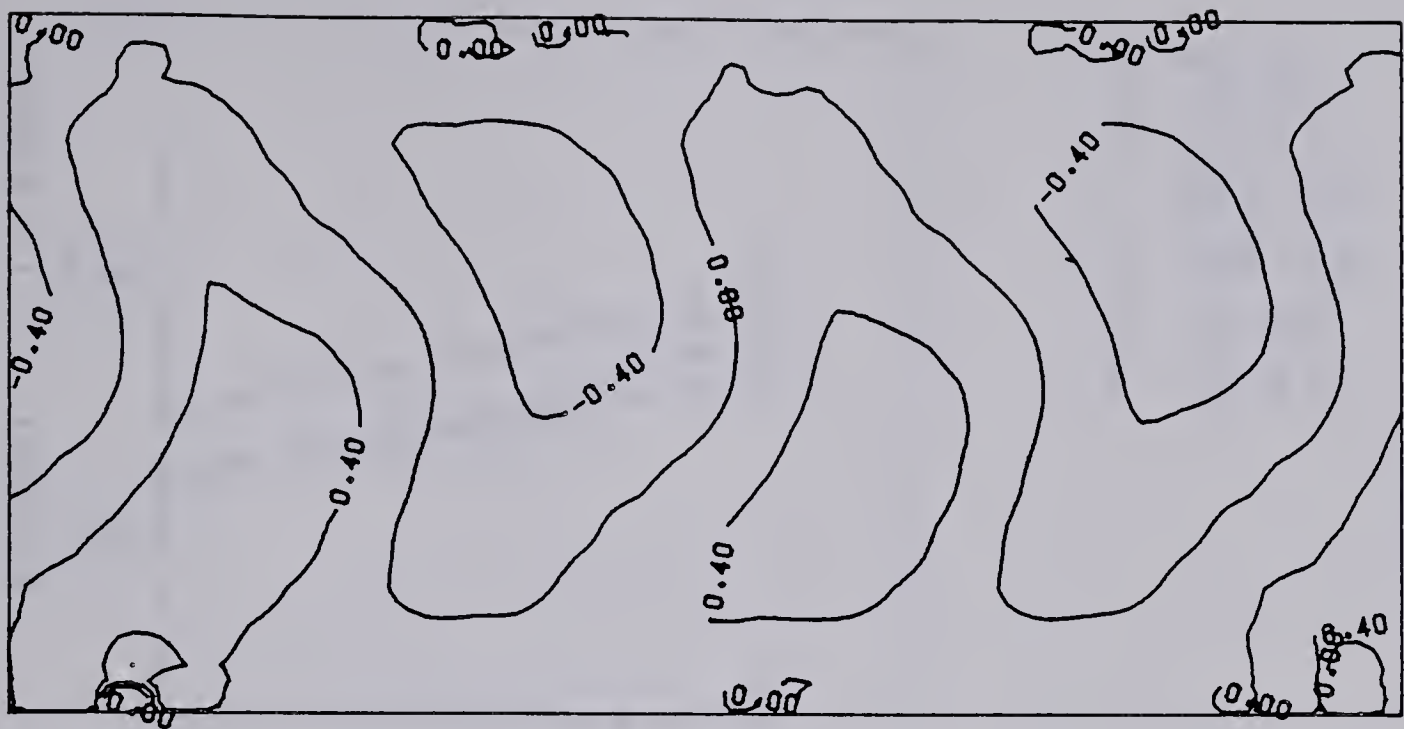


(a)



(b)

Fig. 5.20 The 48-hour finite-element solution for the thickness field in metres for Case II in (a) area B and (b) area A.



(a)

(b)

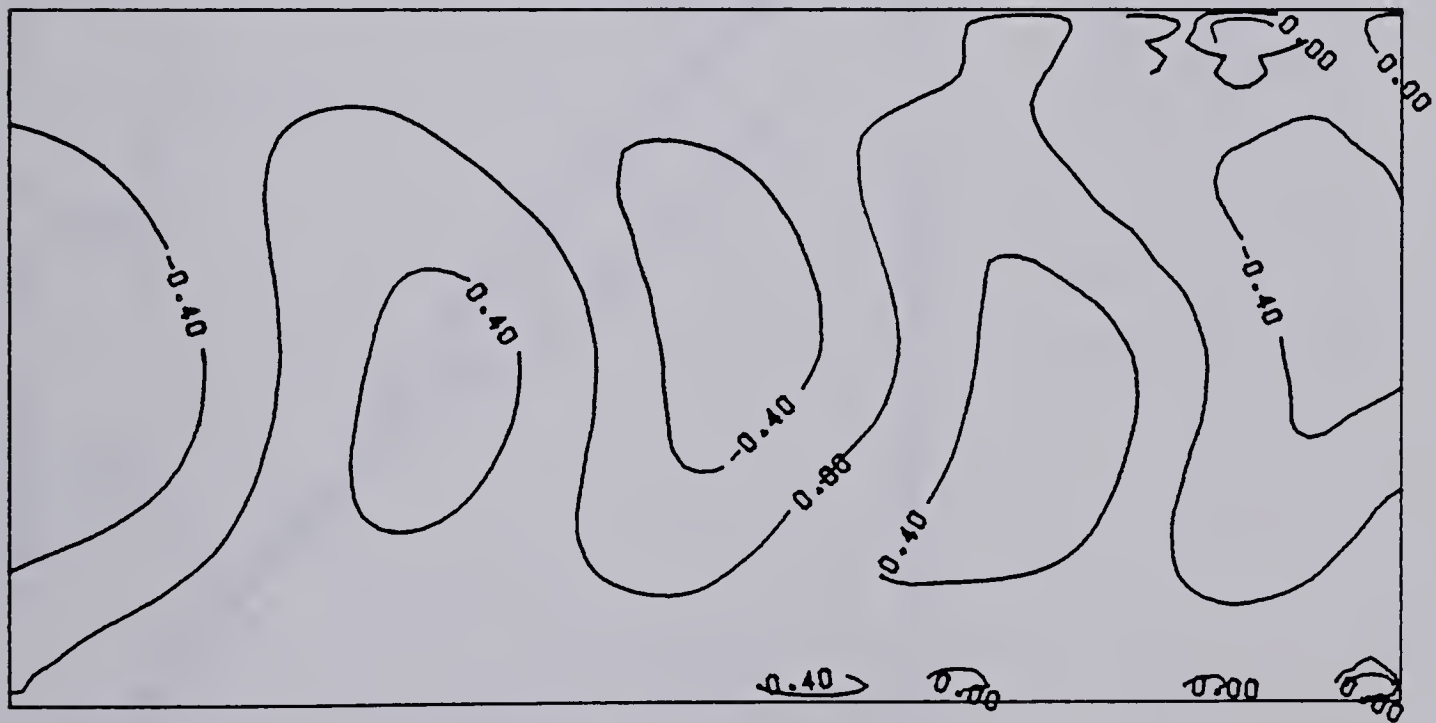
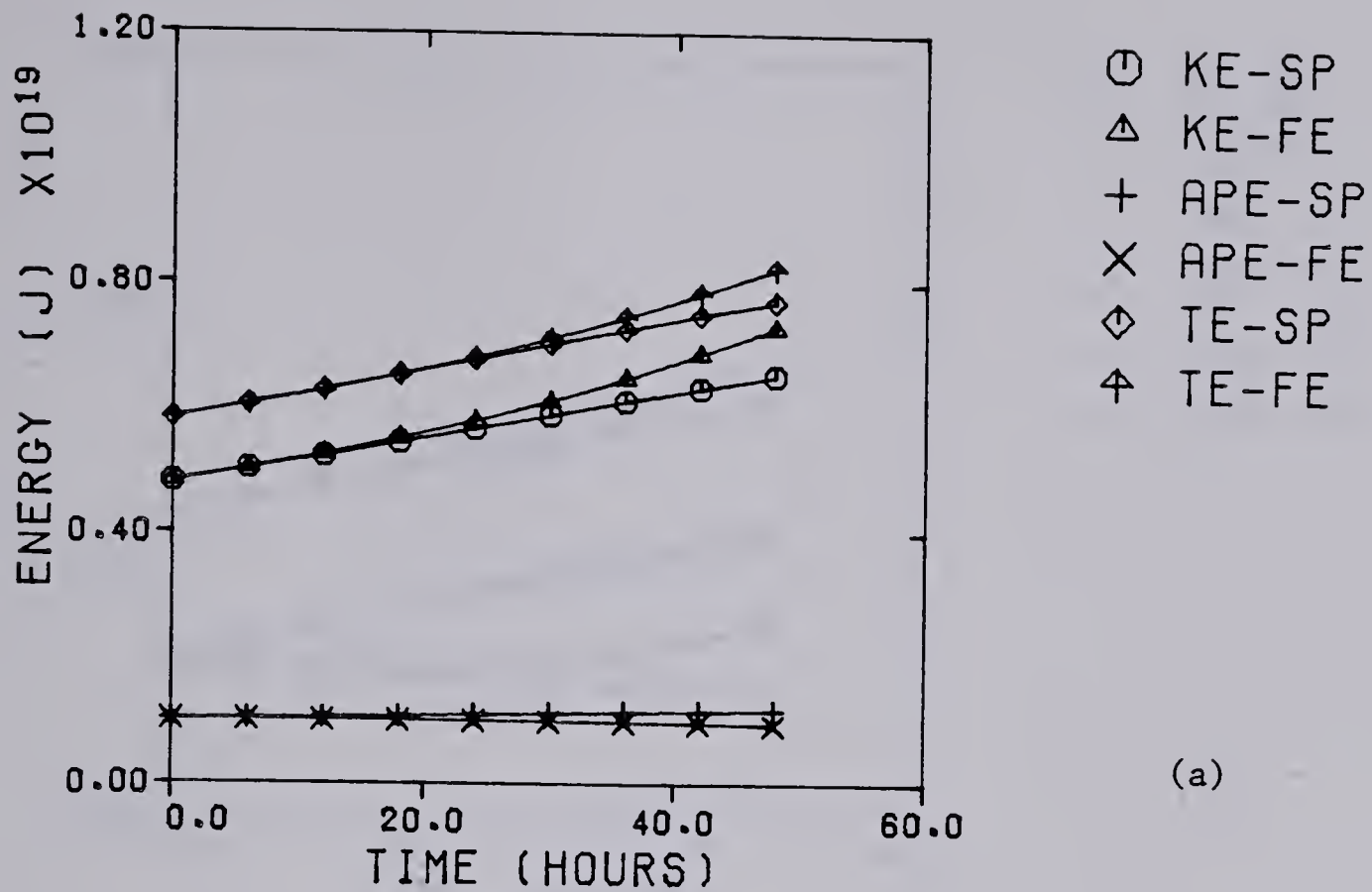
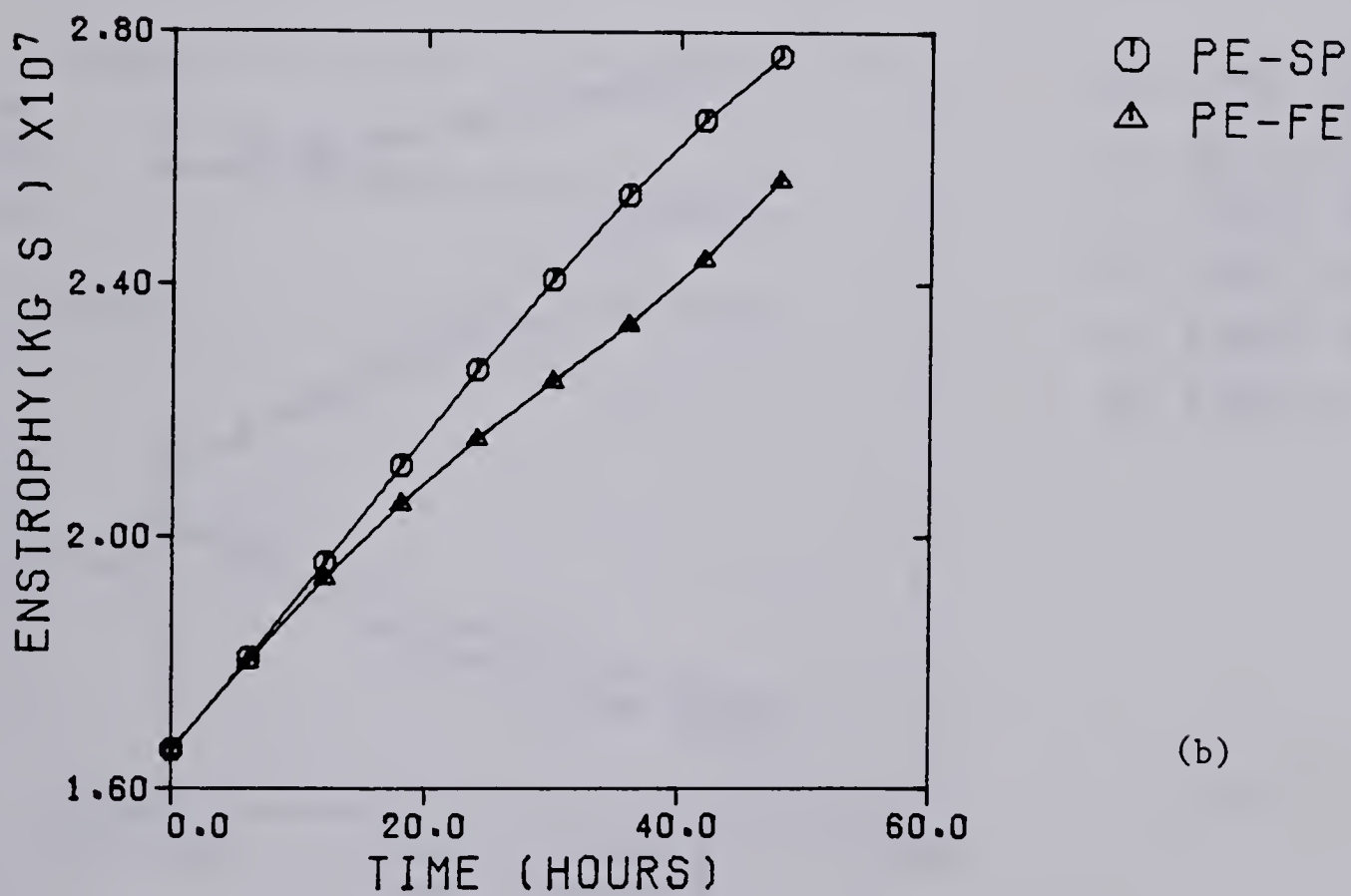


Fig. 5.21 The 48-hour finite-element solution for the omega field in $\mu\text{bar}/\text{sec}$ for Case II in (a) area B and (b) area A.

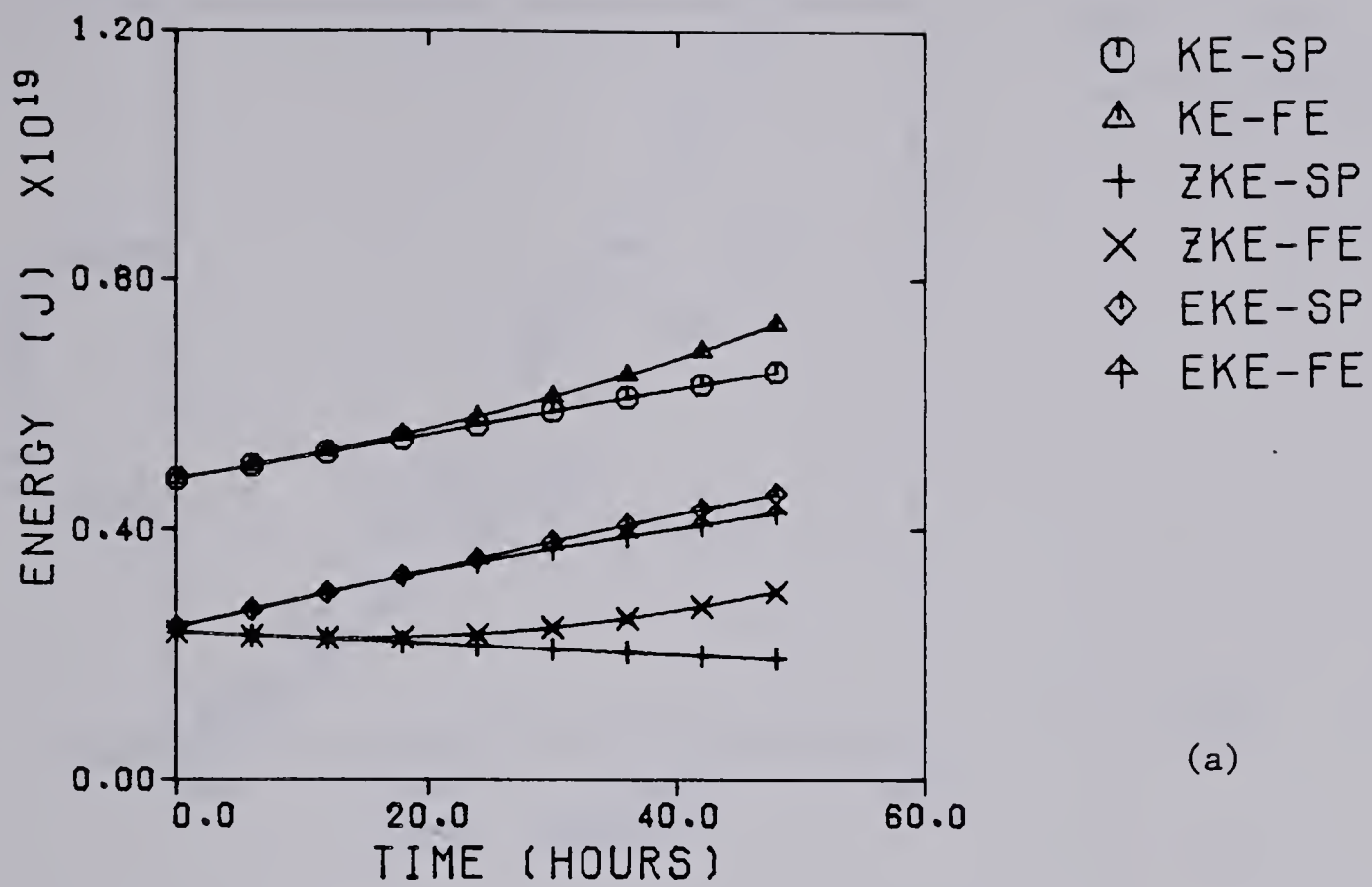


(a)

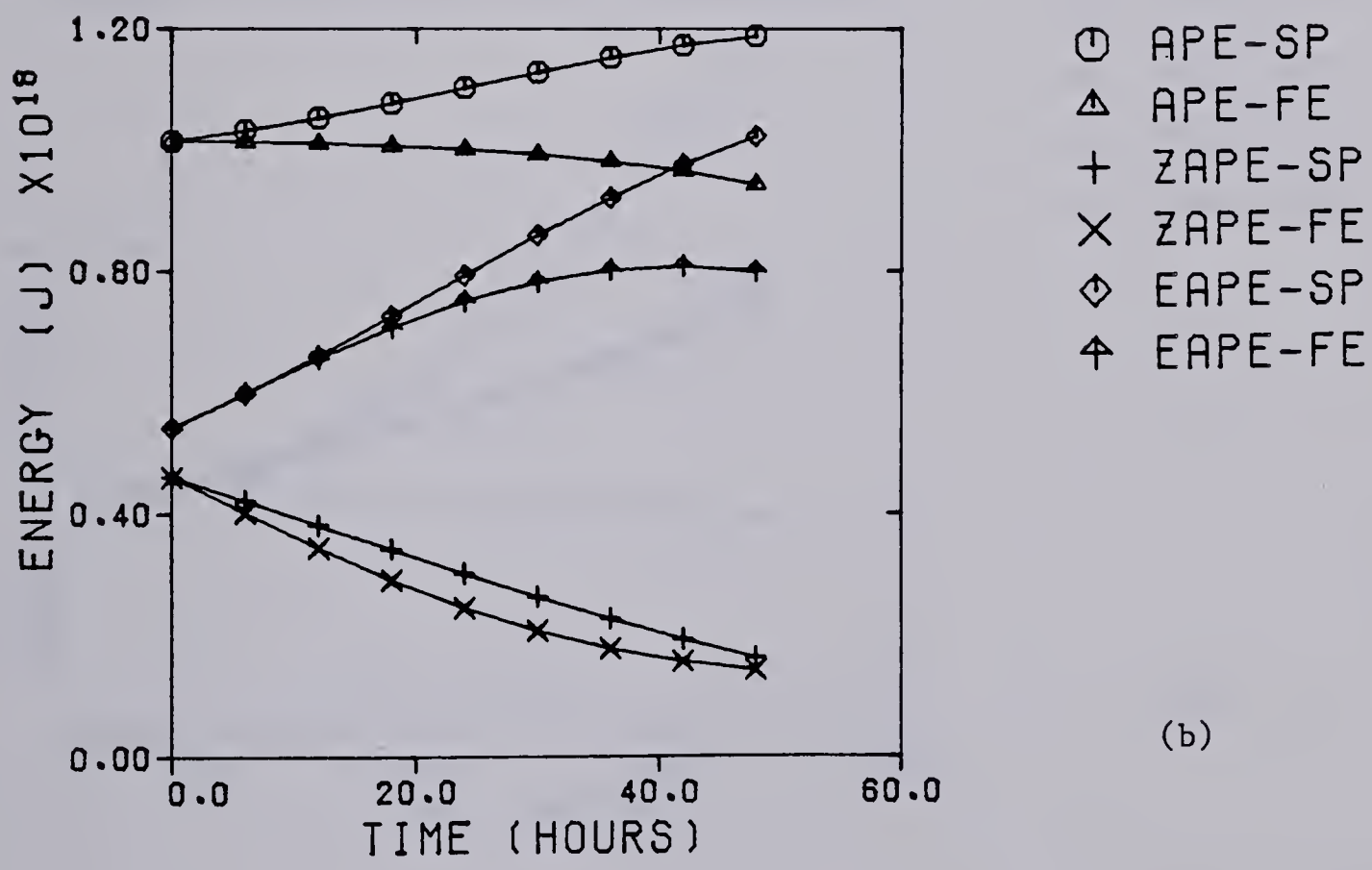


(b)

Fig. 5.22 The (a) energies and (b) potential enstrophy of Case II for the two solutions in the verification area.

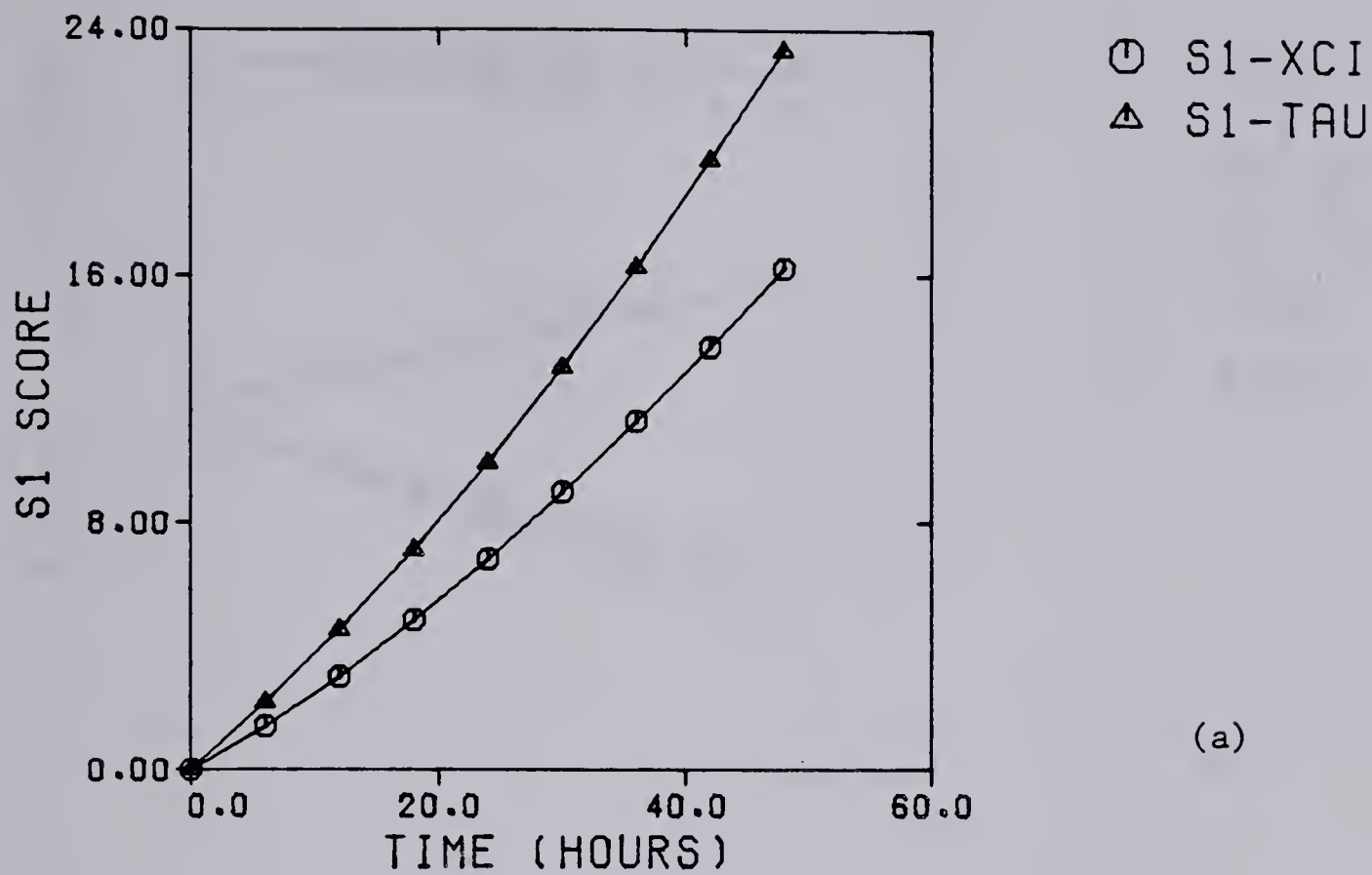


(a)

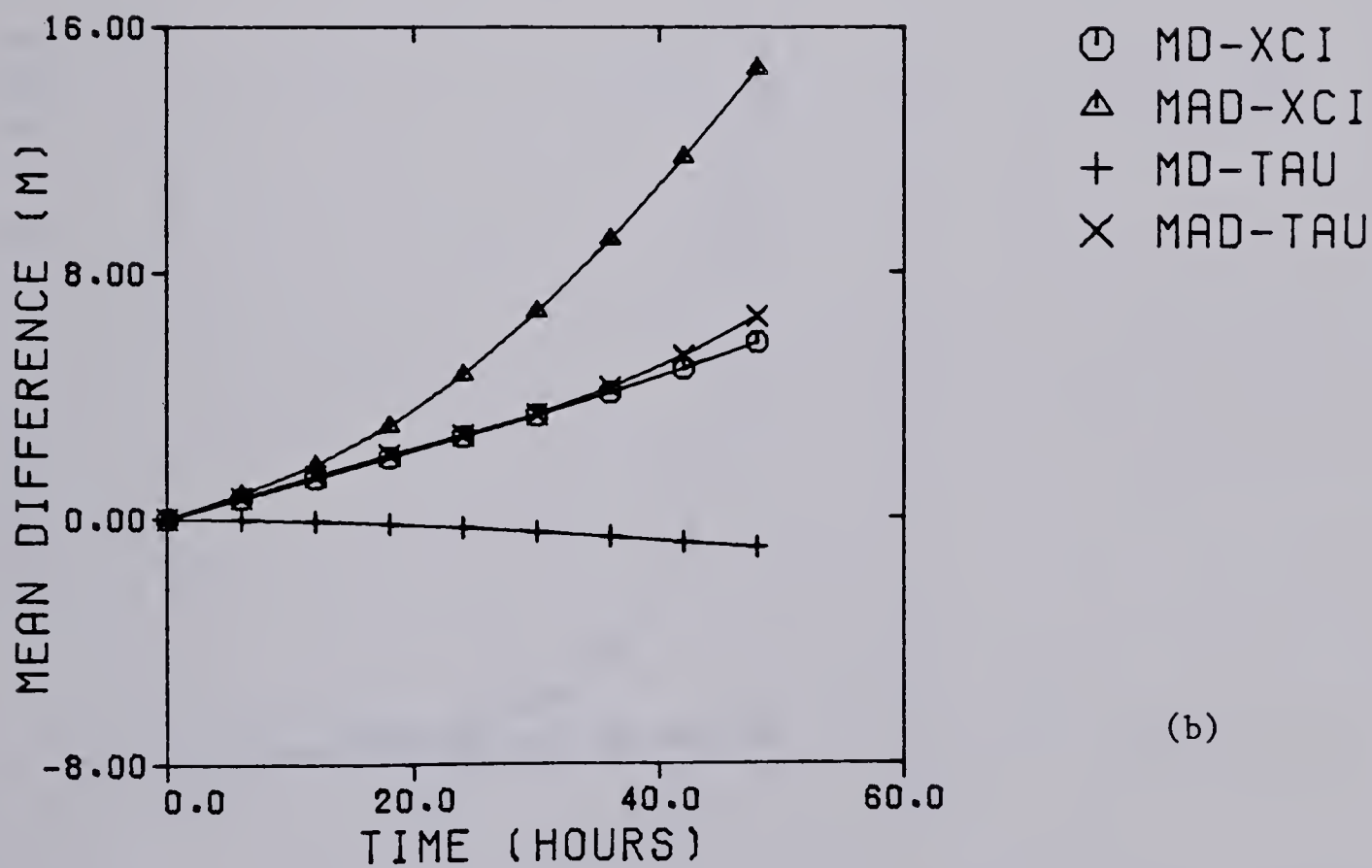


(b)

Fig. 5.23 The (a) kinetic energies and (b) potential energies of Case II for the two solutions in the verification area.

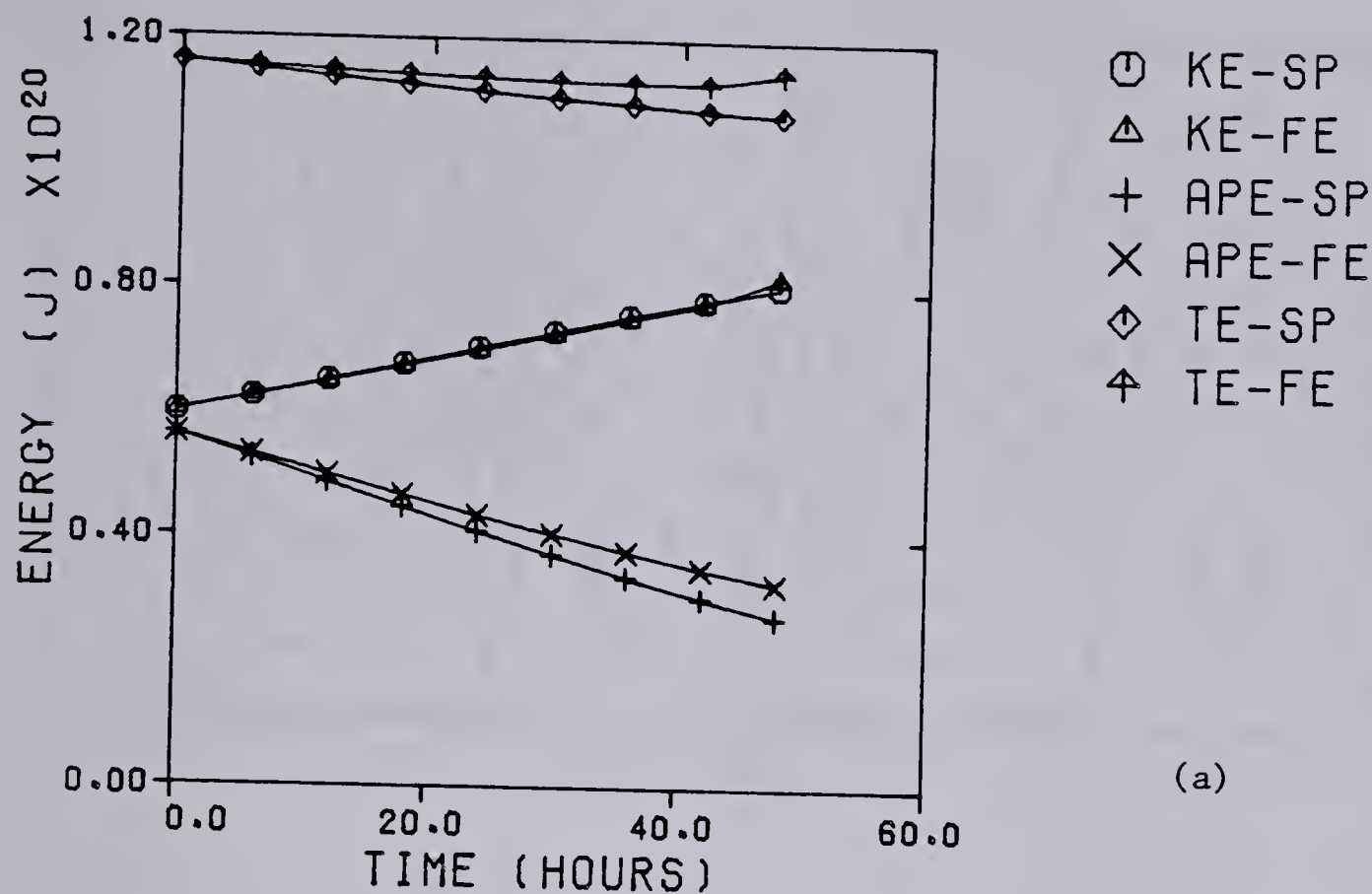


(a)

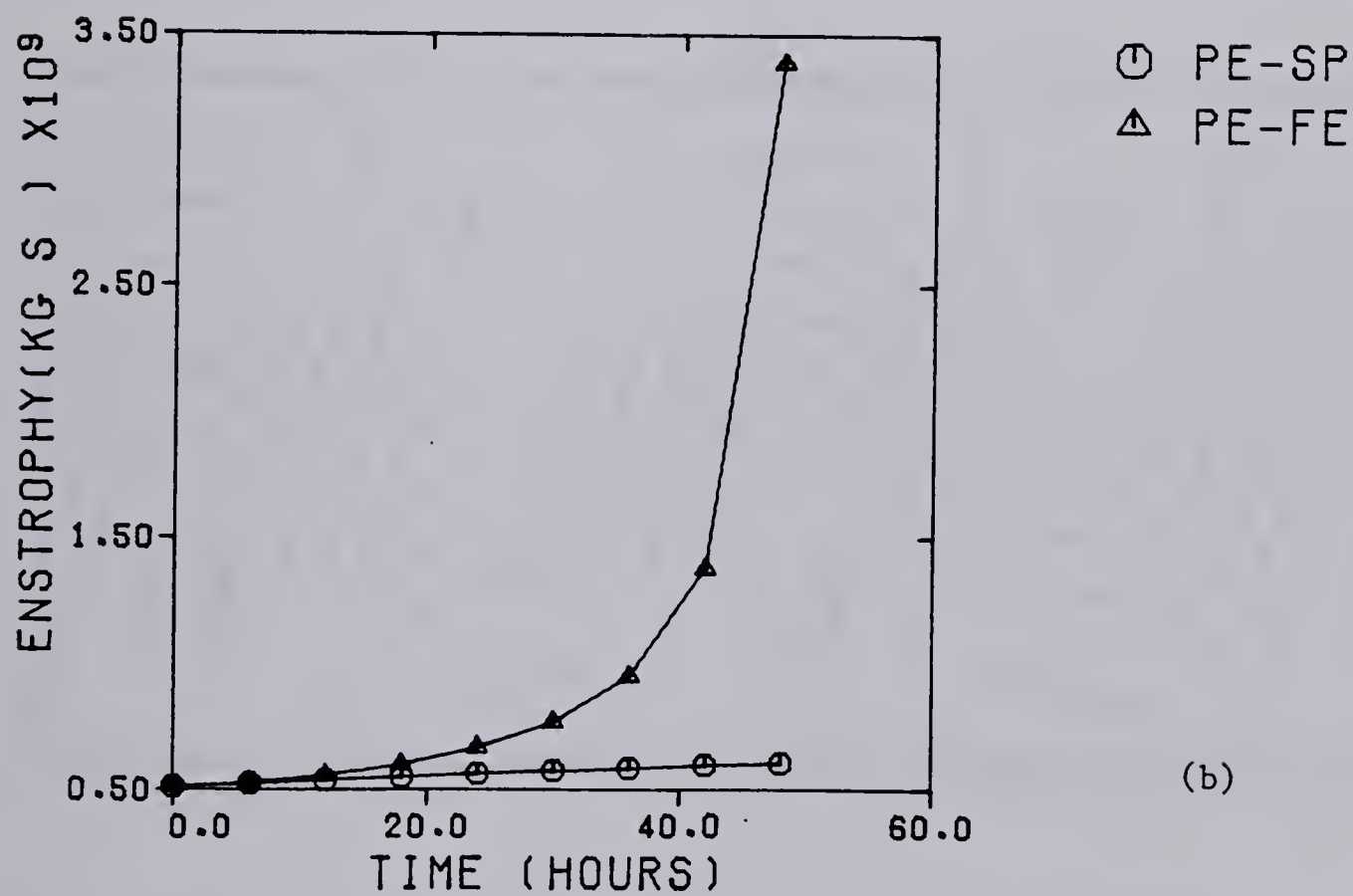


(b)

Fig. 5.24 The (a) S1 scores and (b) MD and MAD curves of Case II for the two solutions in the verification area.

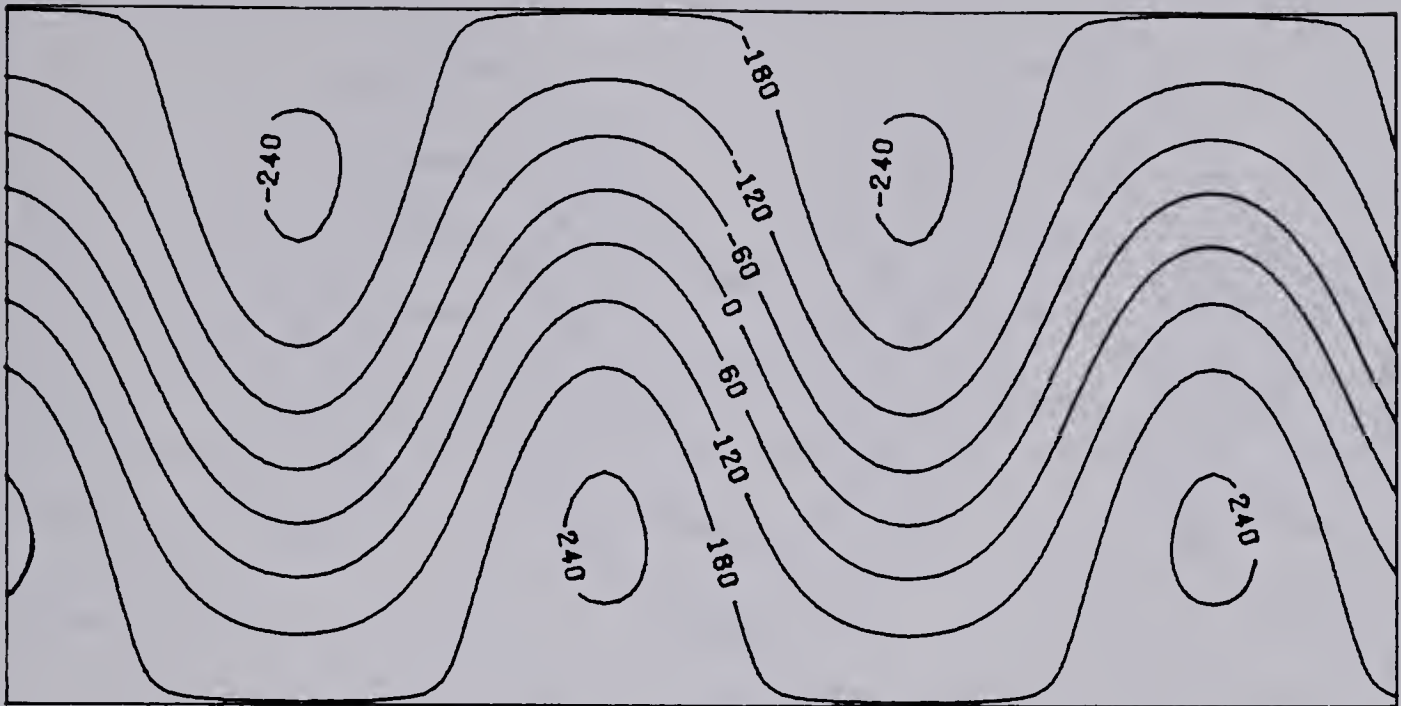


(a)

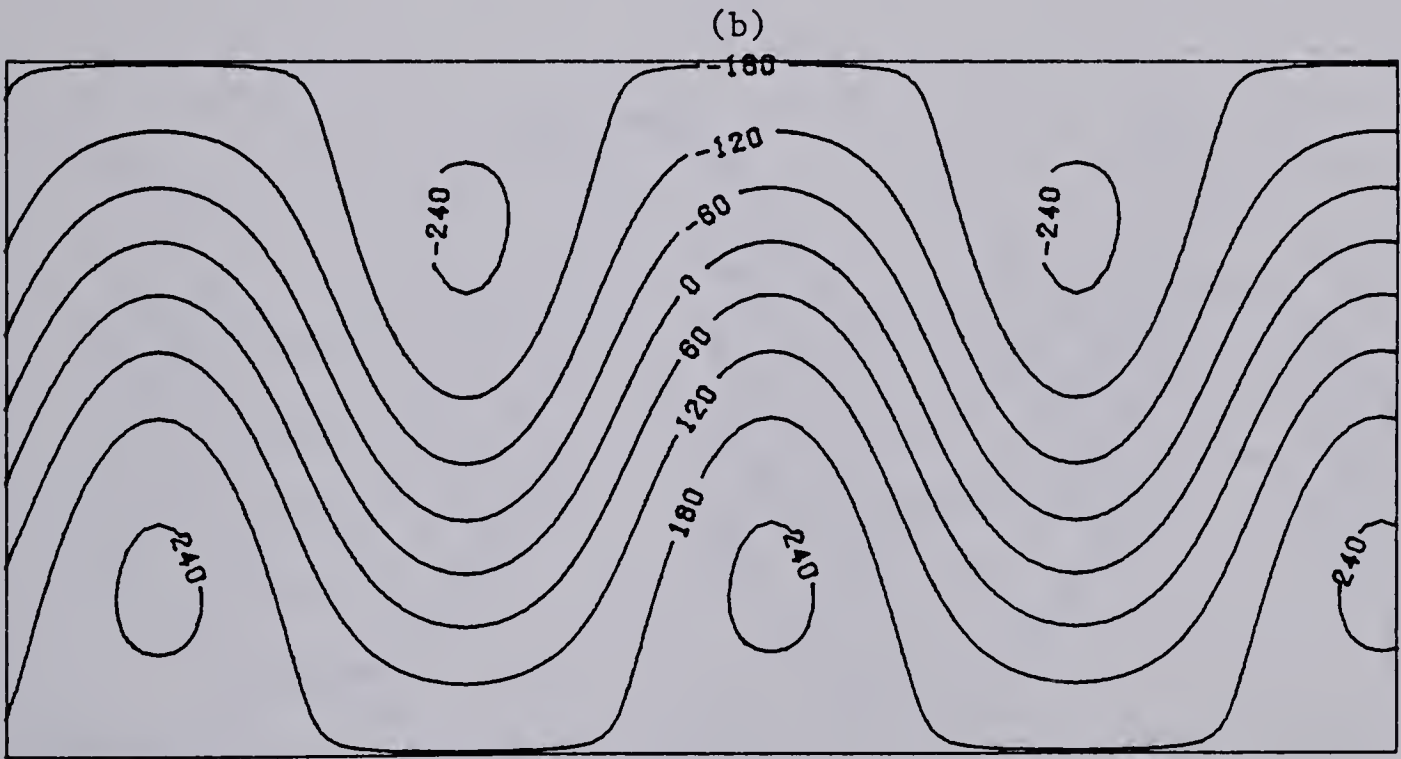


(b)

Fig. 5.25 The (a) energies and (b) potential enstrophy of Case II for the two solutions in the entire domain.

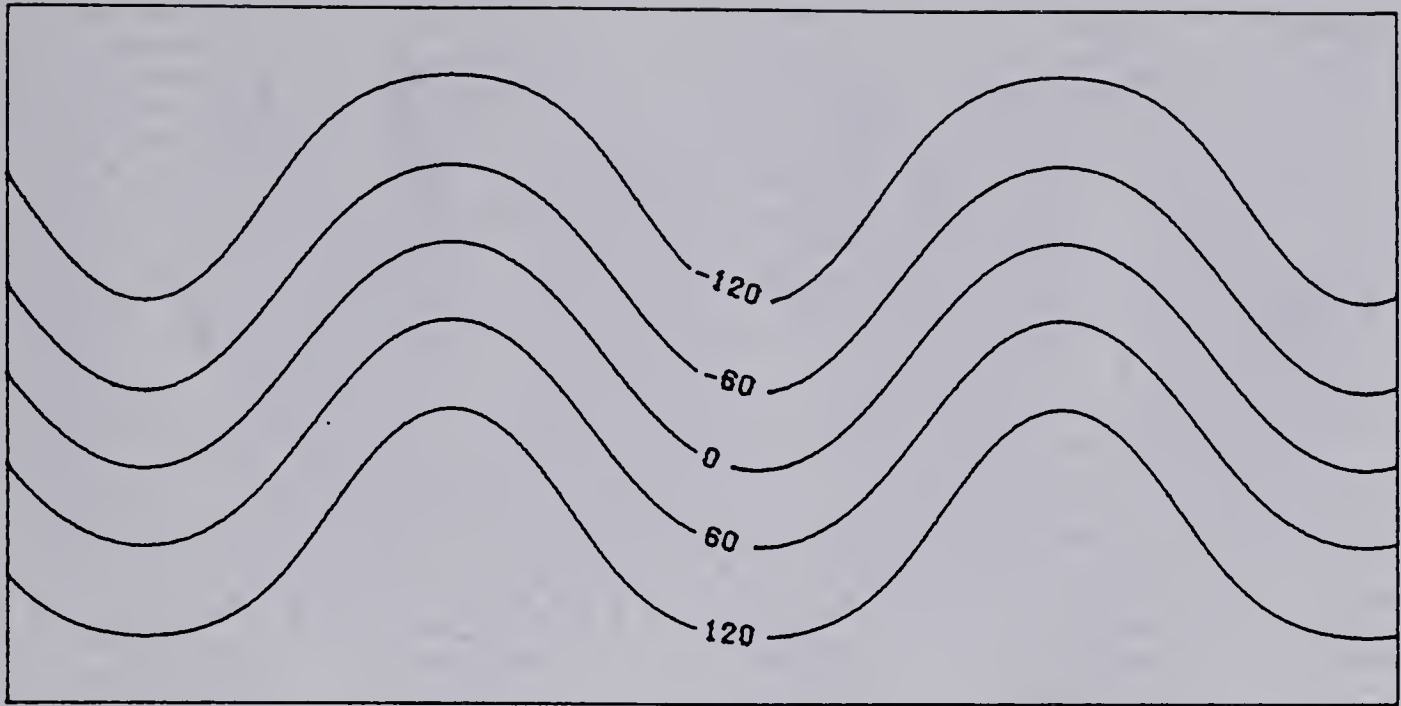


(a)

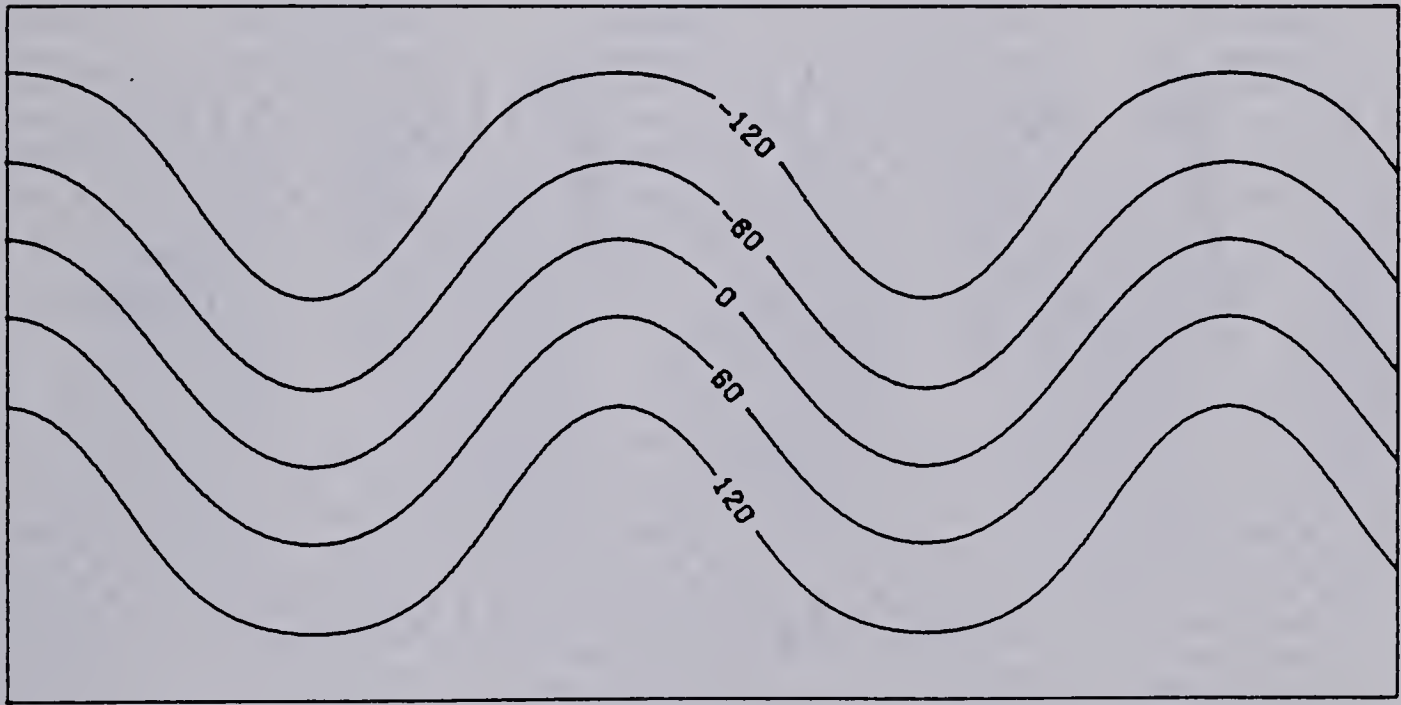


(b)

Fig. 5.26 Initial mean height field in metres for Case III in (a) area B and (b) area A.

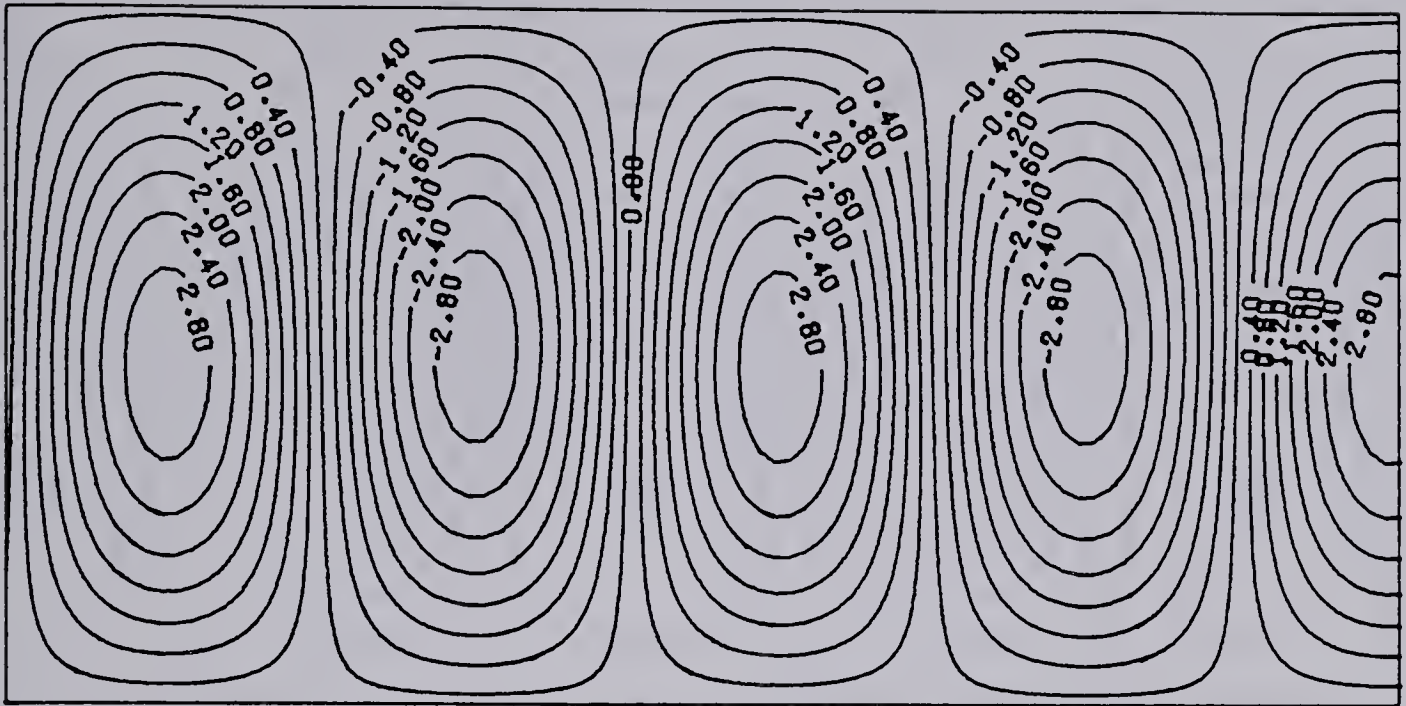


(a)

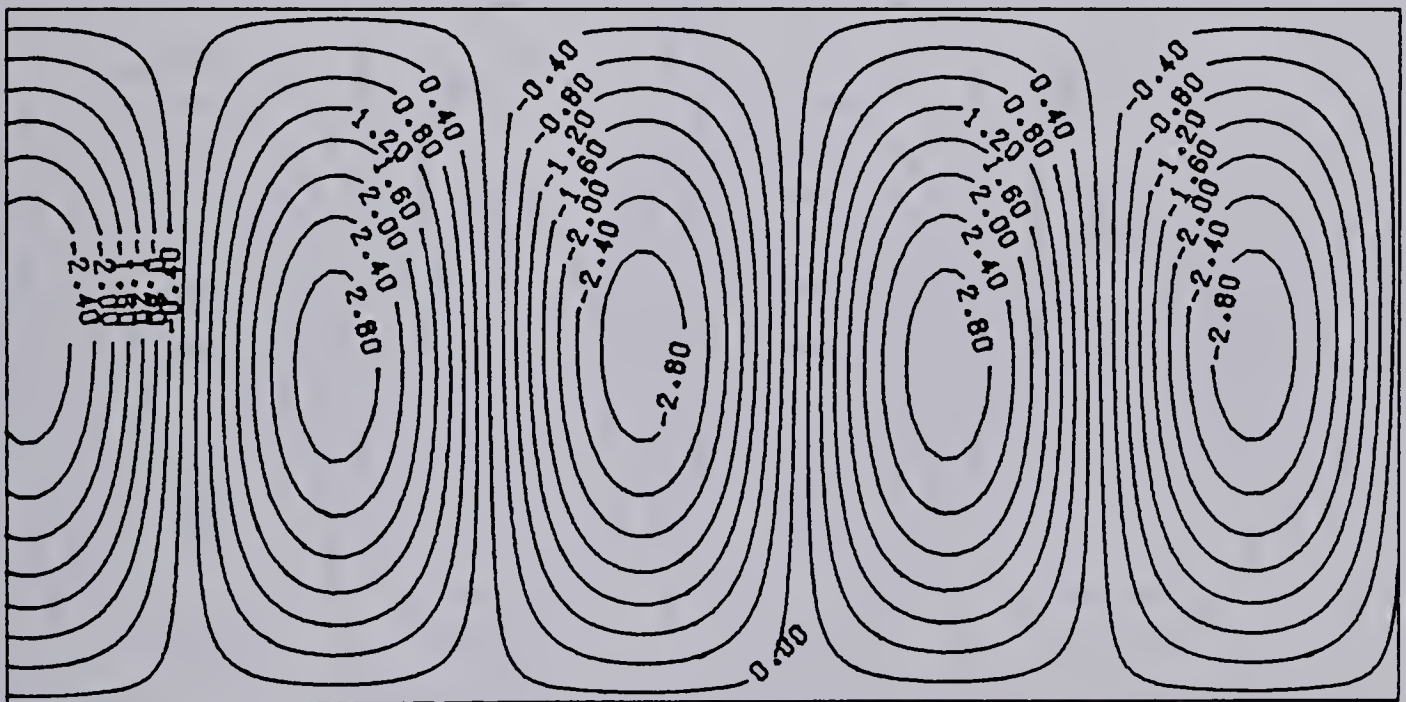


(b)

Fig. 5.27 Initial thickness field in metres for Case III in (a) area B and (b) area A.

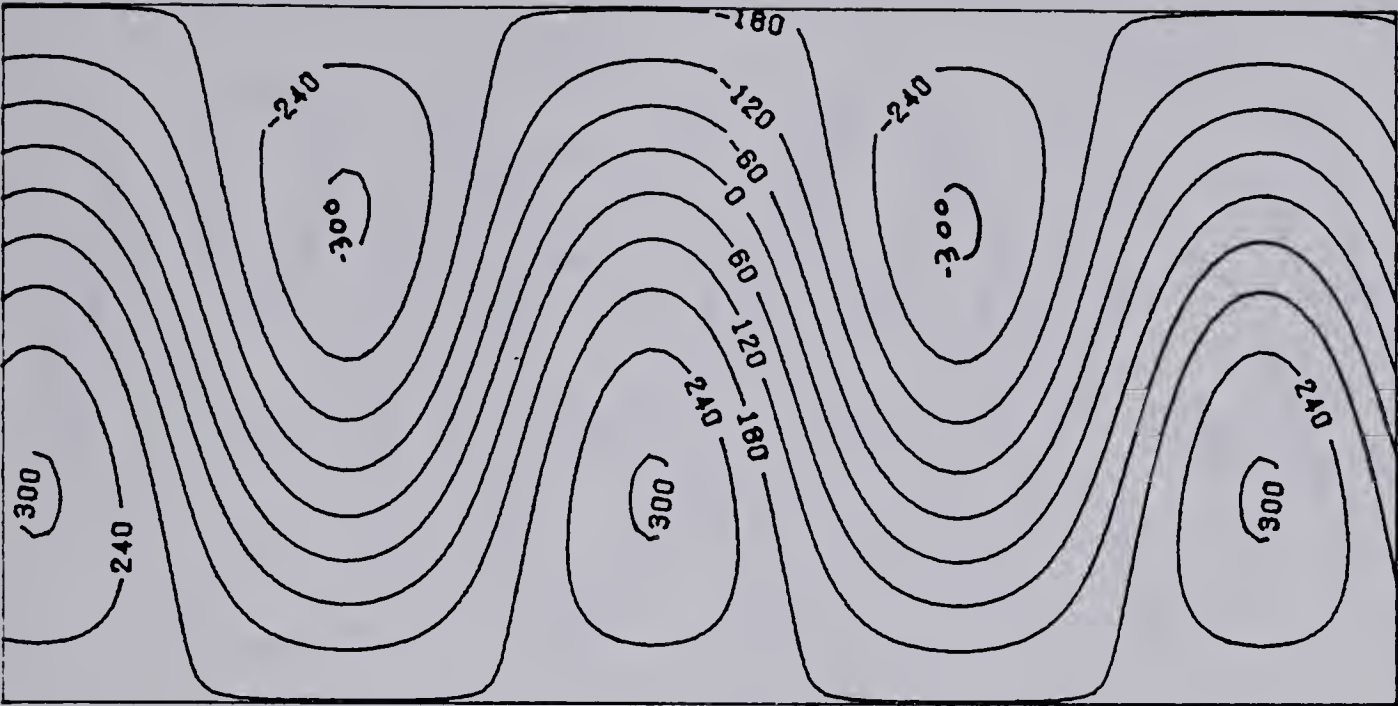


(a)

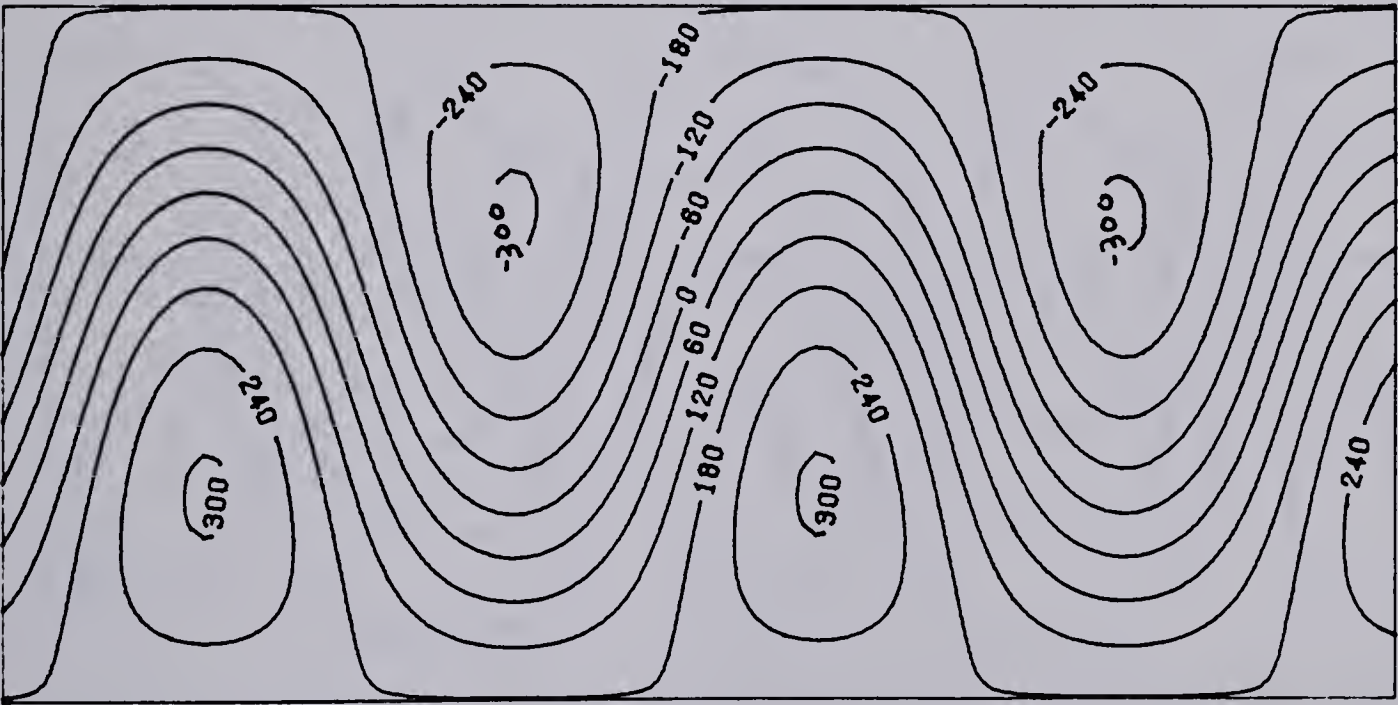


(b)

Fig. 5.28 Initial omega field in $\mu\text{bar}/\text{sec}$ for Case III in
(a) area B and (b) area A.

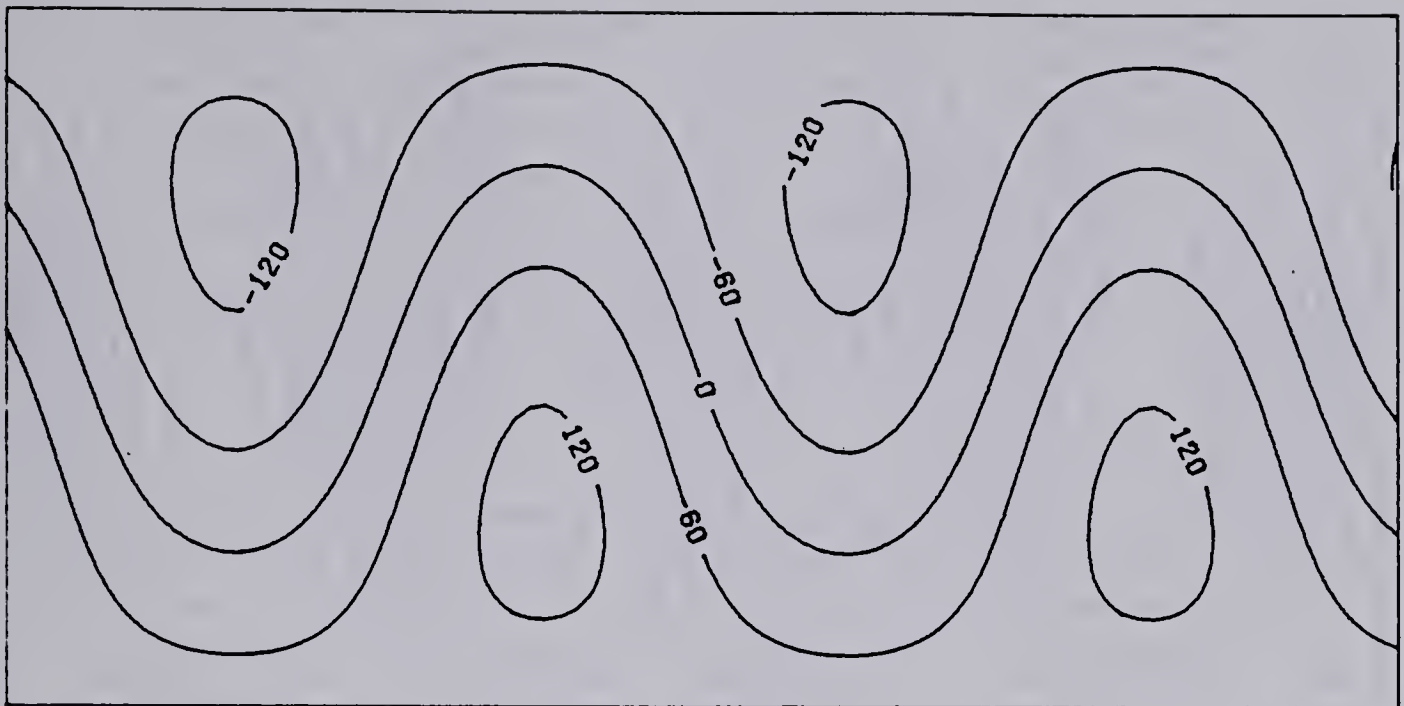


(a)



(b)

Fig. 5.29 The 24-hour spectral solution for the mean height field in metres for Case III in (a) area B and (b) area A.



(a)

(b)

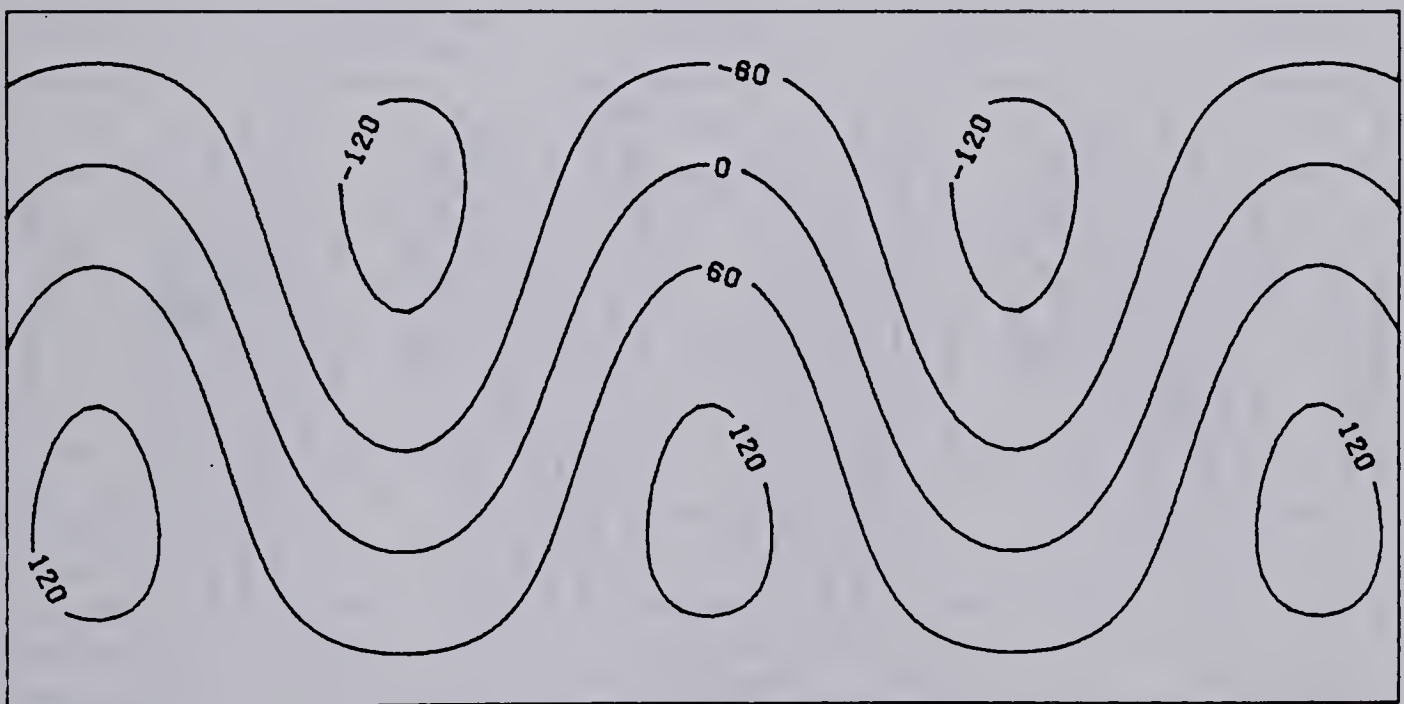
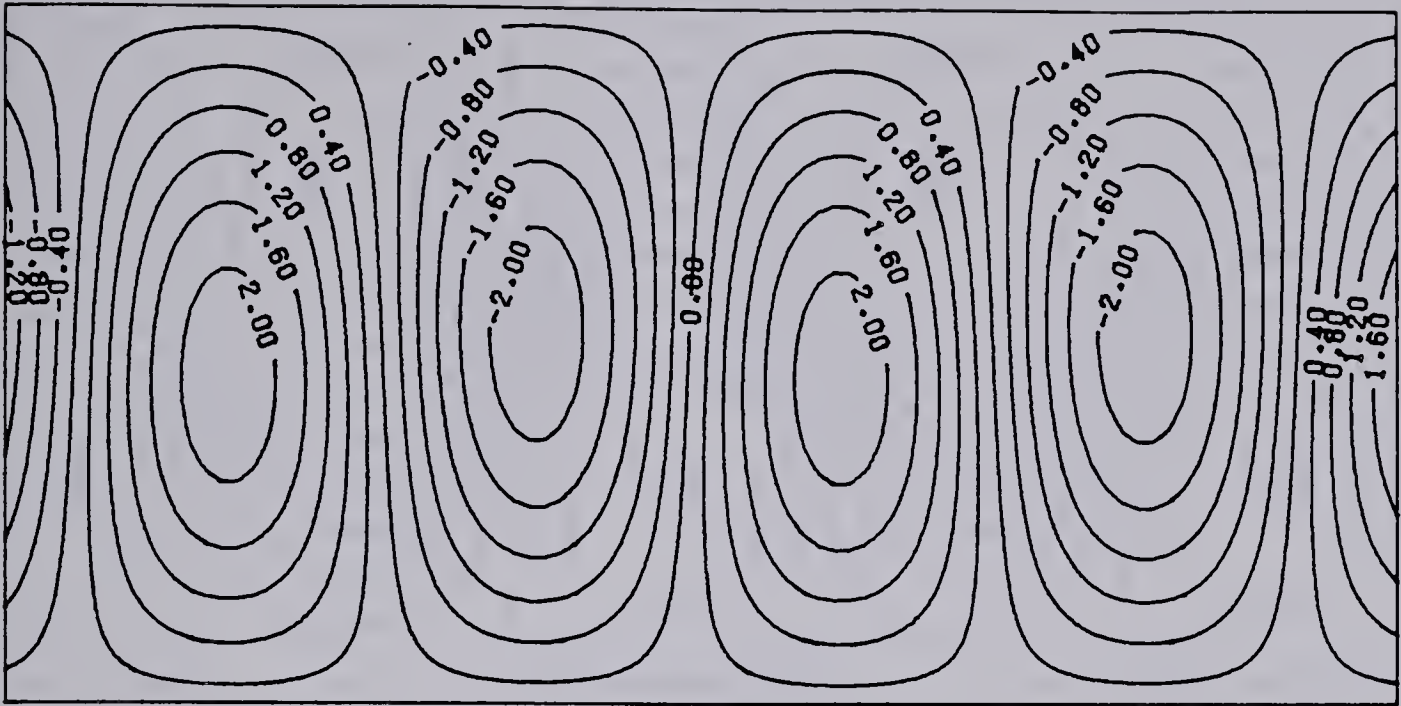
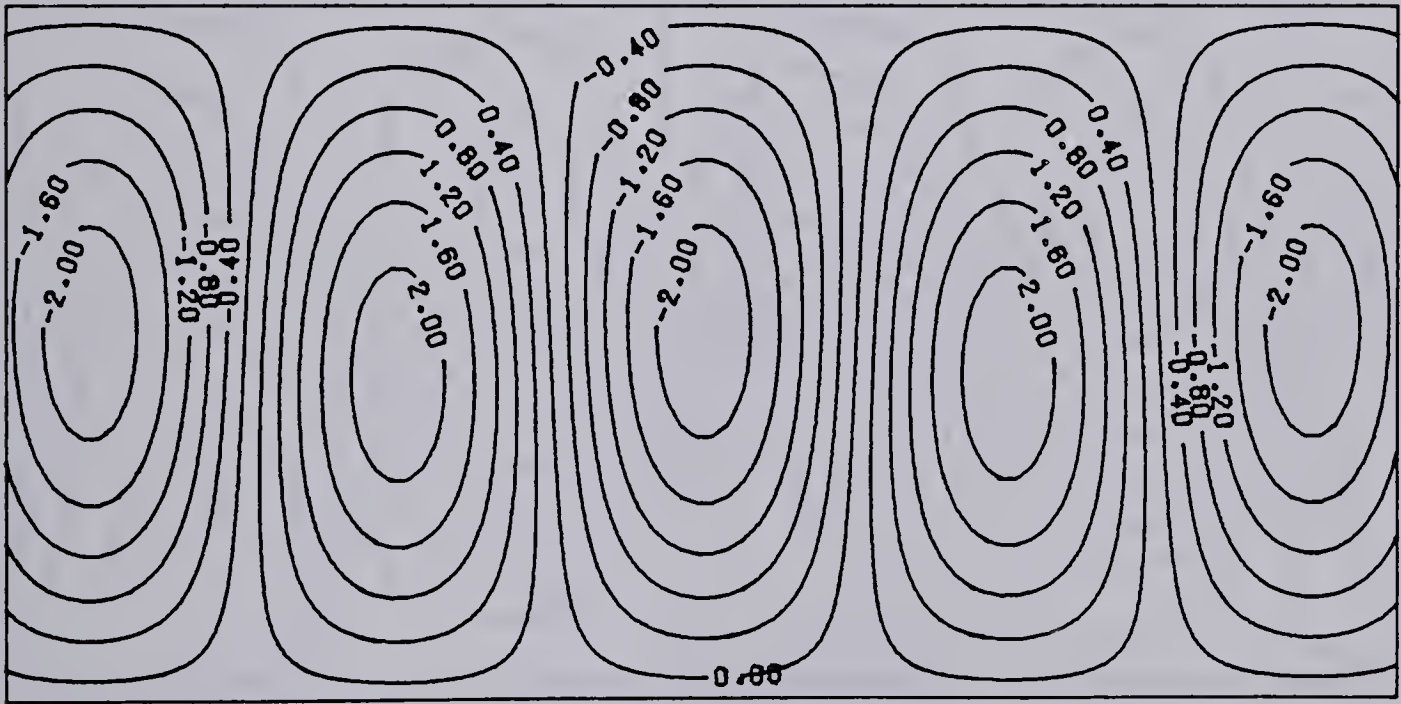


Fig. 5.30 The 24-hour spectral solution for the thickness field in metres for Case III in (a) area B and (b) area A.



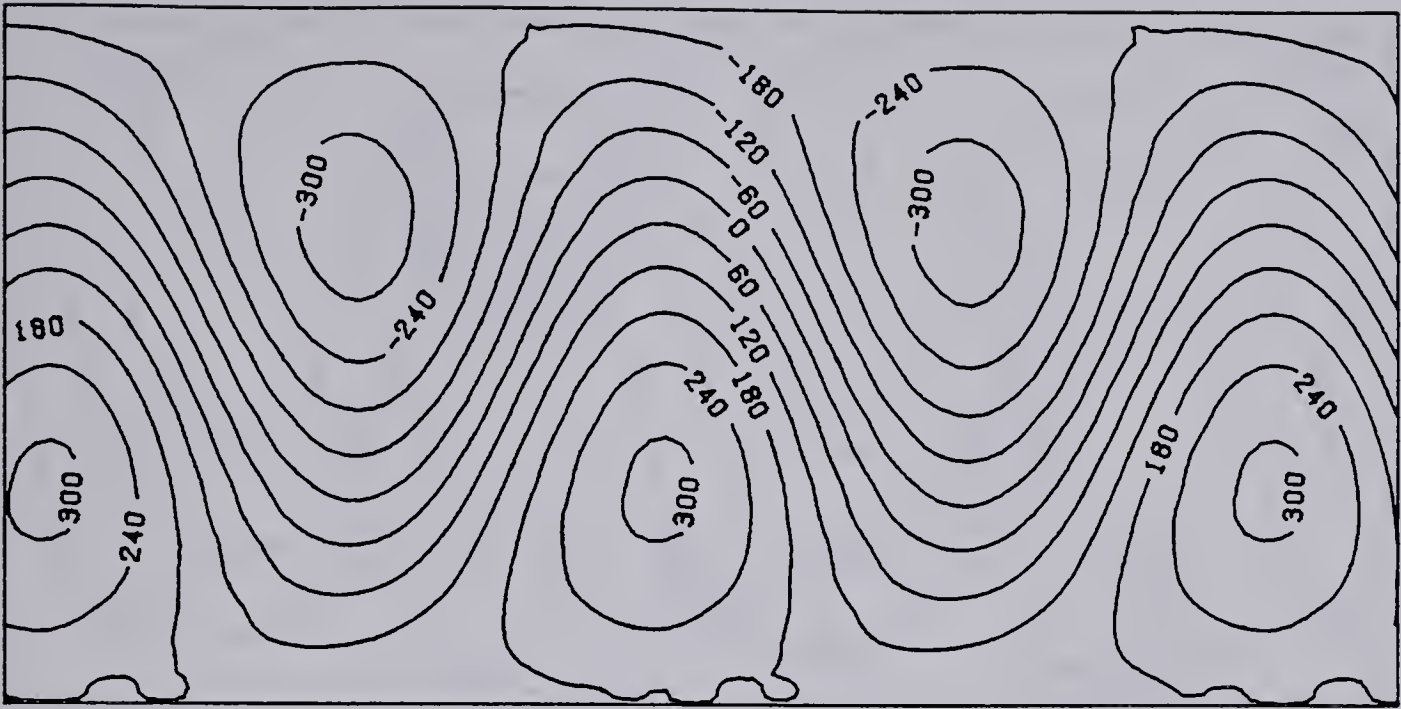
(a)



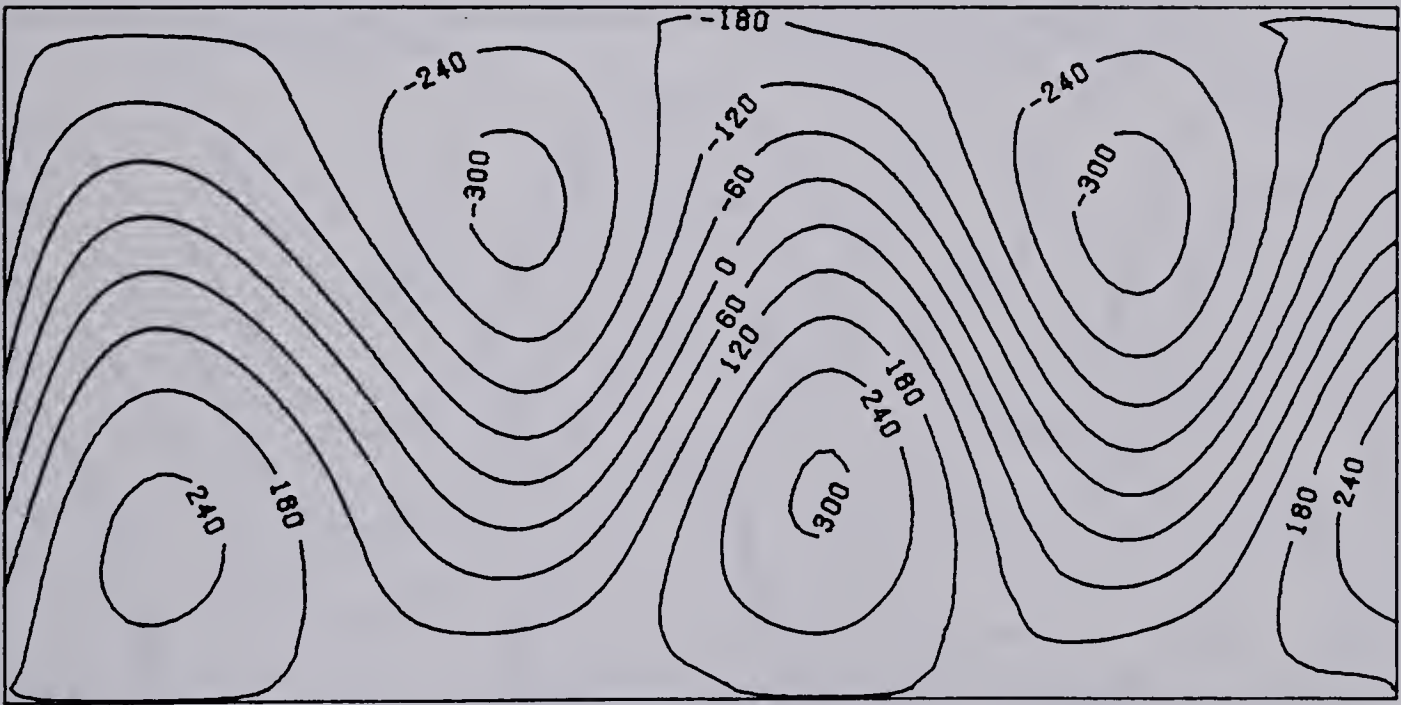
(b)

Fig. 5.31

The 24-hour spectral solution for the omega field in $\mu\text{bar}/\text{sec}$ for Case III in (a) area B and (b) area A.

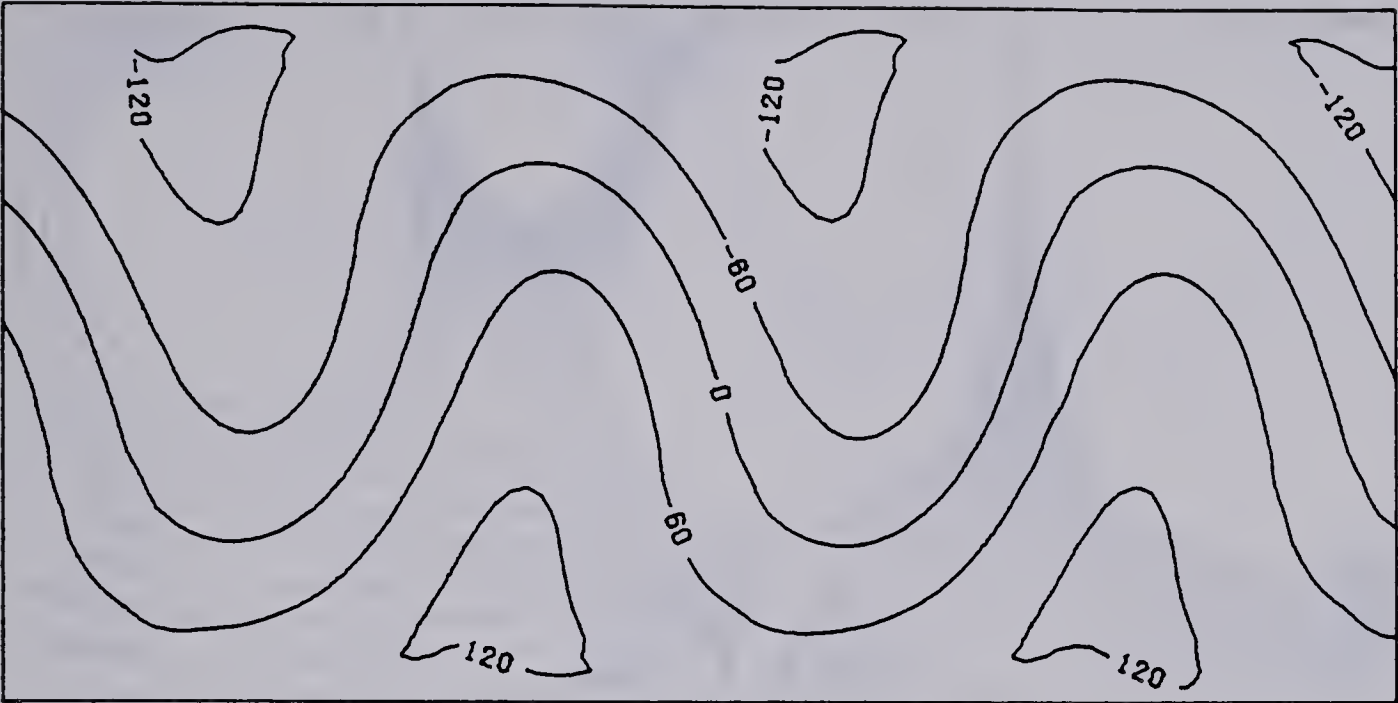


(a)

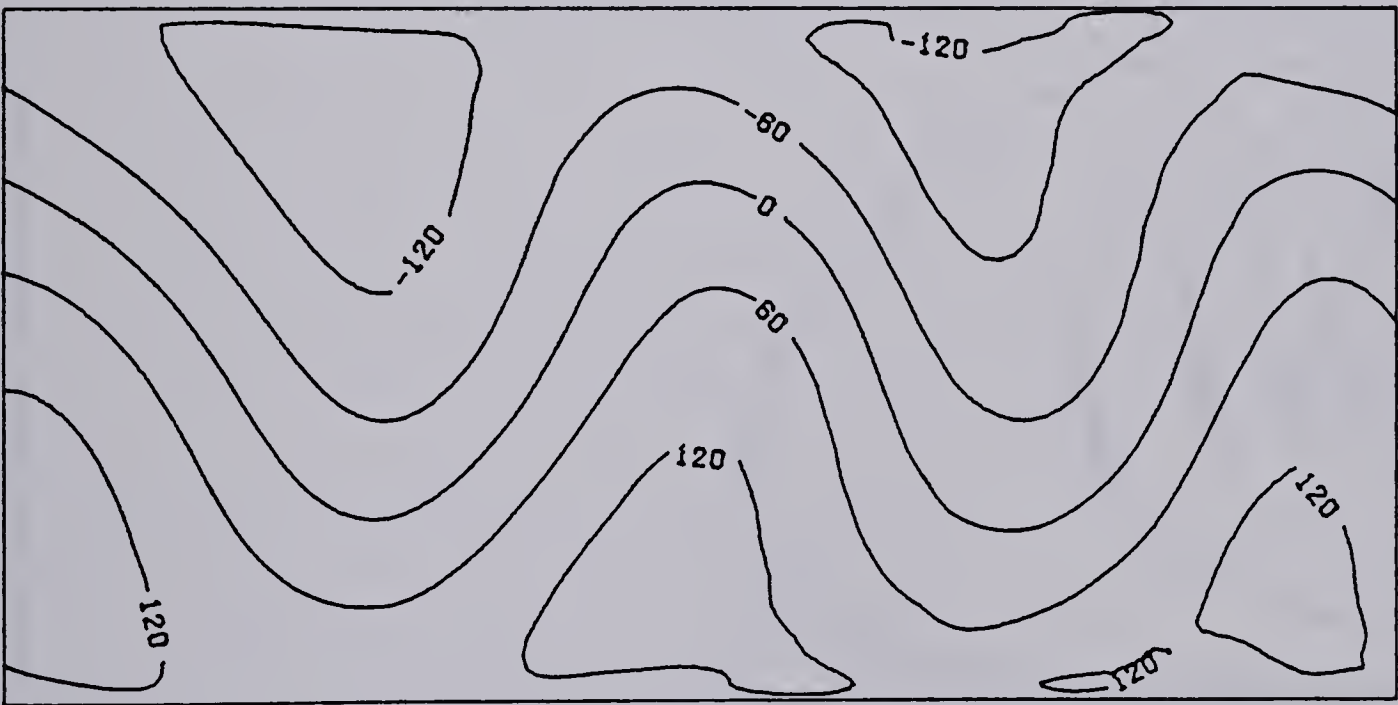


(b)

Fig. 5.32 The 24-hour finite-element solution for the mean height field in metres for Case II in (a) area B and (b) area A.



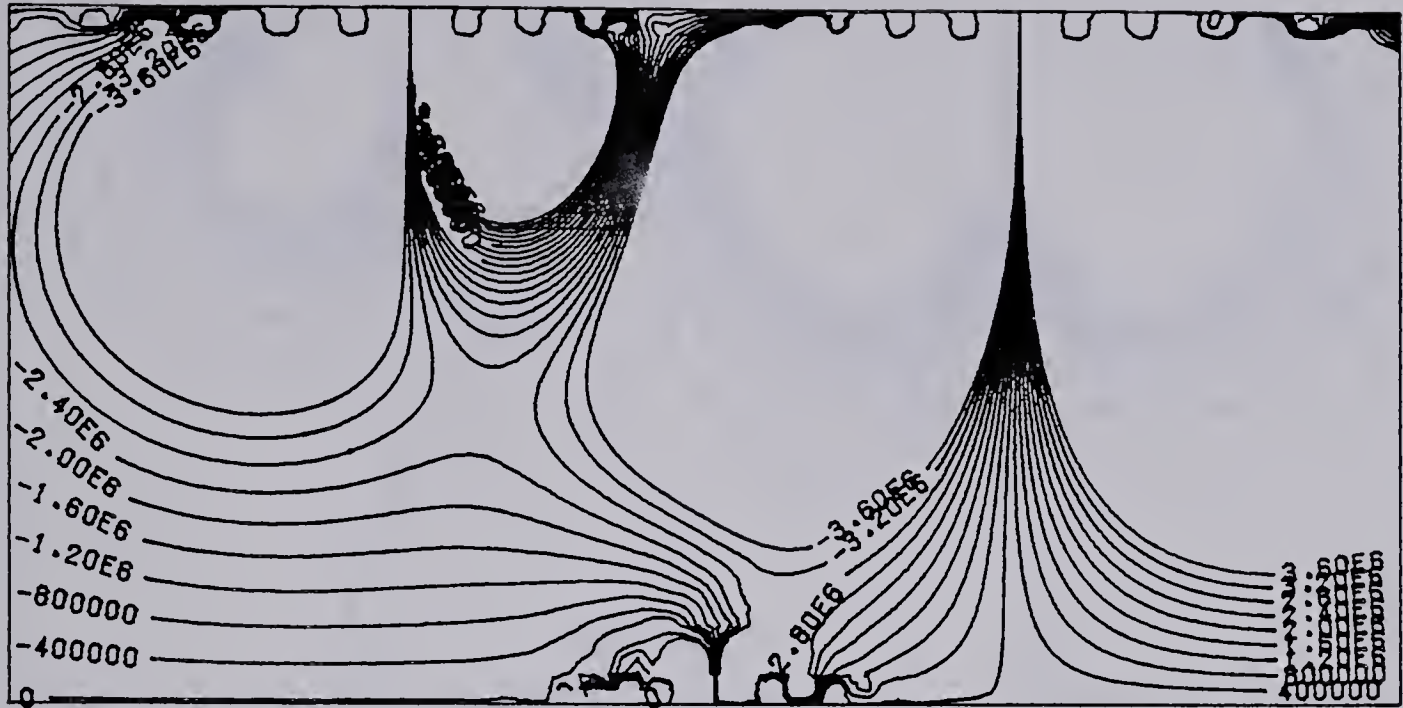
(a)



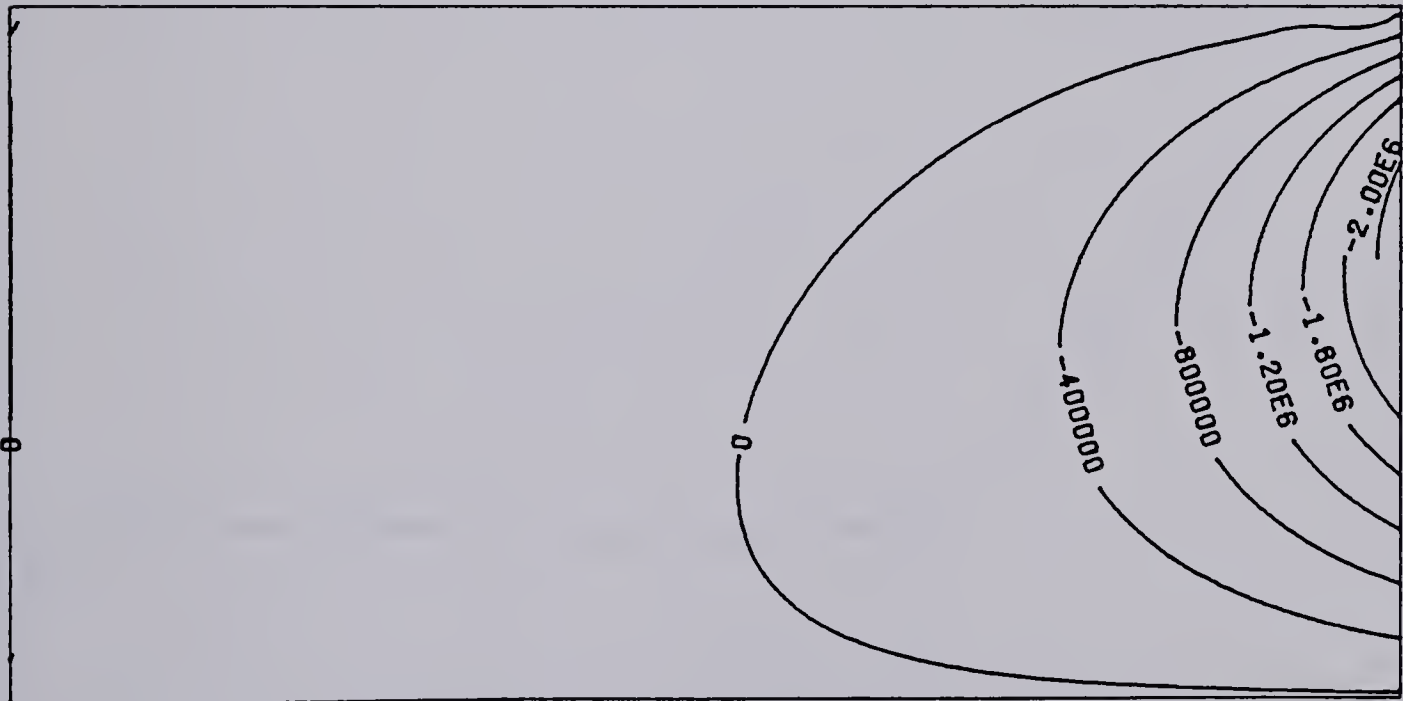
(b)

Fig. 5.33

The 24-hour finite-element solution for the thickness field in metres for Case III in (a) area B and (b) area A.

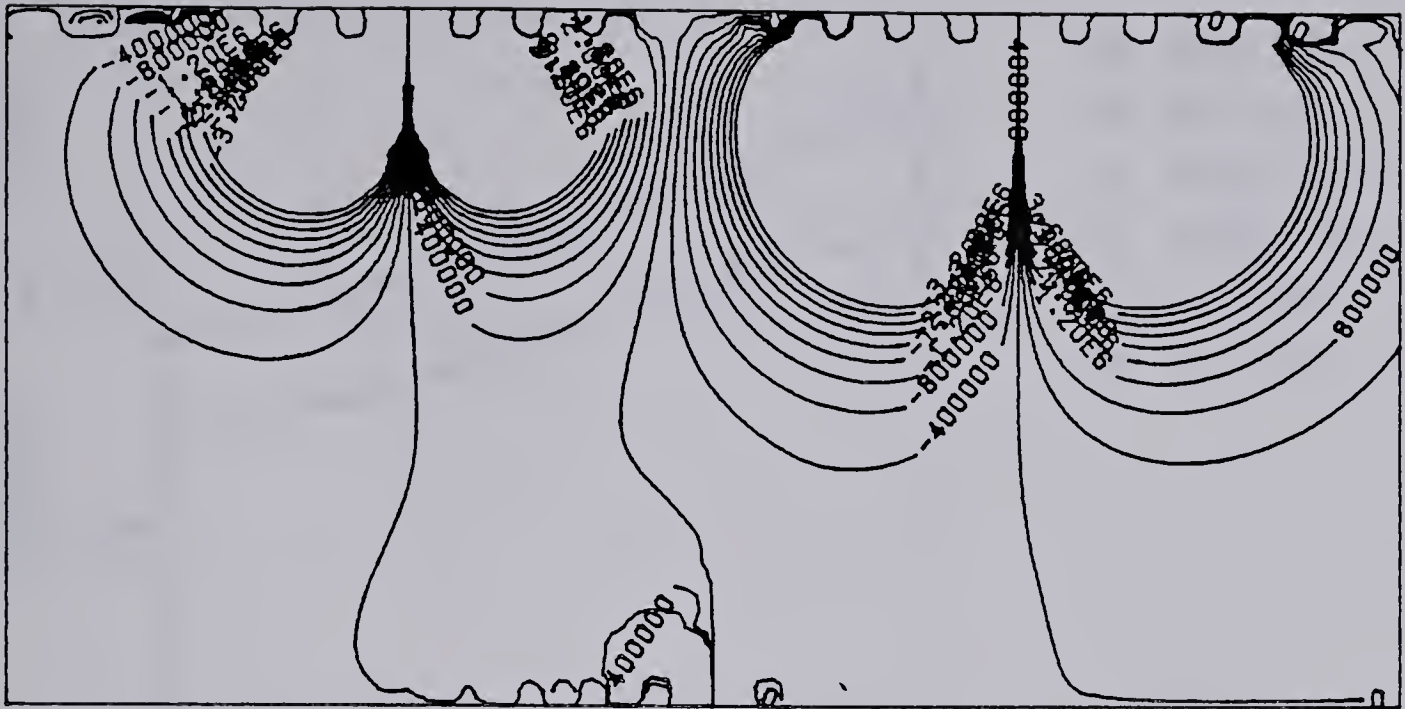


(a)



(b)

Fig. 5.34 The 30-hour finite-element solution for the mean height field in metres for Case III in (a) area B and (b) area A.



(a)

(b)

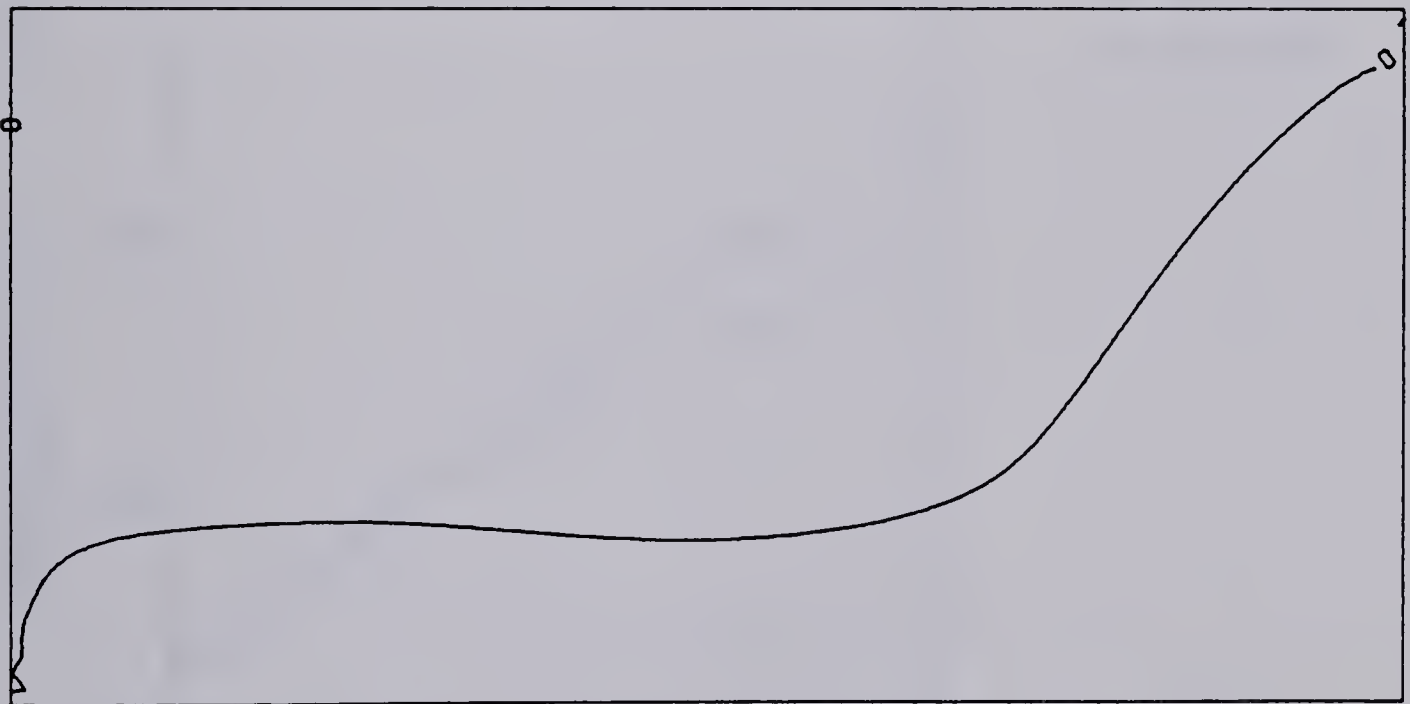
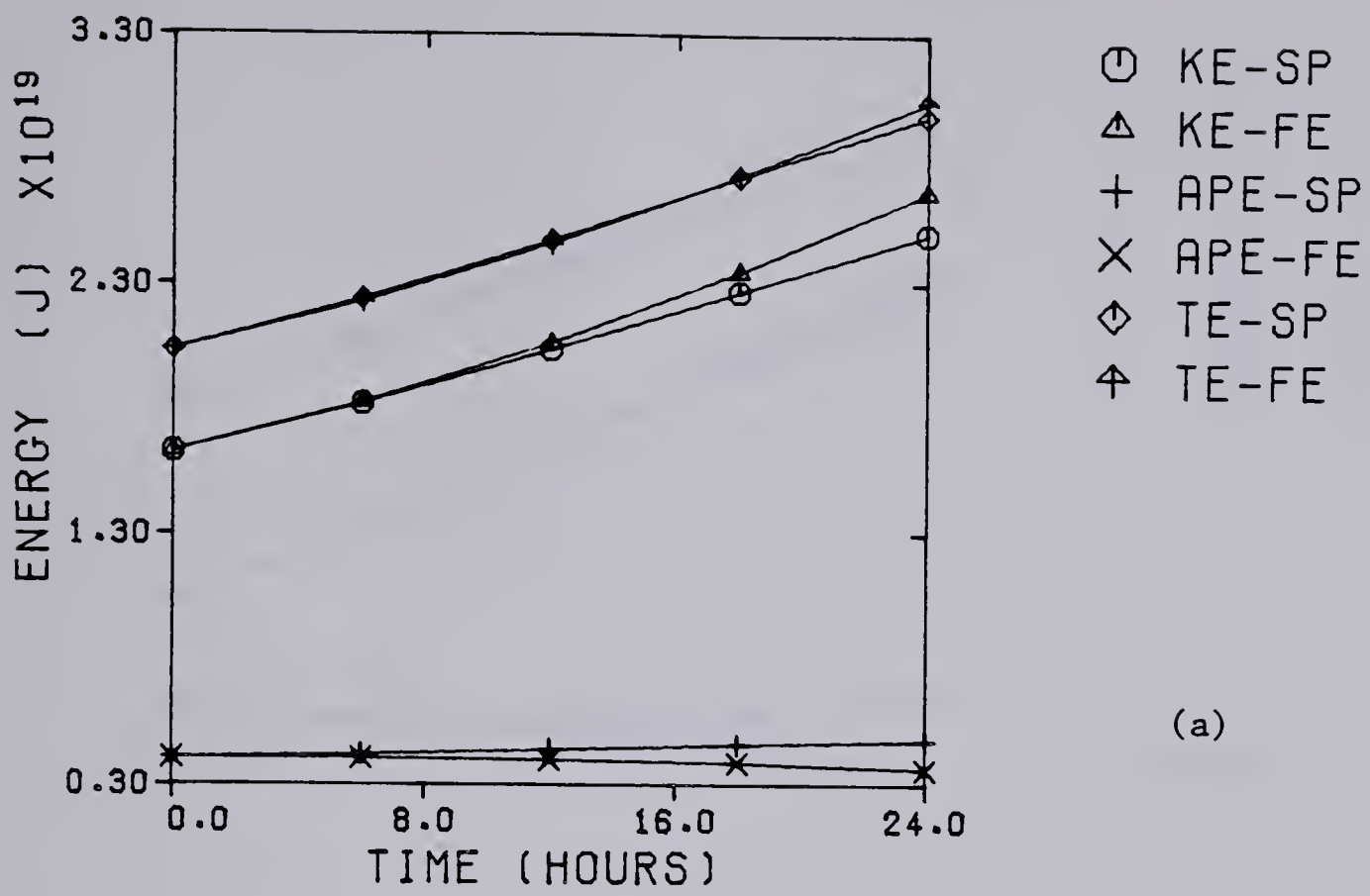
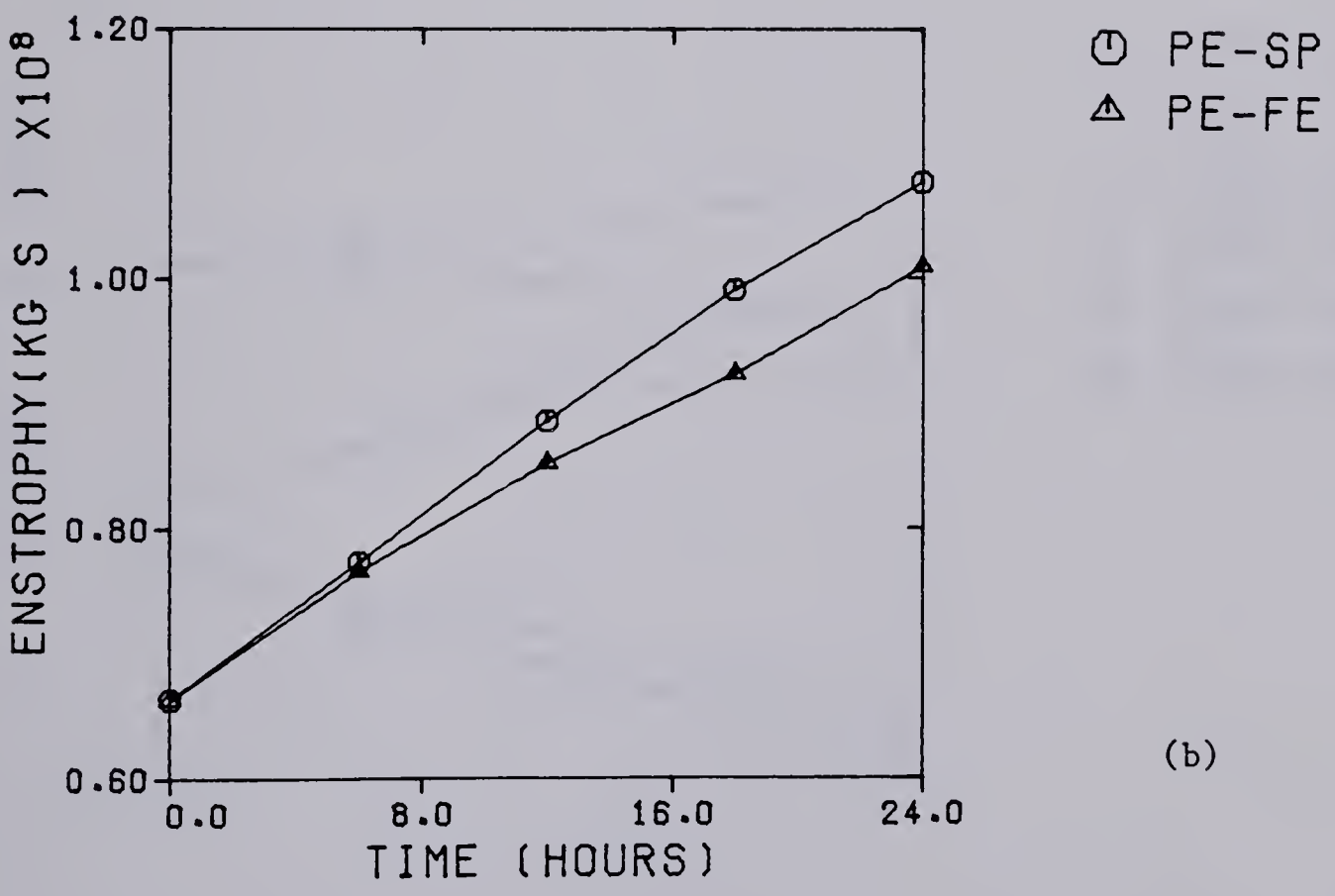


Fig. 5.35 The 30-hour finite-element solution for the thickness field in metres for Case III in (a) area B and (b) area A.

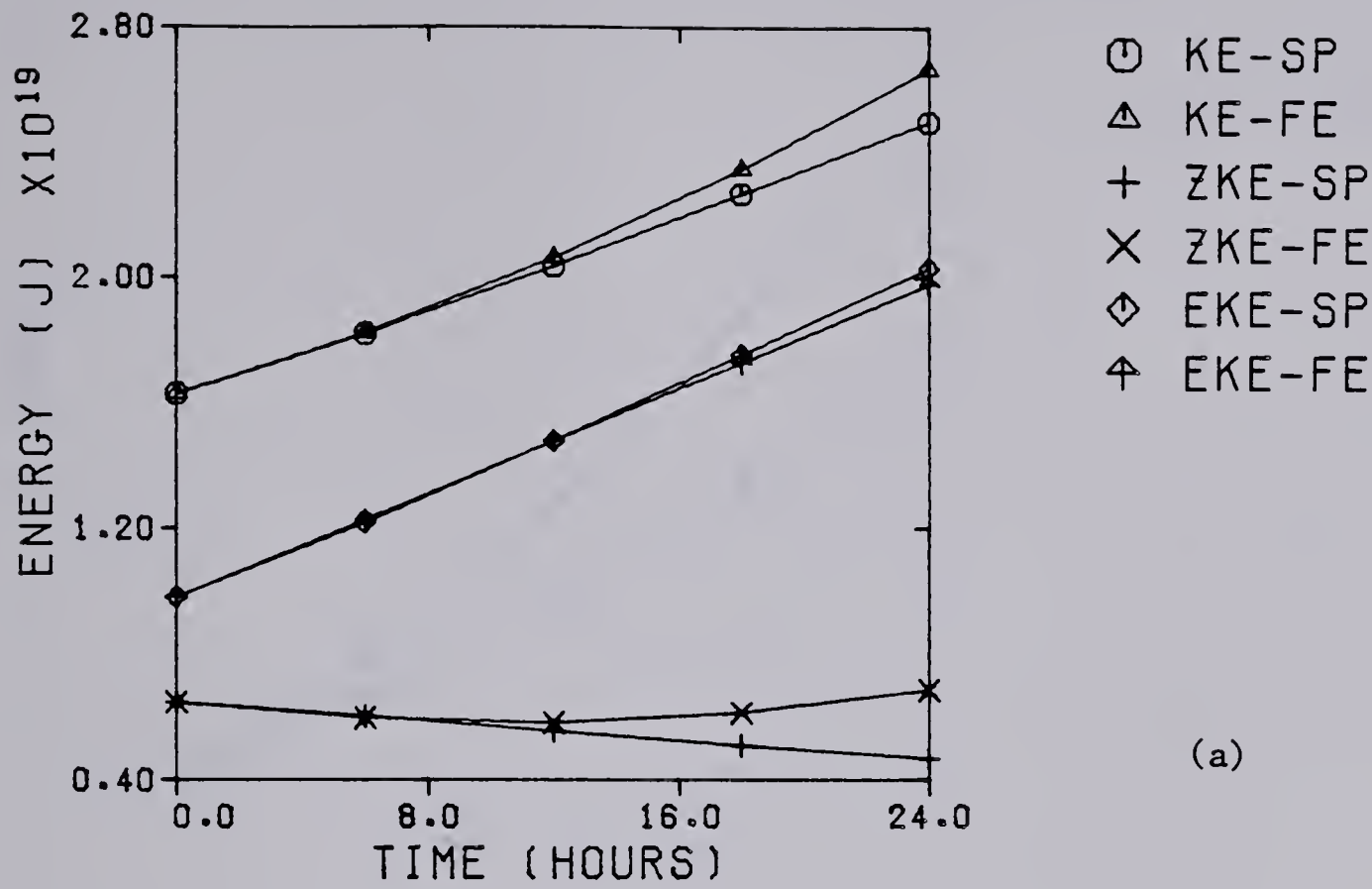


(a)

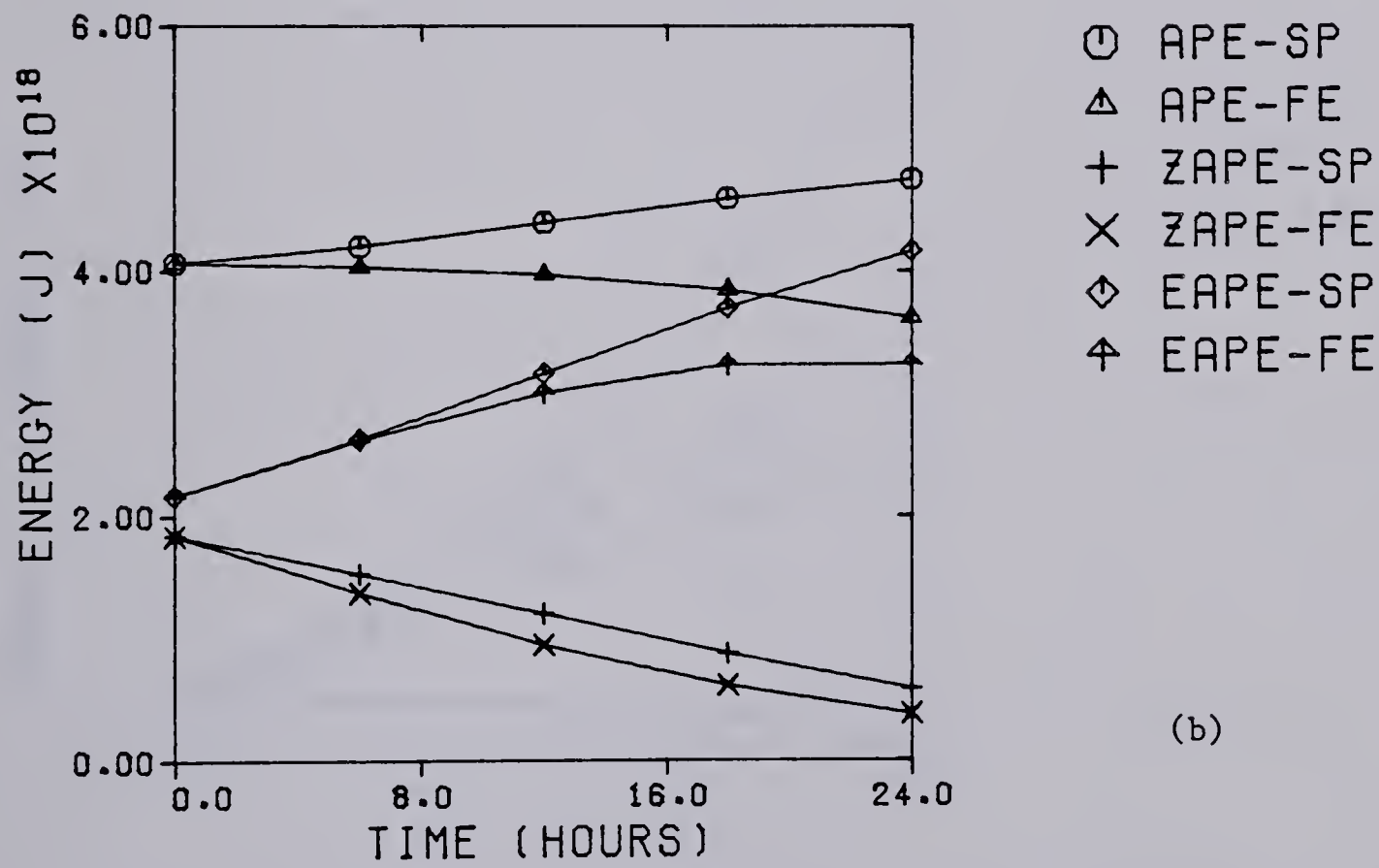


(b)

Fig. 5.36 The (a) energies and (b) potential enstrophy of Case III for the two solutions in the verification area.

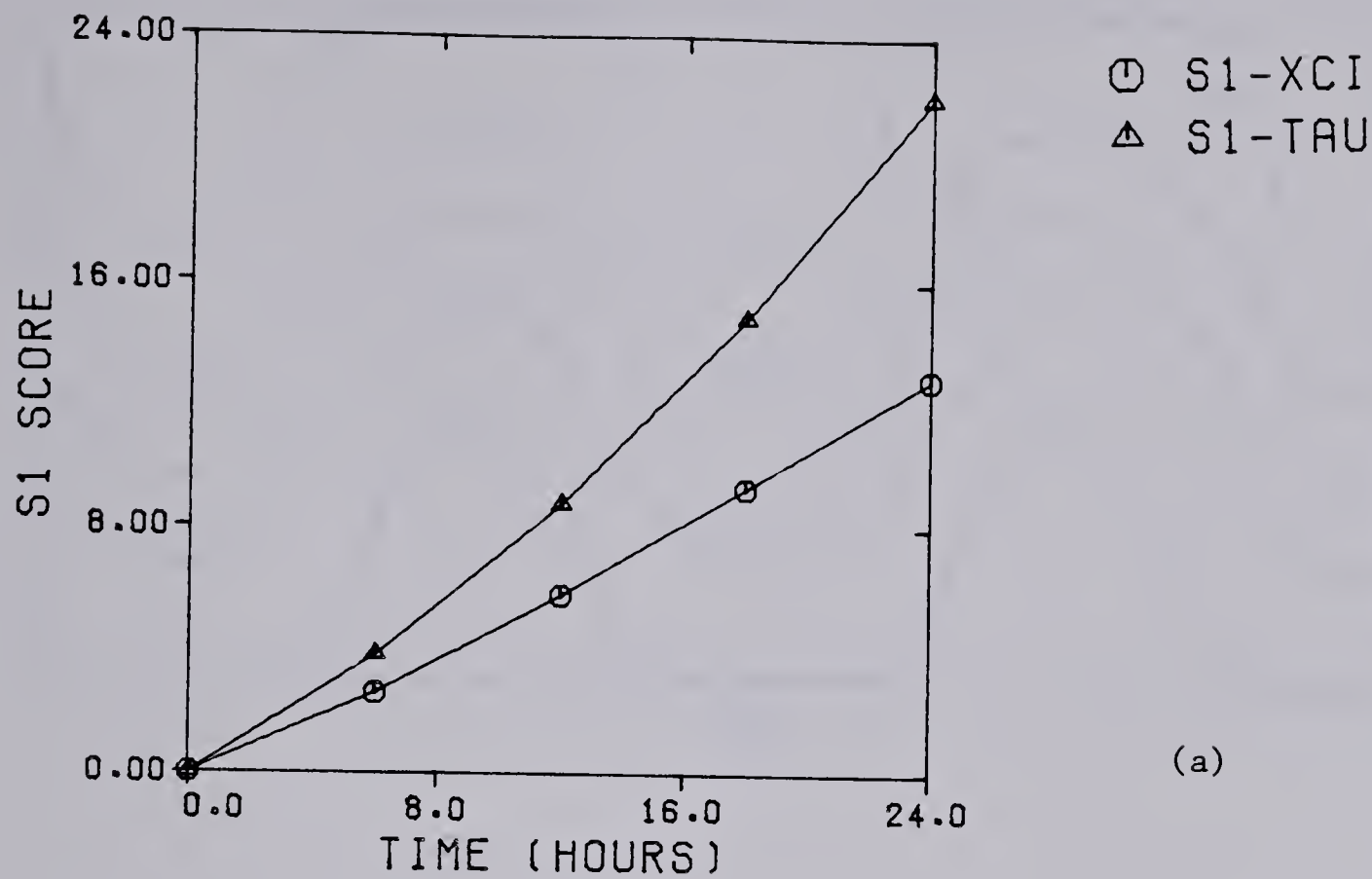


(a)

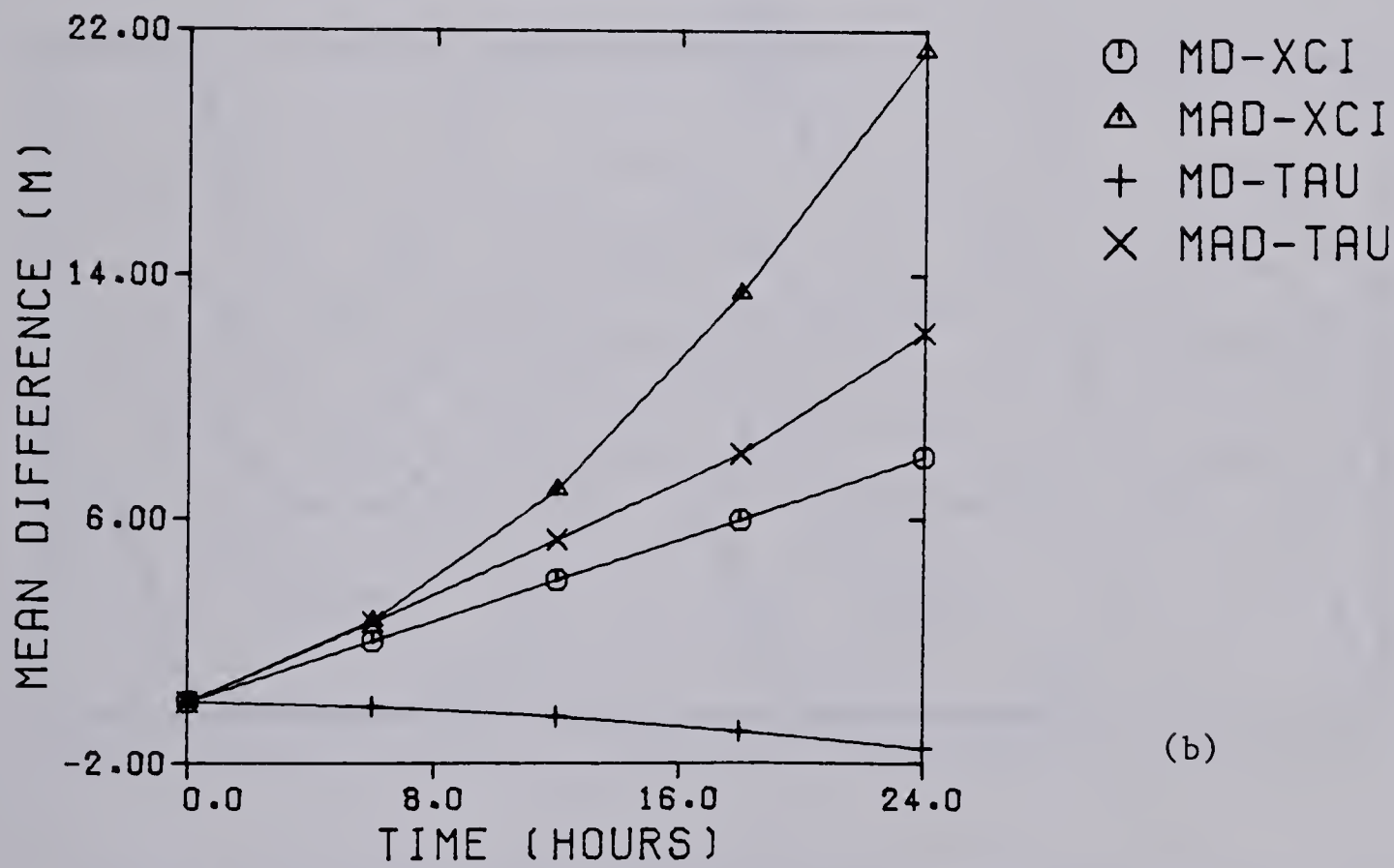


(b)

Fig. 5.37 The (a) kinetic energies and (b) potential energies of Case III for the two solutions in the verification area.

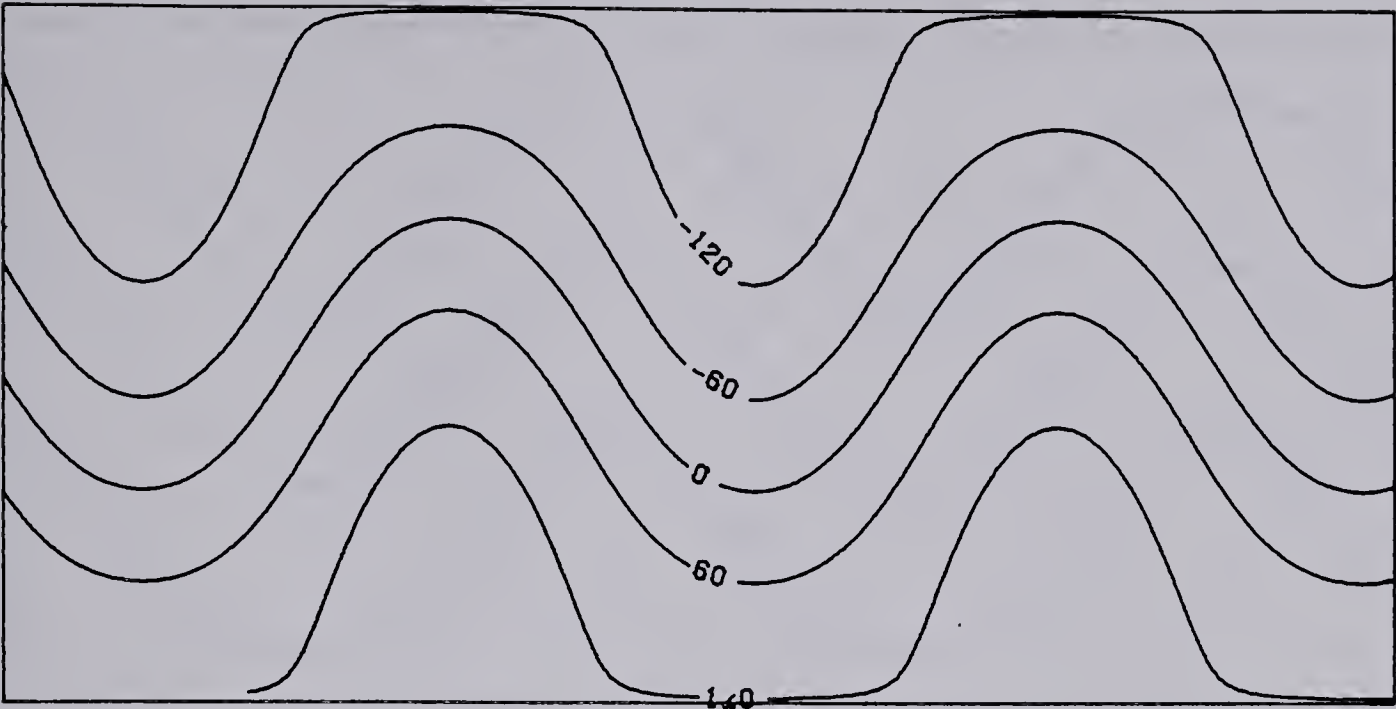


(a)

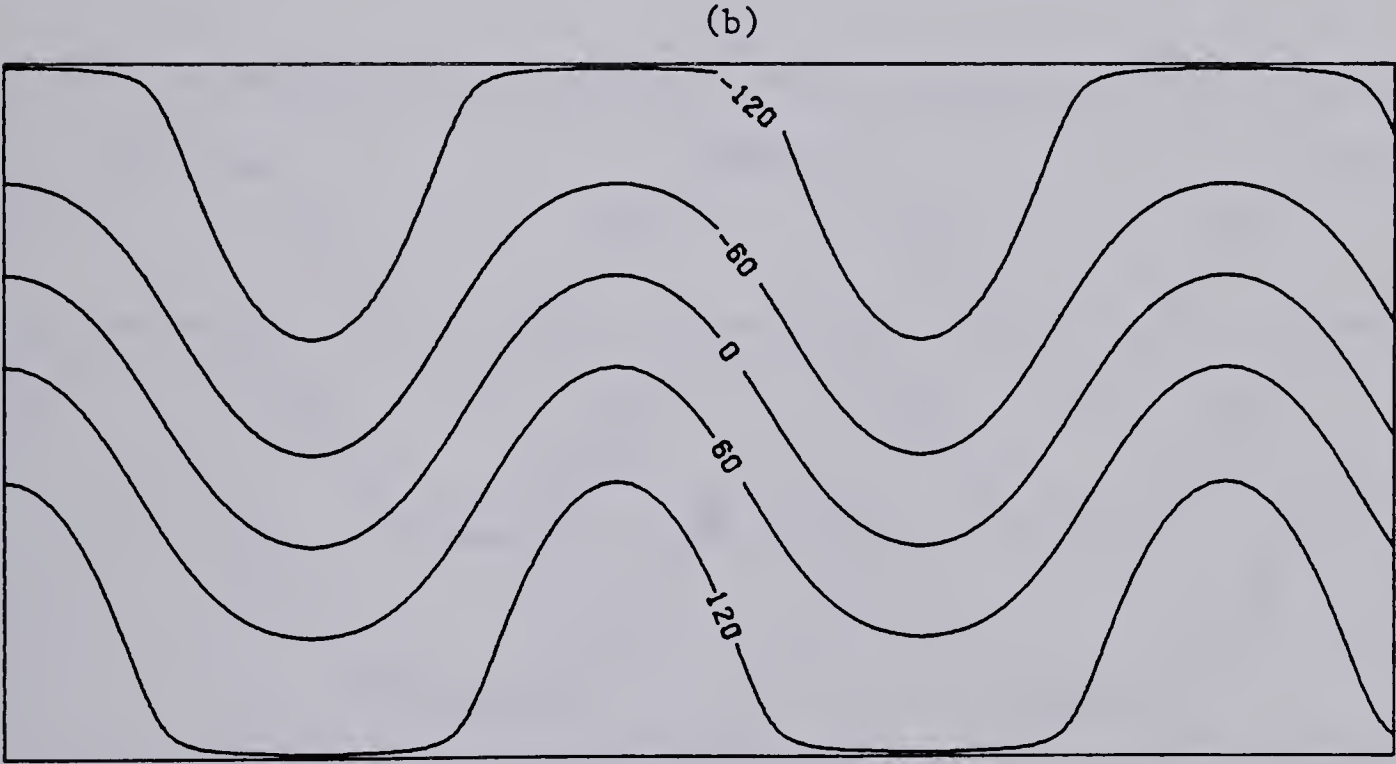


(b)

Fig. 5.38 The (a) S1 scores and (b) MD and MAD curves for Case III for the two solutions in the verification area.

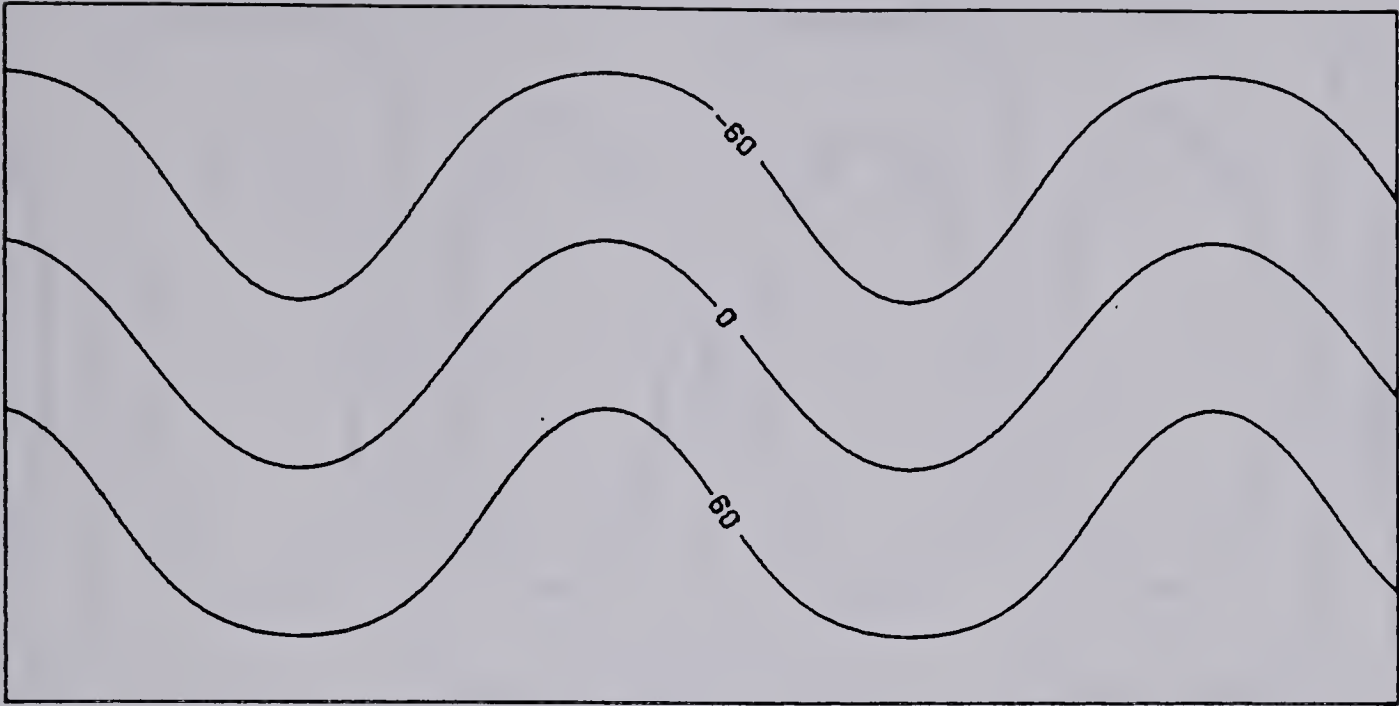


(a)

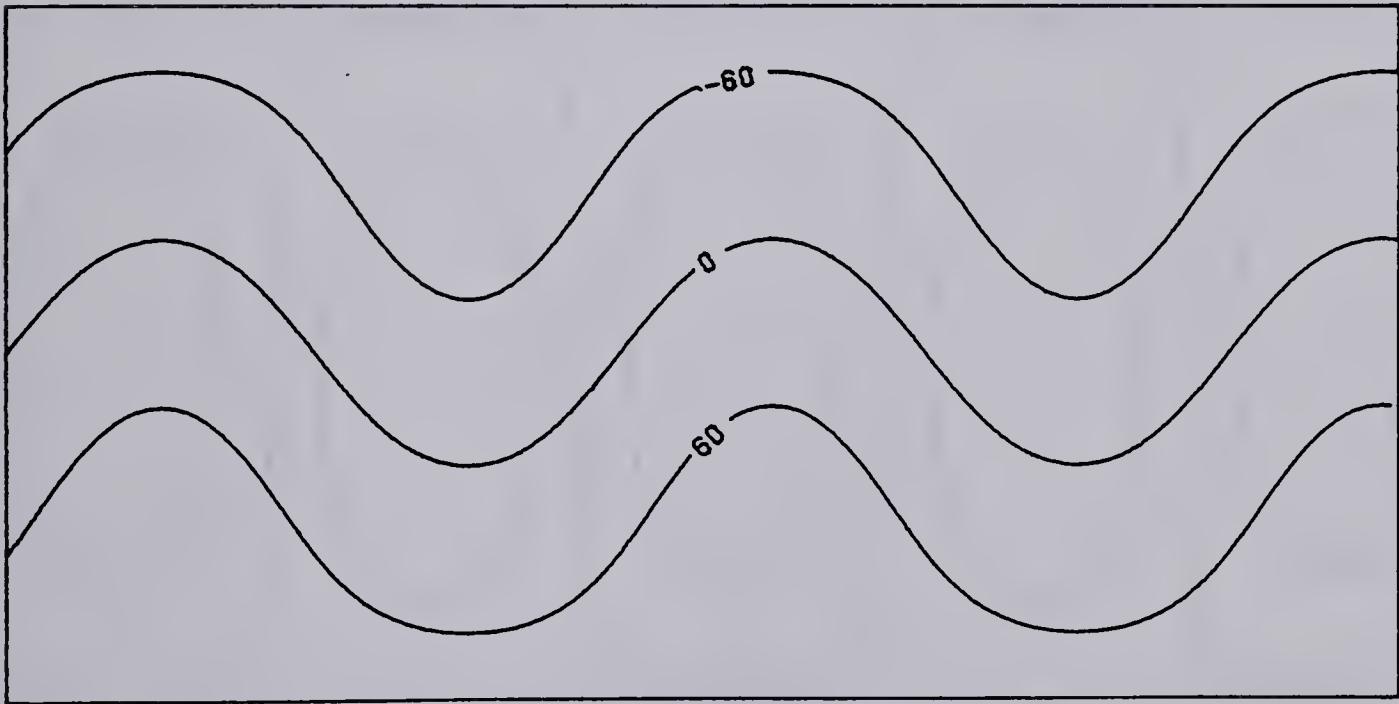


(b)

Fig. 5.39 Initial mean height field in metres for Case IV in (a) area B and (b) area A.

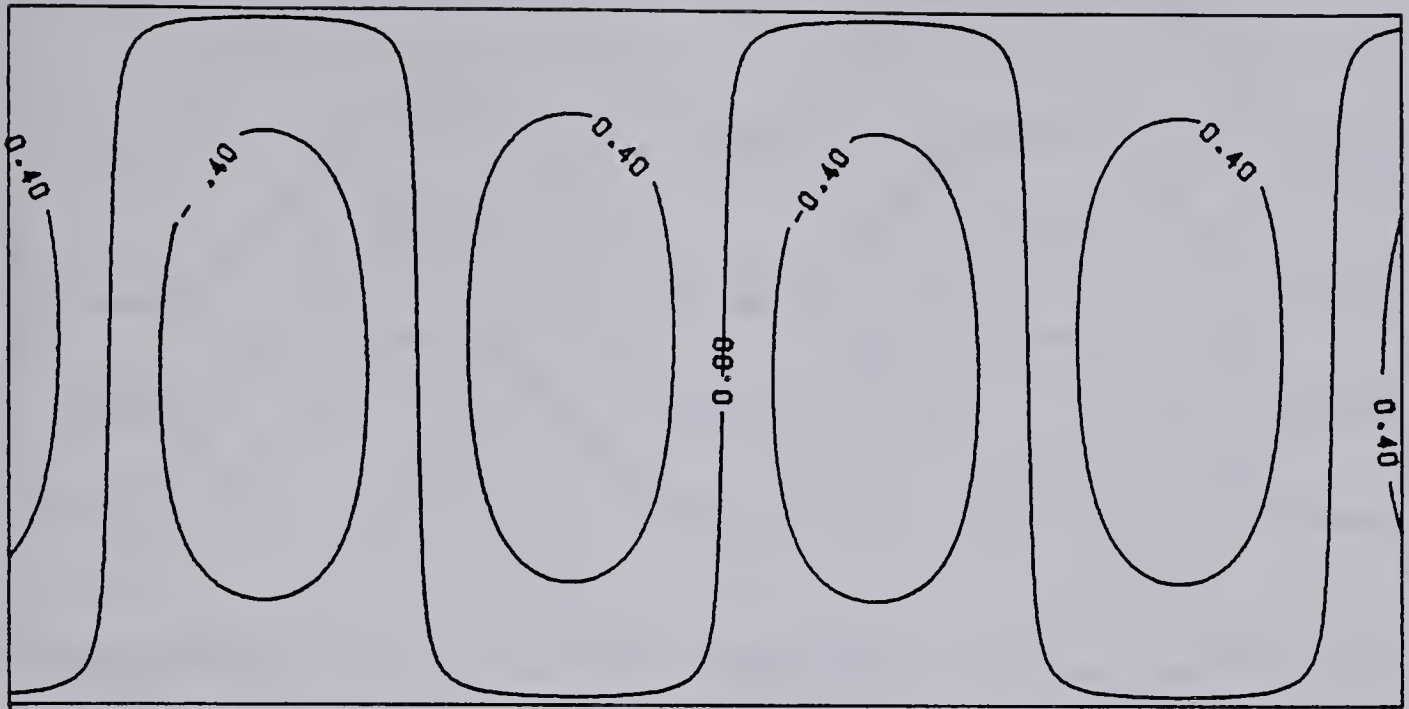


(a)

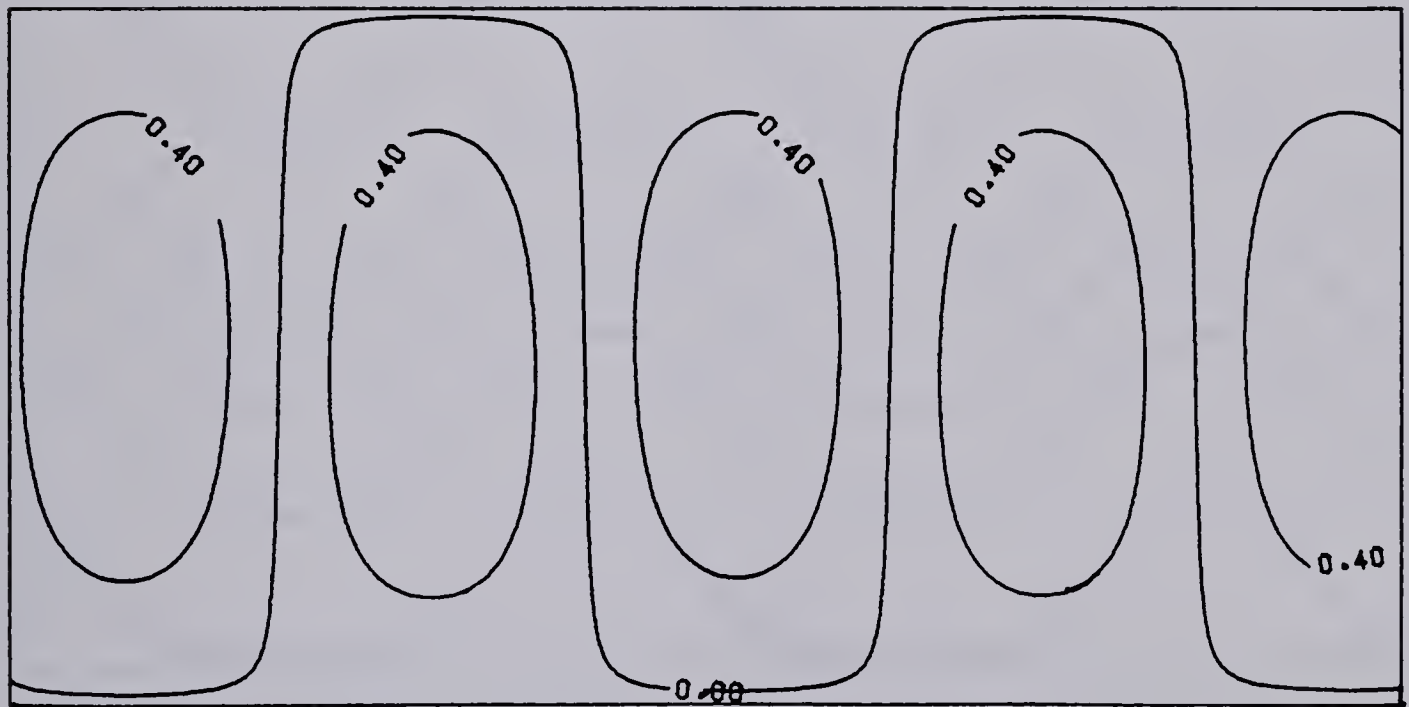


(b)

Fig. 5.40 Initial thickness field in metres for Case IV
in (a) area B and (b) area A.

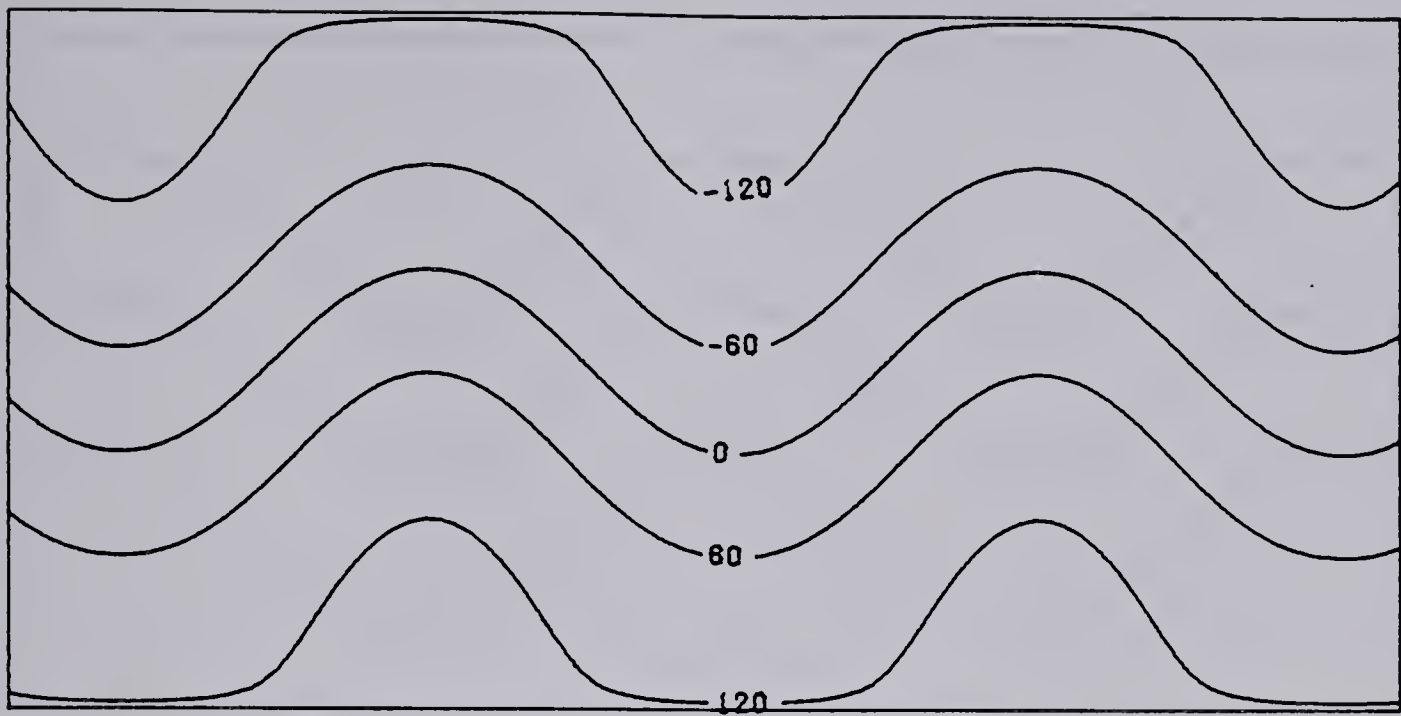


(a)

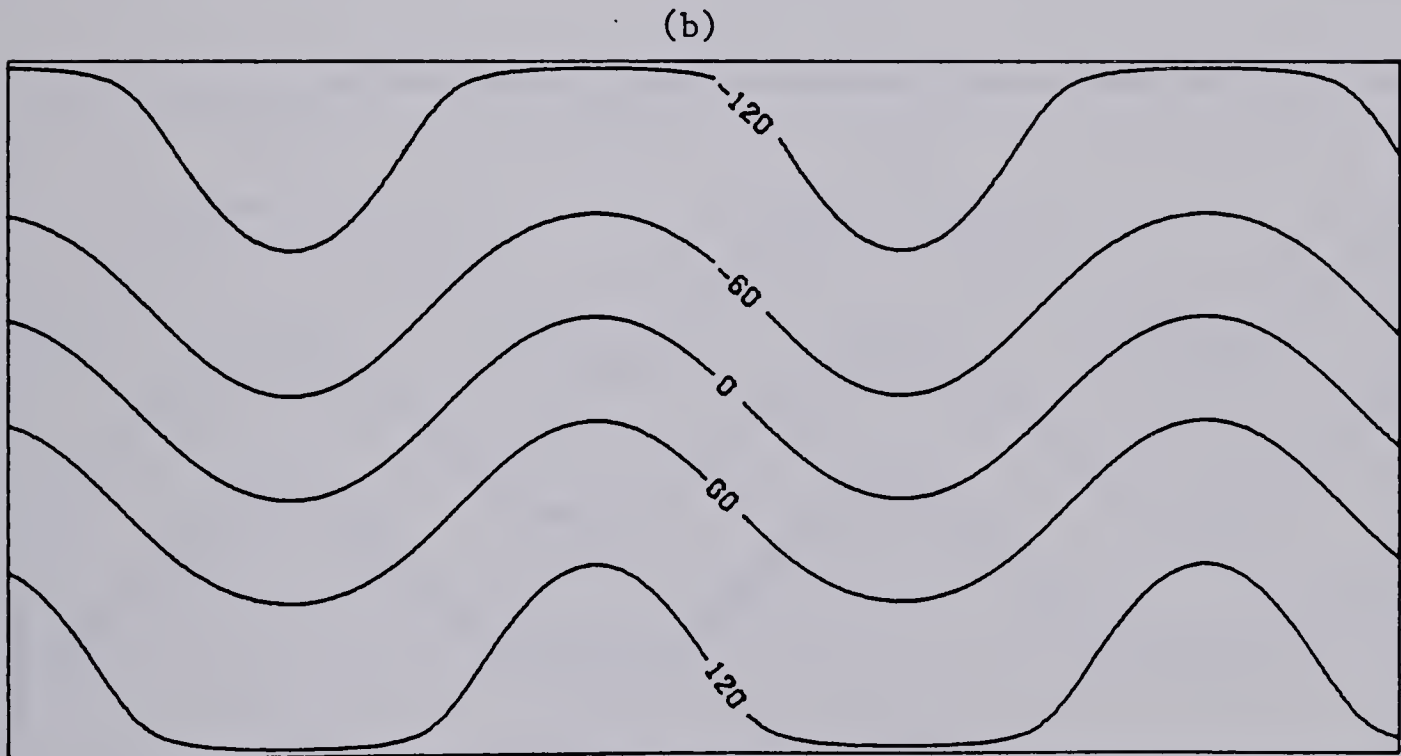


(b)

Fig. 5.41 Initial omega field in $\mu\text{bar}/\text{sec}$ for Case IV in
(a) area B and (b) area A.

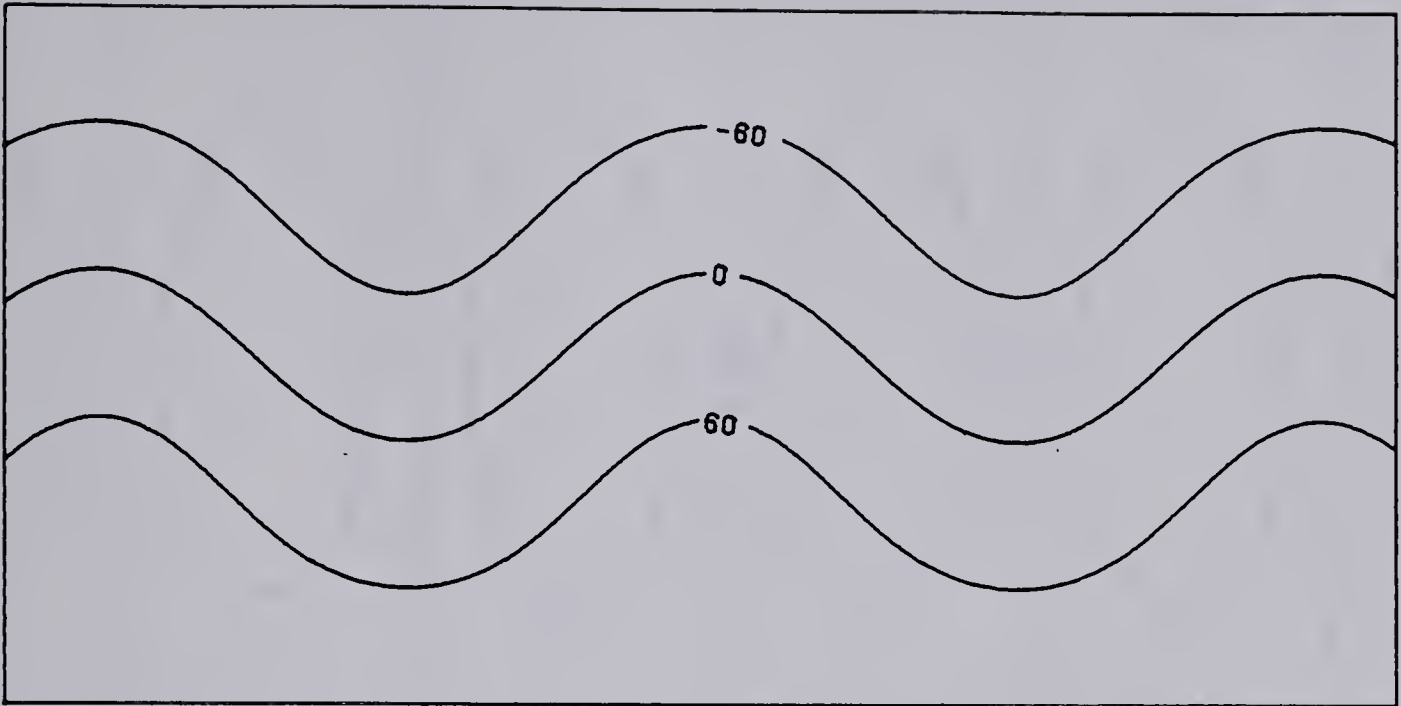


(a)

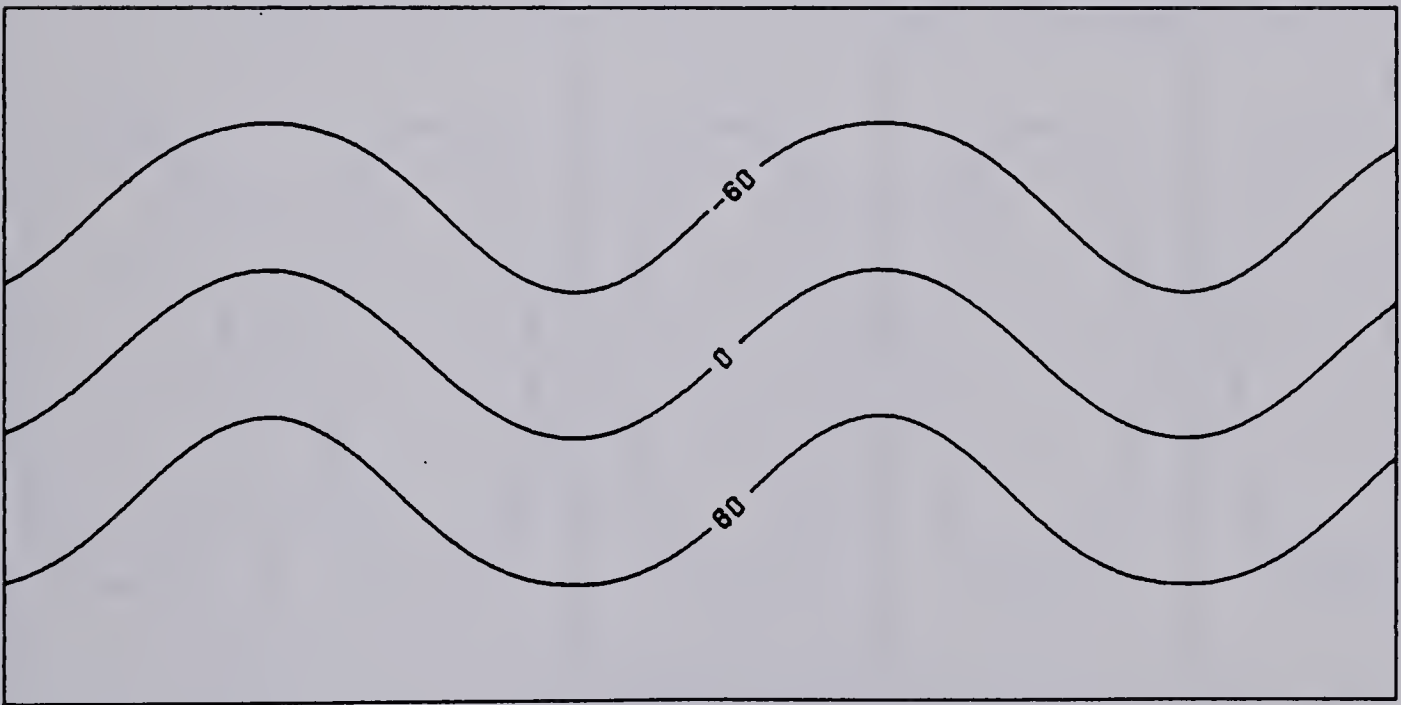


(b)

Fig. 5.42 The 48-hour spectral solution for the mean height field in metres for Case IV in (a) area B and (b) area A.

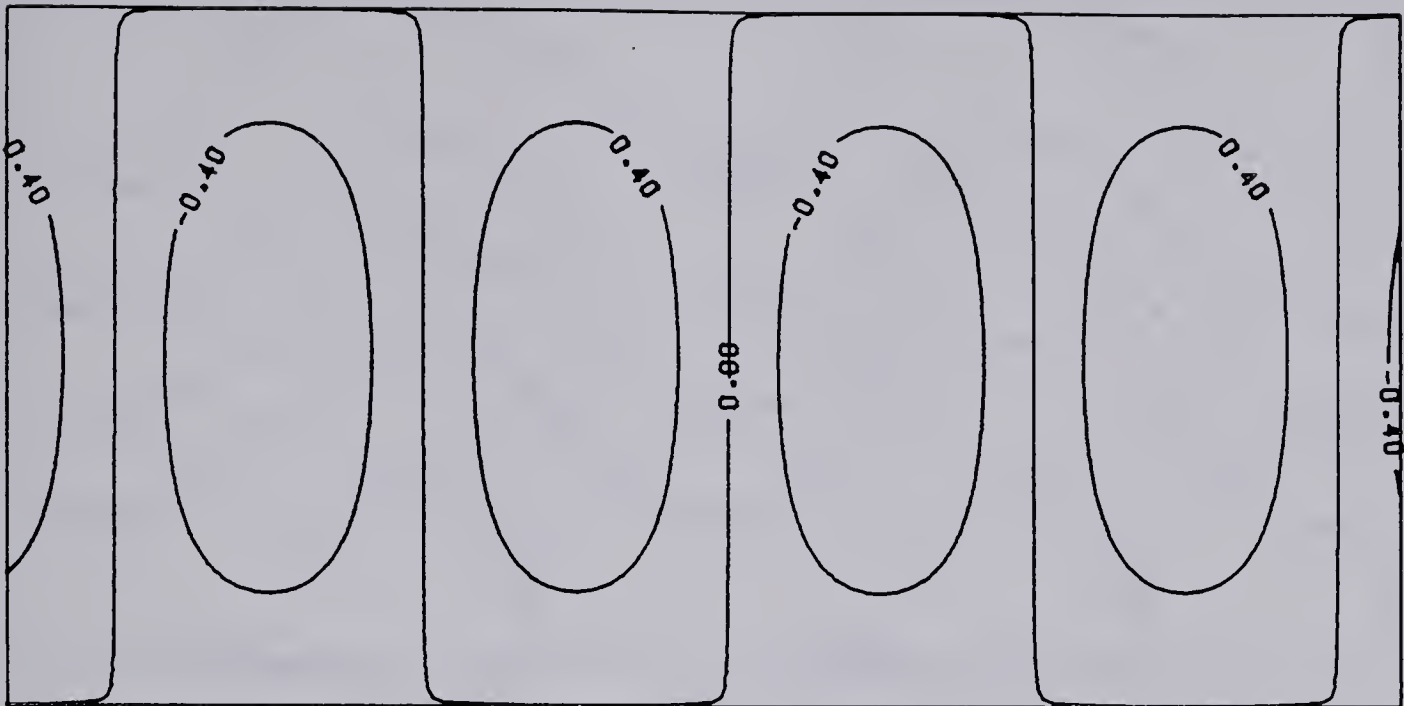


(a)



(b)

Fig. 5.43 The 48-hour spectral solution for the thickness field in metres for Case IV in (a) area B and (b) area A.



(a)

(b)

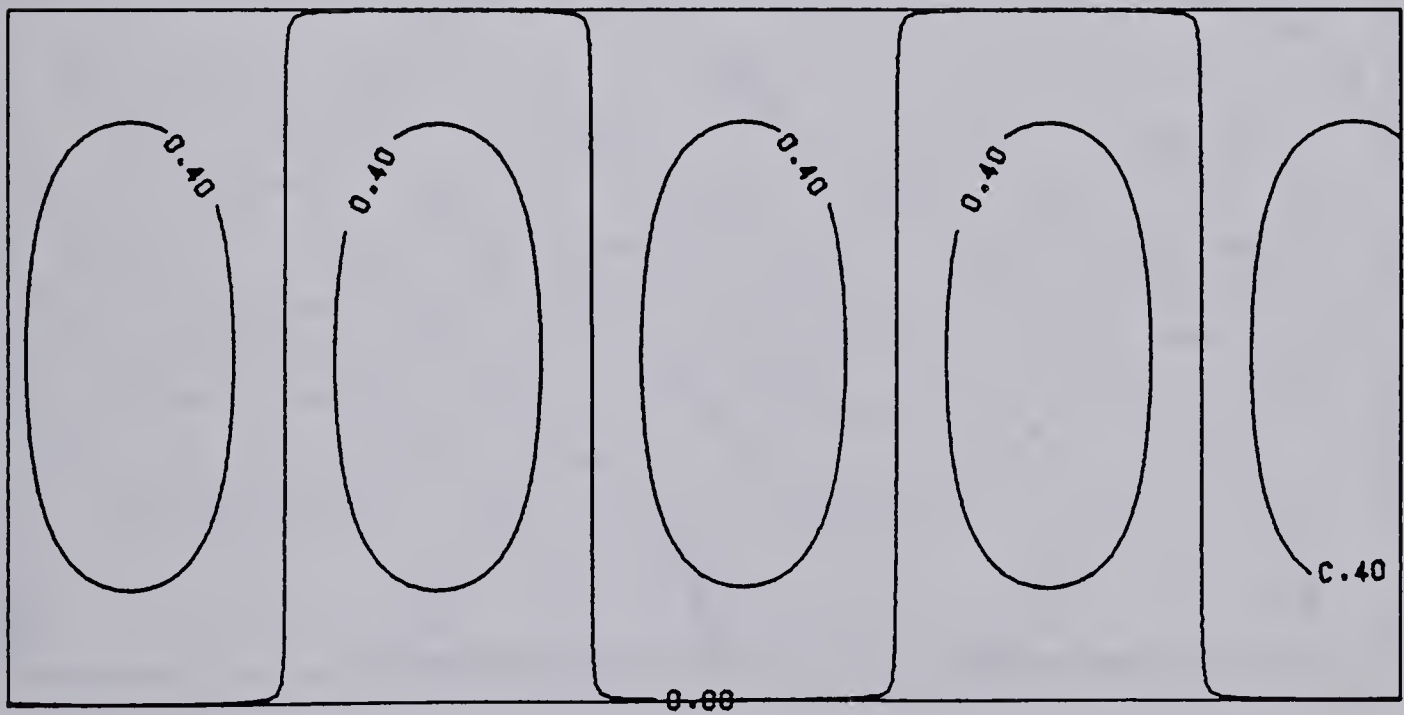
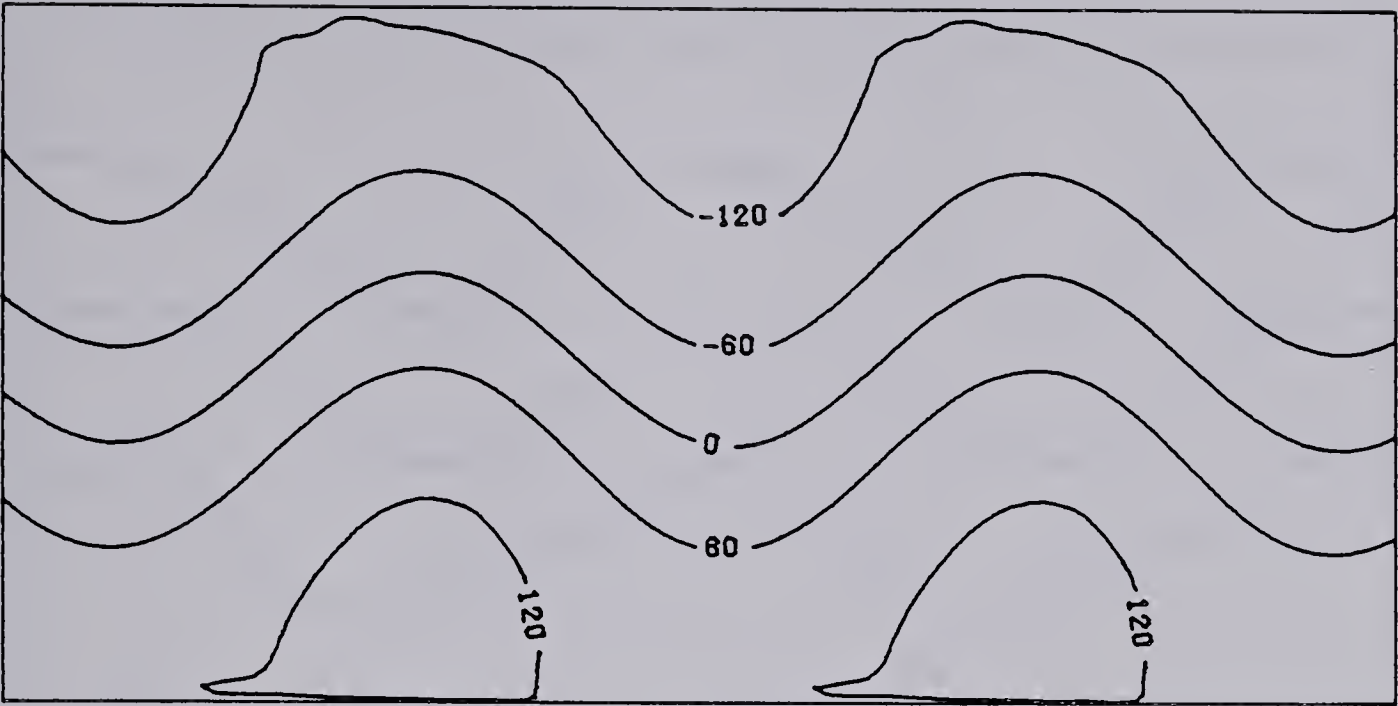


Fig. 5.44 The 48-hour spectral solution for the omega field in $\mu\text{bar/sec}$ for Case IV in (a) area B and (b) area A.



(a)

(b)

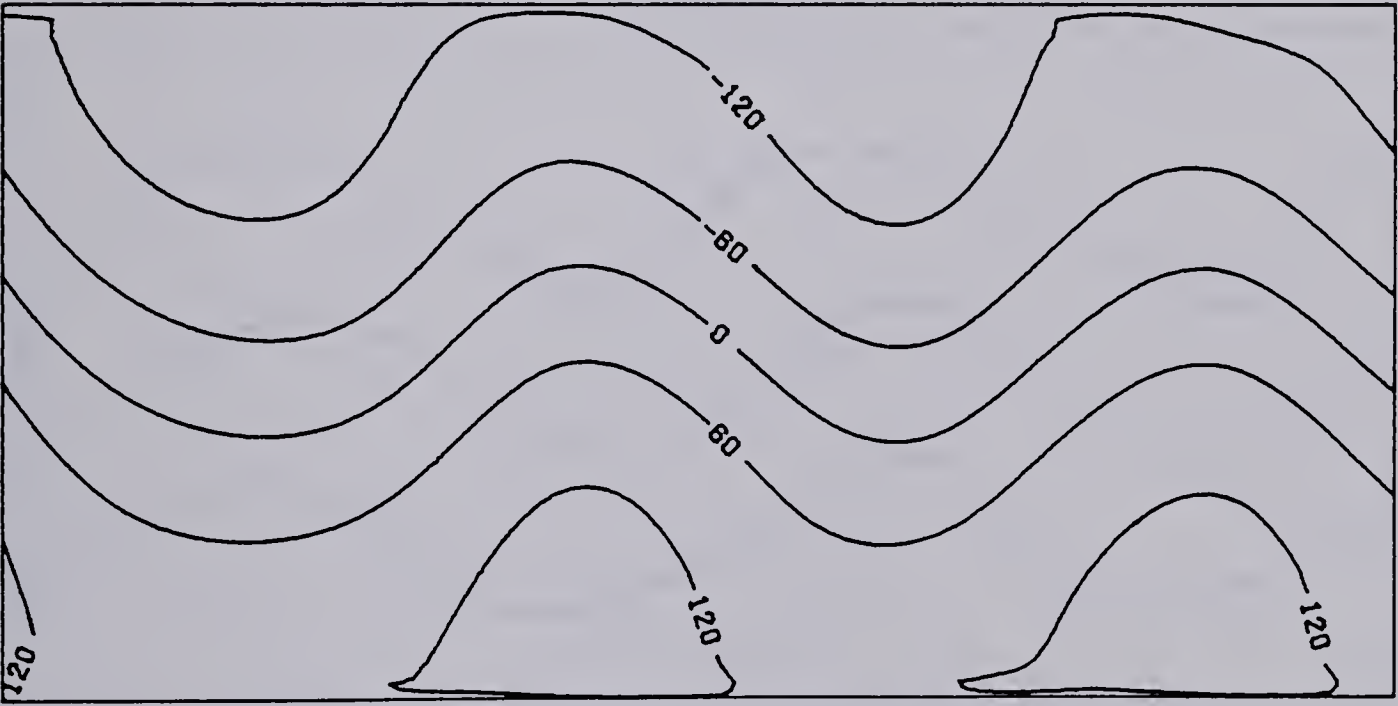
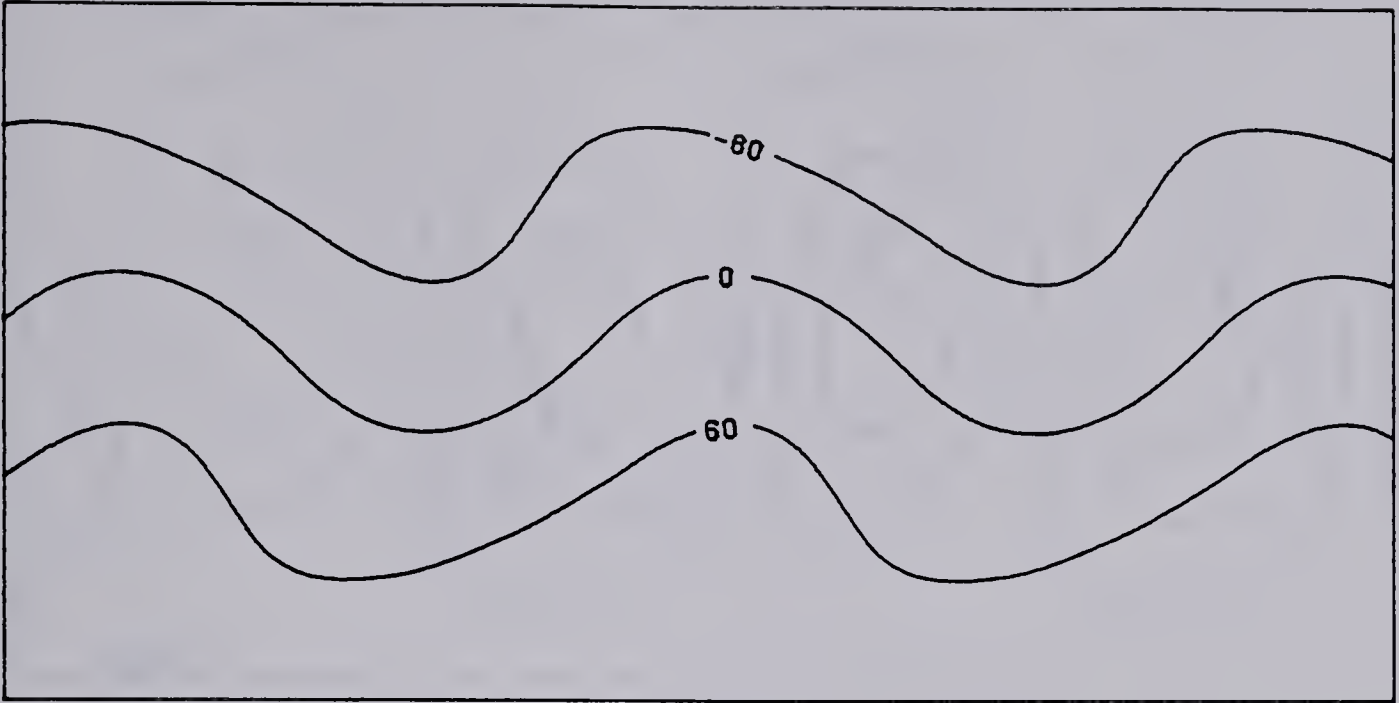


Fig. 5.45 The 48-hour finite-element solution for the mean height field in metres for Case IV in (a) area B and (b) area A.



(a)

(b)

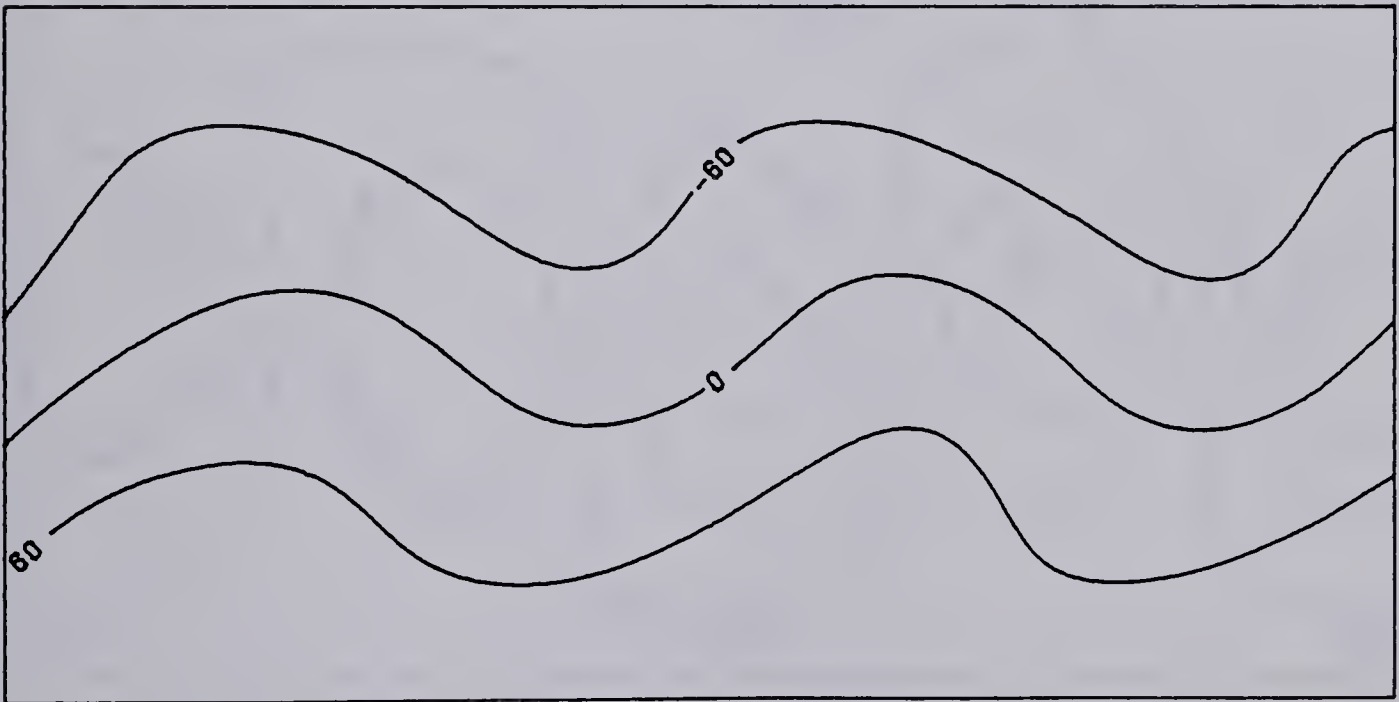
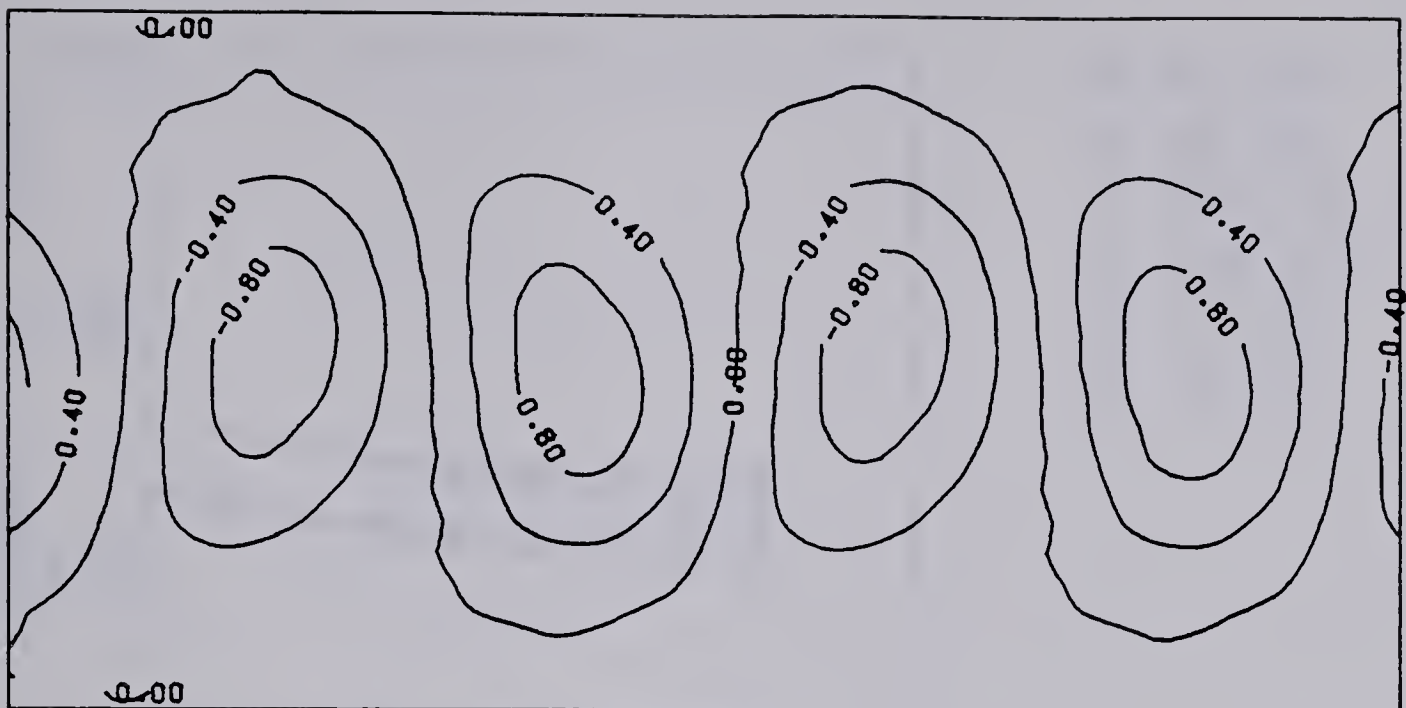
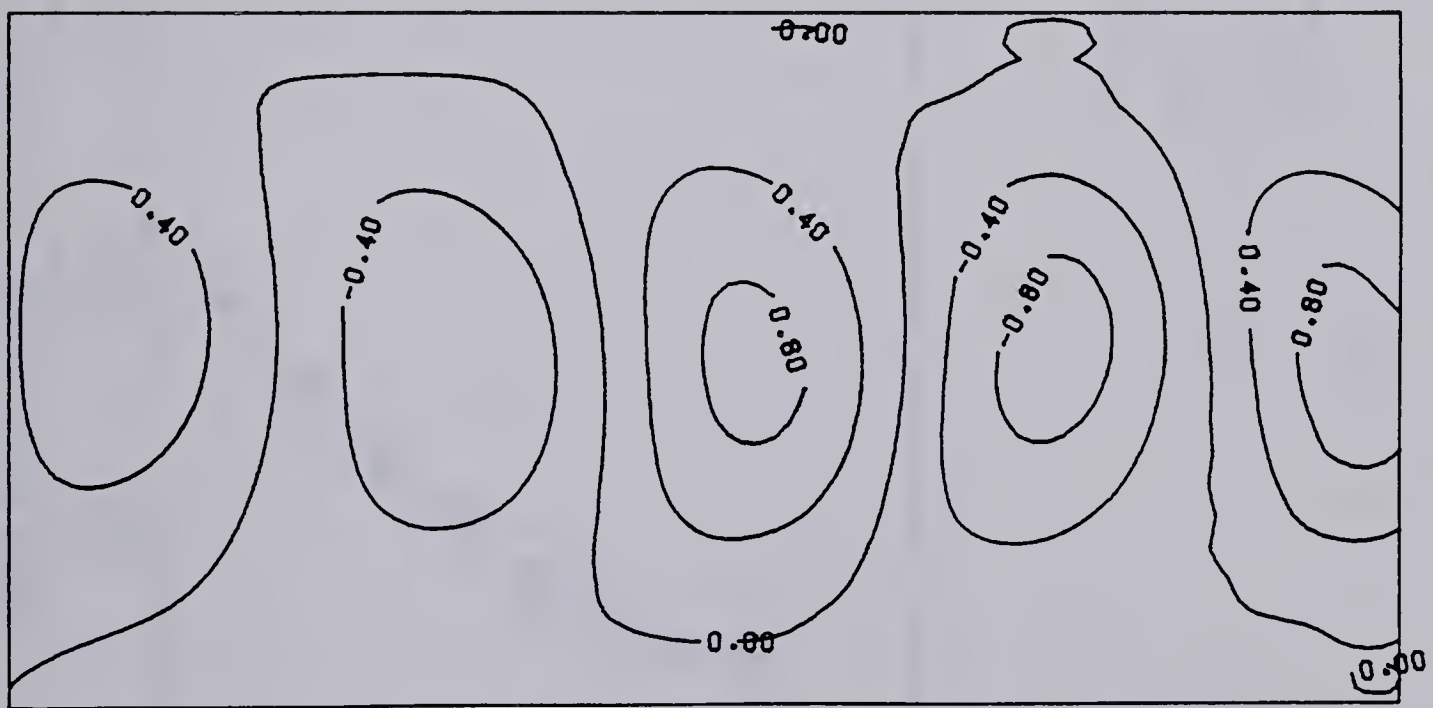


Fig. 5.46 The 48-hour finite-element solution for the thickness field in metres for Case IV in (a) area B and (b) area A.

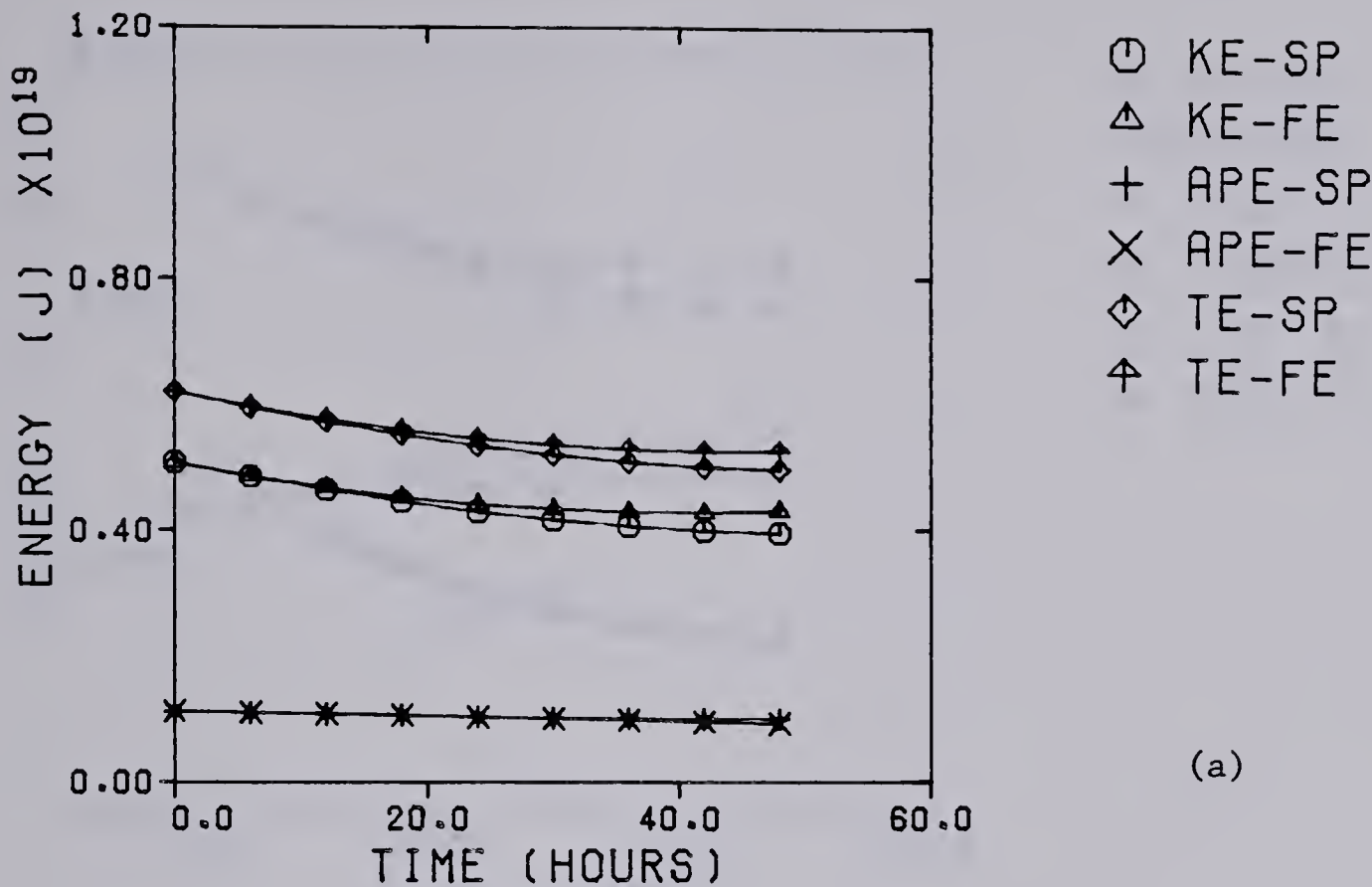


(a)

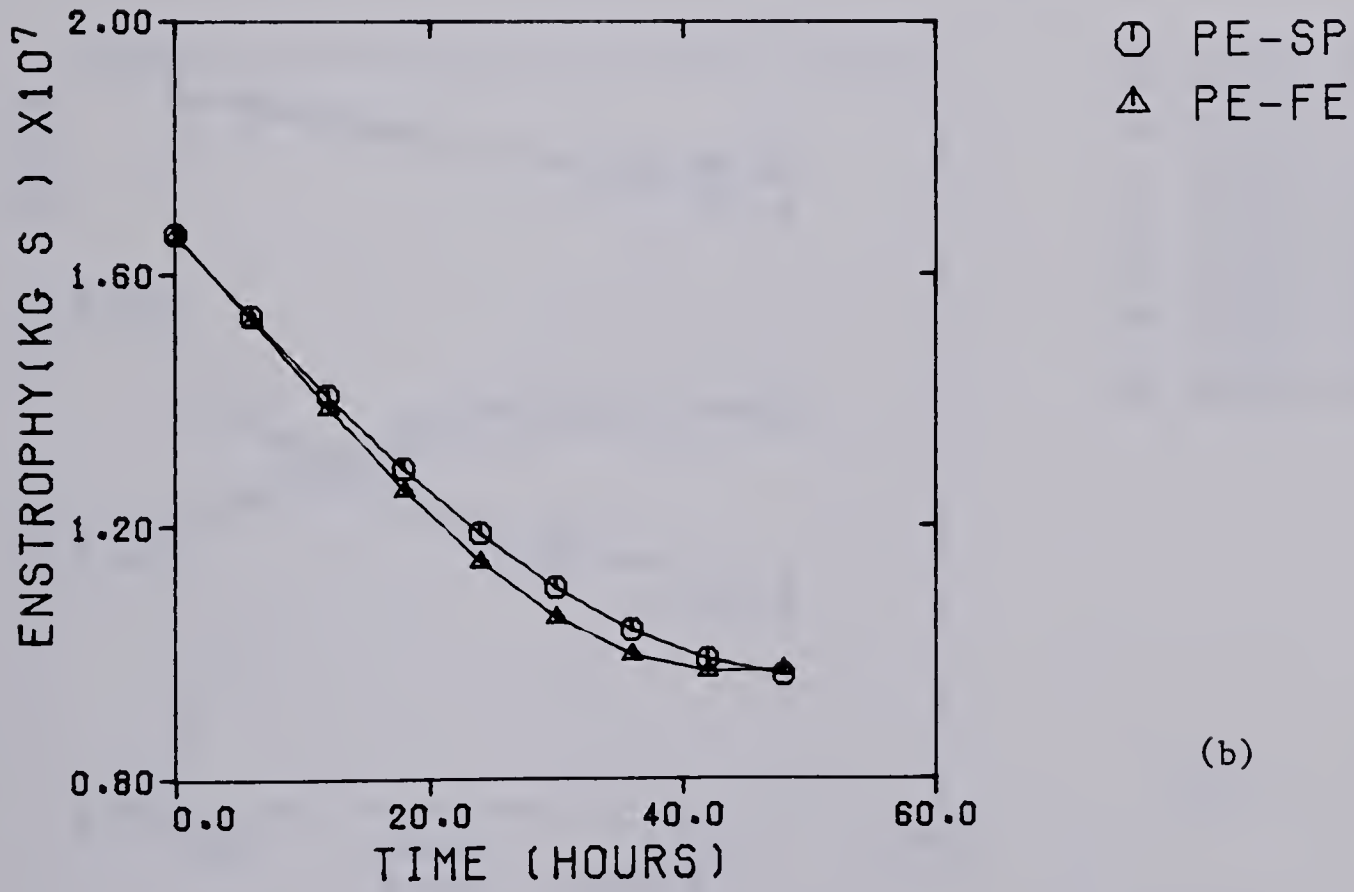


(b)

Fig. 5.47 The 48-hour finite-element solution for the omega field in $\mu\text{bar}/\text{sec}$ for Case IV in (a) area B and (b) area A.

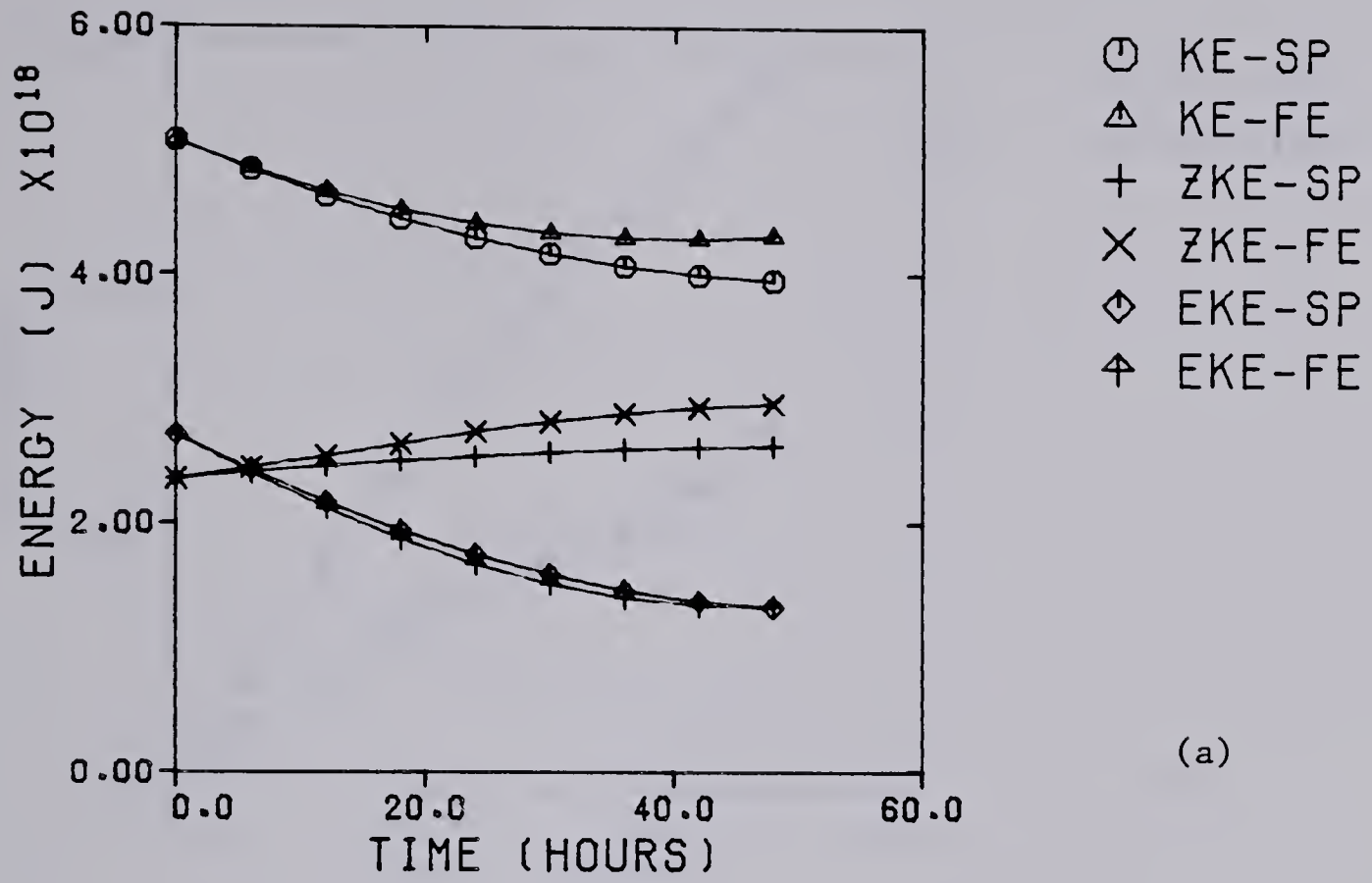


(a)

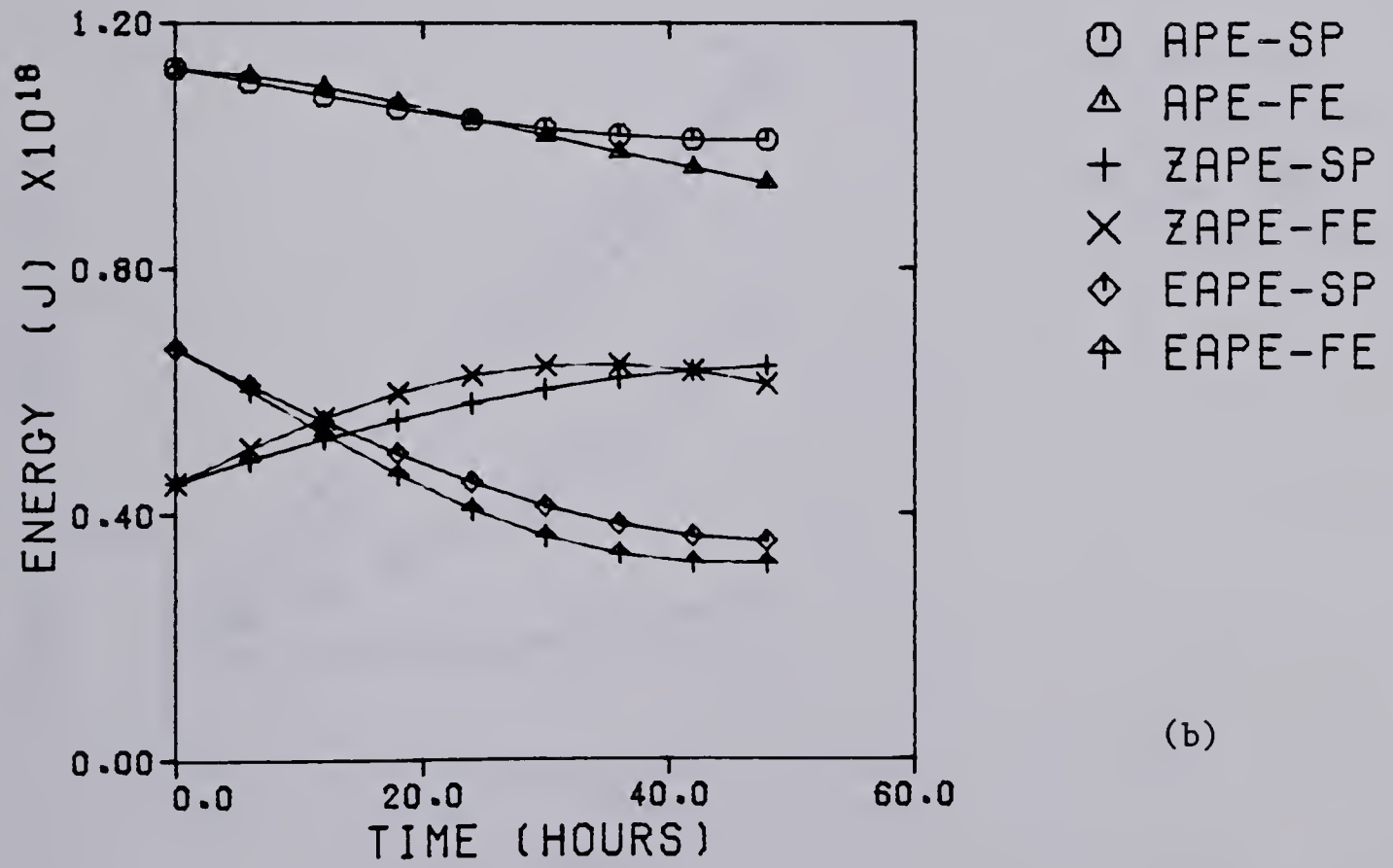


(b)

Fig. 5.48 The (a) energies and (b) potential enstrophy of Case IV for the two solutions in the verification area.

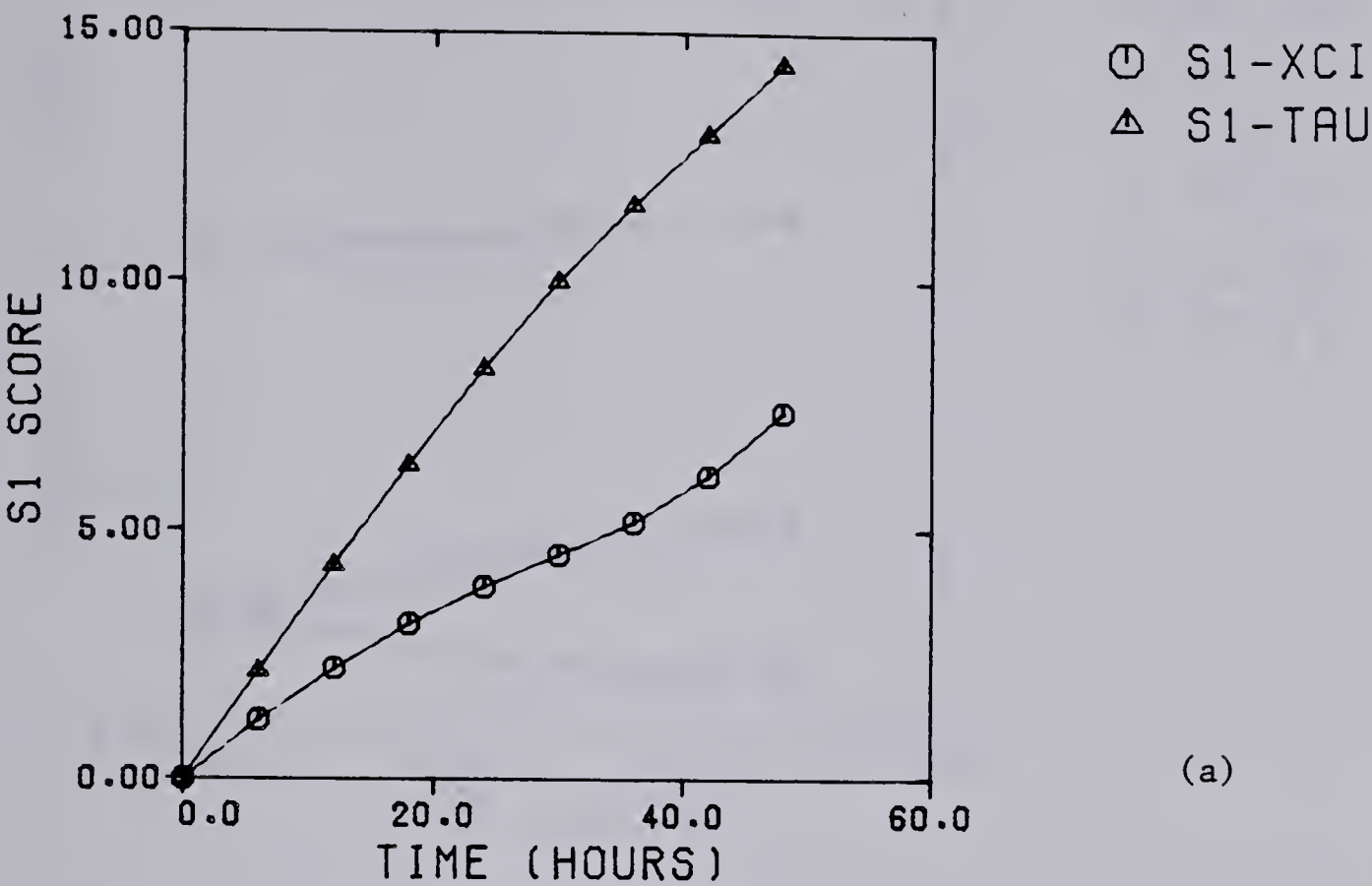


(a)

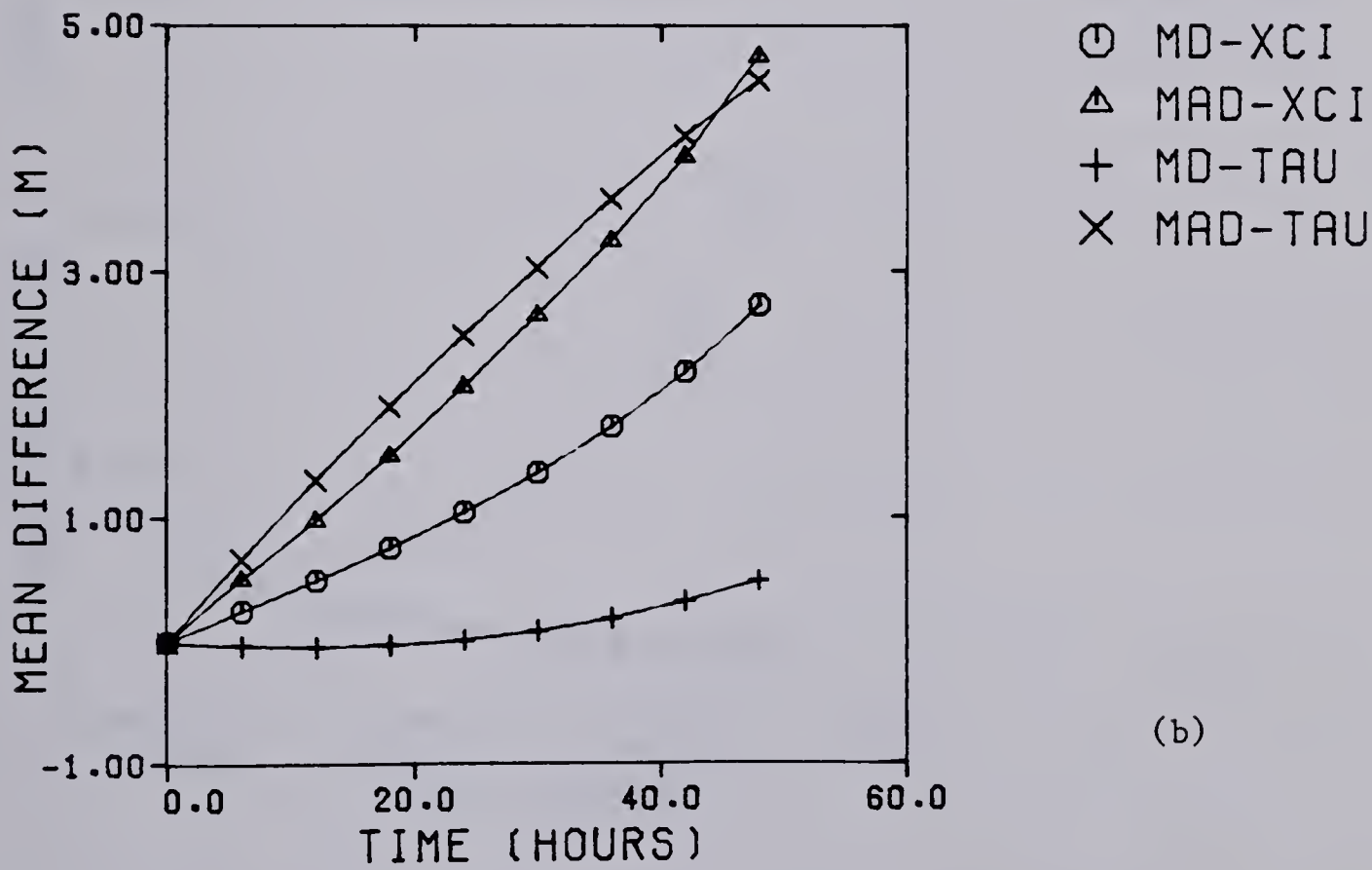


(b)

Fig. 5.49 The (a) kinetic energies and (b) potential energies of Case IV for the two solutions in the verification area.

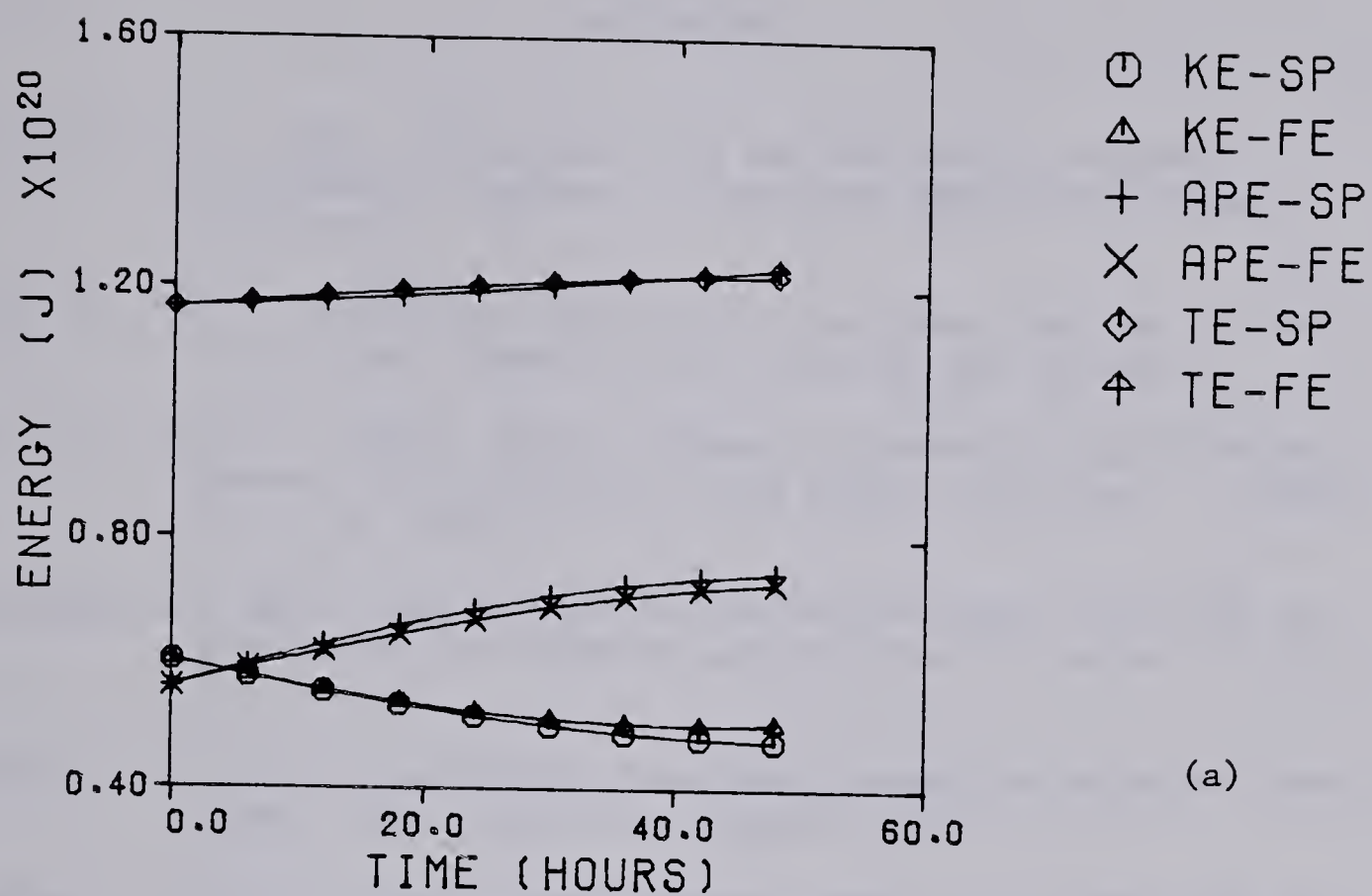


(a)

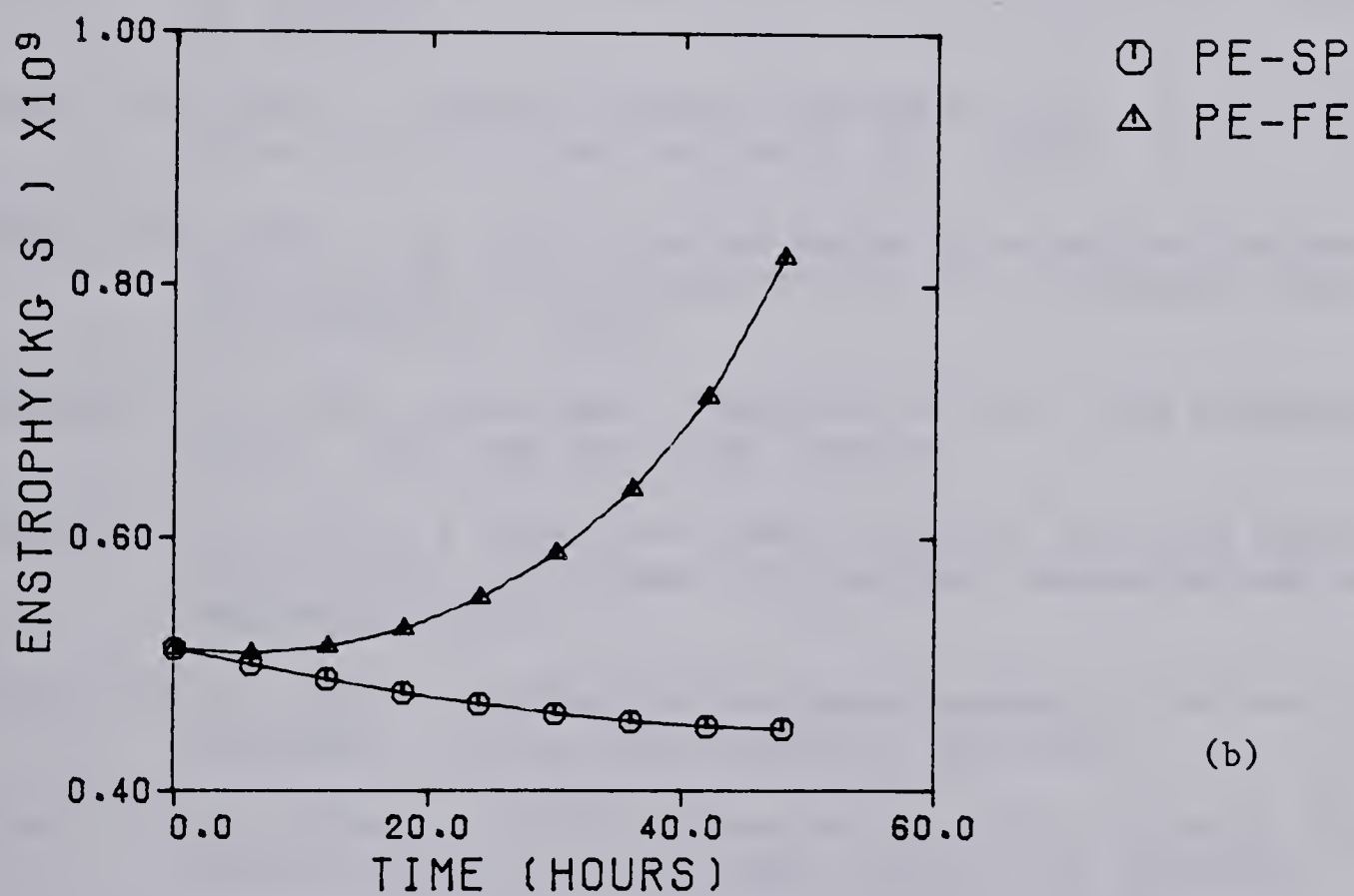


(b)

Fig. 5.50 The (a) S1 scores and (b) MD and MAD curves of Case IV for the two solutions in the verification area.



(a)



(b)

Fig. 5.51 The (a) energies and (b) potential enstrophy of Case IV for the two solutions in the entire domain.

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APPENDIX A

COMPUTER PROGRAM LISTINGS

This appendix contains a listing of the computer programs used and developed for this thesis. In addition, some plotting routines available on the computer were used. All calculations were done using double-precision variables.


```

1  C
2  C
3  C   FILE:      FEMAIN
4  C   FILES:
5  C       OBJECT - FEMAIN.O + INTERP.O + DIFF.O + PROJ.O + PROD.O
6  C       INPUT
7  C           2 = RELDAT
8  C           3 = FEMIN
9  C           4 = MESHD
10 C           5 = VMESH
11 C       OUTPUT
12 C           6 = *PRINT*
13 C           7 = GARBAGE
14 C           8 = FEDIAG
15 C
16 C   PROPER DIMENSIONS
17 C       X(NIU)    ,Y(NJU)    ,XI(NIP)    ,YI(NJP)
18 C       XCI(NI,NJ)    ,TAU(NI,NJ)    ,OMEG(NI,NJ)
19 C       WK1-WK7 DIM. MAX(NI,NJ)
20 C
21 C   THIS IS THE MAIN ROUTINE FOR THE FINITE ELEMENT MODEL.
22 C
23 C
24 C
25 C       IMPLICIT REAL*8 (A-H,O-Z)
26 C       DIMENSION X(281), Y(45), XI(63), YI(27)
27 C       DIMENSION XCI(63,27), TAU(63,27), OMEG(63,27)
28 C       DIMENSION FX(63,27),FY(63,27),FXY(63,27)
29 C       DIMENSION D(4,281,45),WORK(45,8)
30 C       DIMENSION SPAR(6)
31 C       DIMENSION SPARX(6)
32 C       DIMENSION WKA(63,27)
33 C       DIMENSION WKB(63,27)
34 C       DIMENSION WKC(63,27)
35 C       DIMENSION WKD(63,27)
36 C       DIMENSION WKE(63,27)
37 C       DIMENSION WKF(63,27)
38 C       DIMENSION WKG(63,27)
39 C       DIMENSION WKH(63,27)
40 C       DIMENSION WKI(63,27)
41 C       DIMENSION WKJ(63,27)
42 C       DIMENSION WKK(63,27)
43 C       DIMENSION WKL(63,27)
44 C       DIMENSION OUT(63,27),XCI1(63,27),TAU1(63,27)
45 C       DIMENSION XCT2(63,27),TAT2(63,27)
46 C       REAL*8 L,LX,LY,L2,L3
47 C       REAL*8 LYPR,LXPR
48 C       COMMON / WKS1D1 / WK1(63)
49 C       COMMON / WKS1D2 / WK2(63)
50 C       COMMON / WKS1D3 / WK3(63)
51 C       COMMON / WKS1D4 / WK4(63)
52 C       COMMON / WKS1D5 / WK5(63)
53 C       COMMON / WKS1D6 / WK6(63)
54 C       COMMON / WKS1D7 / WK7(63)
55 C       COMMON / WKS1D8 / WK8(63)
56 C       COMMON / WKS1D9 / WK9(63)
57 C       COMMON / WKS1D10 / WK10(63)
58 C       COMMON / WKS1X1 / WK1X(63)
59 C       COMMON / WKS1X2 / WK2X(63)
60 C       COMMON / WKS1X3 / WK3X(63)
61 C       COMMON / WKS1X4 / WK4X(63)
62 C       COMMON / WKS1X5 / WK5X(63)
63 C       COMMON / HXMESH / HX(63)

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```

64      COMMON / HYMESH / HY(27)
65      COMMON / HXMES1 / HX1(63)
66      COMMON / HYMES1 / HY1(27)
67      COMMON / RATIO1 / RAT1(63,27)
68      COMMON / RATI11 / RATI1(63,27)
69      COMMON / RATIO4 / RAT4(63,27)
70      COMMON / RATIOO / RATO(63,27)
71      COMMON / RATIO2 / RAT2(63,27)
72      COMMON / RATIO3 / RAT3(63,27)
73      COMMON / RATIO5 / RAT5(63,27)
74      COMMON / CONSTA / C4,C5
75      LOGICAL FOURTH
76      LOGICAL*1 LFMT(1)/'*/
77      DATA SPAR/.30DO,.50DO,.30DO,.15DO,.08DO,.02DO/
78      DATA SPARX/.10DO,.15DO,.10DO,.05DO,.01DO,.005DO/
79      ICNT=0
80      CALL PLOTS
81      CALL ORGEP(1.0,1.0,1.0)
82      C
83      C SET CONSTANTS
84      C
85      ALP=-12.0DO/5.0DO
86      LX=2.80D7
87      LY=8.80D6
88      PI=3.141592654DO
89      LYPR=2.0DO*PI/LY
90      LXPR=2.0DO*PI/LX
91      FO=1.0312D-4
92      DP=5.00D4
93      BETA=1.62D-11
94      SIG=2.8D-6
95      G=9.806DO
96      L=LX/2.0DO/PI
97      L2=L*L
98      L3=1.0DO/L2
99      ALPHA=LX/LY
100     ALP2=1.0/ALPHA/ALPHA
101     C4=L*L*FO*FO/G
102     C5=DP*FO*10.
103     BETA0=-BETA/FO
104     SIGS=DP*DP/2.0DO/FO/FO/L/L*SIG
105     SIGI=1.0DO/SIGS
106     SIGJ=L2*SIGI
107     OMCT=2.0DO*L*L*FO*FO/(SIG*DP*DP)
108     OMCO=OMCT*L*L
109     BEPR=2.0DO*L*L*FO*BETA/(SIG*DP*DP)
110     HELMCF=2.0DO*FO*FO/(SIG*DP*DP)
111     C
112     C SET FOURTH = .TRUE. FOR FOURTH ORDER SOLUTION AND
113     C FOURTH = .FALSE. FOR SECOND ORDER SOLUTION.
114     C
115     FOURTH =.TRUE.
116     C
117     C READ NO. OF POINTS IN X AND Y DIRNS OF UNDERLYING GRID.
118     C
119     READ(4,LFMT) NIU,NJU
120     C
121     C READ NO. OF PTS IN X AND Y DIRNS. OF NON UNIFORM GRID
122     C
123     READ(4,LFMT) NI,NJ
124     NIP=NI
125     NJP=NJ
126     C
127     C READ NO. OF PTS IN X AND Y DIRNS OF UNIFORM PART OF NON-UNIFORM GRID
128     C
129     READ(4,LFMT) NILA,NJLA

```



```

130      C
131      C READ X AND Y MESH LENGTHS OF UNDERLYING UDIFORM GRID
132      C
133          READ (4,LFMT) HXU,HYU
134      C
135      C READ X AND Y MESH LENGTHS OF UNIFORM PART OF NON-UNIFORM GRID
136      C
137          READ(4,LFMT) HXN,HYN
138      C
139      C READ X-POSN'S OF UNDERLYING GRID
140      C
141          READ(5,LFMT)(X(I),I=1,NIU)
142      C
143      C READ Y-POSN'S OF UNDERLYING GRID
144      C
145          READ(5,LFMT)(Y(I),I=1,NJU)
146      C
147      C READ X-POSN'S OF NON UNIFORM GRID
148      C
149          READ(5,LFMT)(XI(I),I=1,NIP)
150      C
151      C READ Y-POSN'S OF NON UNIFORM GRID
152      C
153          READ(5,LFMT)(YI(I),I=1,NJP)
154      C
155      C READ NO. OF WAVES, NO. OF HOURS FOR PROG., TIMESTEP, NO. OF
156      C HOURS BETWEEN OUTPUT OF MAPS.
157      C
158          READ(3,LFMT) N,M,DT,IDT2
159      C
160      C READ INITIAL VALUES OF HEIGHT, THICKNESS AND OMEGA FIELDS
161      C
162          READ(3,LFMT) XCII1,XCI2,XCI3,TAUU1,TAU2,TAU3,OMEG1,OMEG2,OMEG3
163      C
164      C READ INITIAL HEIGHT CHANGE GUESSES TO SAVE CPU TIME
165      C
166          READ(3,LFMT)DXC1,DXC2,DXC3,DTA1,DTA2,DTA3
167      C
168      C READ RELAXATION PARAMETERS.
169      C
170          READ(2,LFMT) RELF,XTOL
171          READ(2,LFMT) RELFH,XTOLH
172          DT1=DT*.5ODO
173          NIPM=NIP-1
174          NJPM=NJP-1
175          NIM=NI-1
176          NJM=NJ-1
177      C
178      C CALCULATE GRID LENGTHS OF NON UNIFORM GRID
179      C
180          CALL CALH(HX1,XI,NIP)
181          CALL CALH(HY1,YI,NJP)
182          HX1(NIP)=HX1(NIPM)
183          HY1(NJP)=HY1(NJPM)
184          DO 60 I=1,NI
185      60  HX(I)=HX1(I)
186          DO 65 J=1,NJ
187      65  HY(J)=HY1(J)
188      C
189      C FORM RATIO MATRICES FOR SOLUTION
190      C
191          HELP=5.ODO*HELMCF/12.ODO
192          DO 848 J=1,NJ
193      J1=J-1
194          HY2=HY(J)
195          IF(J.EQ.1) GO TO 849

```



```

196      HYJ=HY2+HY(J1)
197      849 DO 848 I=1,NI
198          I1=I-1
199          HX2=HX(I)
200          RAT1(I,J)=HY2/HX2
201          RATI1(I,J)=1.0DO/RAT1(I,J)
202          IF(I.LT.2.OR.J.LT.2) GO TO 848
203          HXI=HX2+HX(I1)
204          RATO(I,J)=HY2*HX2
205          RAT2(I,J)=HY2*HXI
206          RAT3(I,J)=HYJ*HX2
207          RAT4(I,J)=RAT1(I,J)+RAT1(I1,J)+RAT1(I,J1)+RAT1(I1,J1)
208          RAT4(I,J)=RAT4(I,J)+RATI1(I,J)+RATI1(I1,J)+RATI1(I,J1)
209          RAT4(I,J)=RAT4(I,J)+RATI1(I1,J1)
210          RAT5(I,J)=RAT4(I,J)+HELP*HXI*HYJ
211          RAT5(I,J)=1.0DO/RAT5(I,J)
212          RAT4(I,J)=1.0DO/RAT4(I,J)
213      848 CONTINUE
214      C
215      C CALCULATE INITIAL VALUES AT GRID POINTS
216      C
217          XCII1=XCII1*DSQRT(2.0DO)
218          XCI2=XCI2*2.0DO
219          XCI3=XCI3*2.0DO
220          DXC1=DXC1*DSQRT(2.0DO)
221          DXC2=DXC2*2.0DO
222          DXC3=DXC3*2.0DO
223          TAUU1=TAUU1*DSQRT(2.0DO)
224          TAU2=TAU2*2.0DO
225          TAU3=TAU3*2.0DO
226          DTA1=DTA1*DSQRT(2.0DO)
227          DTA2=DTA2*2.0DO
228          DTA3=DTA3*2.0DO
229          OMEG1=OMEG1*DSQRT(2.0DO)
230          OMEG2=OMEG2*2.0DO
231          OMEG3=OMEG3*2.0DO
232          CX=2.0DO*PI*DFLOAT(N)/LX
233          CY=2.0DO*PI/LY
234          SY=0.0DO
235          DO 100 J=1,NJ
236              J1=J-1
237              IF(J.EQ.1) GO TO 110
238              SY=SY+HY(J1)
239      110      Y2=SY*CY
240              SIY=DSIN(Y2)
241              COY=DCOS(Y2)
242              SX=0.0DO
243              DO 100 I=1,NI
244                  I1=I-1
245                  IF(I.EQ.1) GO TO 120
246                  SX=SX+HX(I1)
247      120      X2=SX*CX
248                  SIX=DSIN(X2)
249                  COX=DCOS(X2)
250                  S=SIY*SIX
251                  C=SIY*COX
252                  XCI(I,J)=XCII1*COY+XCI2*S+XCI3*C
253                  TAU(I,J)=TAUU1*COY+TAU2*S+TAU3*C
254                  OMEG(I,J)=OMEG1*COY+OMEG2*S+OMEG3*C
255                  XCI1(I,J)=DXC1*COY+DXC2*S+DXC3*C
256                  TAU1(I,J)=DTA1*COY+DTA2*S+DTA3*C
257      100 CONTINUE
258      C
259      C BEGINNING OF THE TIME LOOP
260      C
261          DO 1000 II=1,M

```



```

262      I2=(II-1)/IDT2*IDT2
263      I3=I2-(II-1)
264      IF(I3.EQ.O) ICNT=ICNT+1
265      I4=(ICNT-1)/4*4-(ICNT-1)
266      C
267      C SET ARRAYS TO O.
268      C
269      C
270      C FORM X-DERIV. OF XCI FIELD.
271      C
272      CALL DXDYDS(WKI,XCI,1.ODO,.TRUE...FALSE.,NI,NJ)
273      CALL PSOLVE(WKA,WKI,NI,NJ,.TRUE...FALSE...FALSE.)
274      C
275      C FORM Y-DERIV. OF XCI FIELD.
276      C
277      CALL DXDYDS(WKI,XCI,1.ODO,.FALSE...TRUE.,NI,NJ)
278      CALL PSOLVE(WKB,WKI,NI,NJ,.FALSE...TRUE...FALSE.)
279      C
280      C FORM X-DERIV. OF TAU FIELD.
281      C
282      CALL DXDYDS(WKI,TAU,1.ODO,.TRUE...FALSE.,NI,NJ)
283      CALL PSOLVE(WKC,WKI,NI,NJ,.TRUE...FALSE...FALSE.)
284      C
285      C FORM Y-DERIV. OF TAU FIELD.
286      C
287      CALL DXDYDS(WKI,TAU,1.ODO,.FALSE...TRUE.,NI,NJ)
288      CALL PSOLVE(WKD,WKI,NI,NJ,.FALSE...TRUE...FALSE.)
289      C
290      C ADD TERM 5 TO WKG, THE R.H.S. OF EQN. 1
291      C
292      CALL GDADGD(WKG,O.ODO,WKG,BETAO,WKA,NI,NJ)
293      C
294      C ADD TERM 5 TO WKH, THE R.H.S. OF EQN. 2
295      C
296      CALL GDADGD(WKH,O.ODO,WKH,BETAO,WKC,NI,NJ)
297      CALL GDADGD(WKL,O.ODO,WKL,-BEPR,WKC,NI,NJ)
298      C
299      C FORM LAPLACIAN OF XCI
300      C
301      CALL D2XYS(WKI,XCI,1.ODO,.TRUE...FALSE.,NI,NJ)
302      CALL PSOLVE(WKE,WKI,NI,NJ,.TRUE...FALSE.,FOURTH)
303      CALL D2XYS(WKI,XCI,1.ODO,.FALSE...TRUE.,NI,NJ)
304      CALL PSOLVE(WKF,WKI,NI,NJ,.FALSE...TRUE.,FOURTH)
305      CALL GDADGD(WKE,1.ODO,WKE,1.ODO,WKF,NI,NJ)
306      CALL SMY(WKE,NI,NJ,SPAR,6,WKK)
307      CALL SMX(WKE,NI,NJ,SPARX,6,WKK)
308      C
309      C FORM Y-DERIV. OF LAPLAC OF XCI.
310      C
311      CALL DXDYDS(WKI,WKE,1.ODO,.FALSE...TRUE.,NI,NJ)
312      CALL PSOLVE(WKF,WKI,NI,NJ,.FALSE...TRUE...FALSE.)
313      C
314      C FORM TERM 1 OF EQN. 1
315      C
316      CALL NLLOOP(WKI,WKA,WKF,NI,NJ)
317      CALL PSOLVE(WKJ,WKI,NI,NJ,.TRUE...TRUE...FALSE.)
318      C
319      C ADD TO R.H.S. OF EQN. 1 (I.E. WKG)
320      C
321      CALL GDADGD(WKG,1.ODO,WKG,-L2,WKJ,NI,NJ)
322      C
323      C FORM TERM 1 OF EQN. 2
324      C
325      CALL NLLOOP(WKI,WKC,WKF,NI,NJ)
326      CALL PSOLVE(WKJ,WKI,NI,NJ,.TRUE...TRUE...FALSE.)
327      C

```



```

328 C ADD TO R.H.S. OF EQN. 2 (I.E. WKH)
329 C
330 CALL GDADGD(WKH,1.ODO,WKH,-L2,WKJ,NI,NJ)
331 CALL GDADGD(WKL,1.ODO,WKL,-OMCO,WKJ,NI,NJ)
332 C
333 C FORM X-DERIV. OF LAPLACIAN OF XCI
334 C
335 CALL DXDYDS(WKI,WKE,1.ODO,.TRUE.,.FALSE.,NI,NJ)
336 CALL PSOLVE(WKE,WKI,NI,NJ,.TRUE.,.FALSE.,.FALSE.)
337 C
338 C FORM TERM 2 OF EQN 1.
339 C
340 CALL NLLOOP(WKF,WKB,WKE,NI,NJ)
341 CALL PSOLVE(WKI,WKF,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
342 C
343 C ADD TO R.H.S. OF EQN. 1
344 C
345 CALL GDADGD(WKG,1.ODO,WKG,L2,WKI,NI,NJ)
346 C
347 C FORM TERM 2 OF EQN. 2.
348 C
349 CALL NLLOOP(WKF,WKD,WKE,NI,NJ)
350 CALL PSOLVE(WKI,WKF,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
351 C
352 C ADD TO R.H.S. OF EQN. 2.
353 C
354 CALL GDADGD(WKH,1.ODO,WKH,L2,WKI,NI,NJ)
355 CALL GDADGD(WKL,1.ODO,WKL,OMCO,WKI,NI,NJ)
356 C
357 C FORM LAPLACIAN OF TAU
358 C
359 CALL D2XYS(WKE,TAU,1.ODO,.TRUE.,.FALSE.,NI,NJ)
360 CALL PSOLVE(WKI,WKE,NI,NJ,.TRUE.,.FALSE.,FOURTH)
361 CALL D2XYS(WKF,TAU,1.ODO,.FALSE.,.TRUE.,NI,NJ)
362 CALL PSOLVE(WKJ,WKF,NI,NJ,.FALSE.,.TRUE.,FOURTH)
363 CALL GDADGD(WKE,1.ODO,WKI,1.ODO,WKJ,NI,NJ)
364 CALL SMY(WKE,NI,NJ,SPAR,6,WKK)
365 CALL SMX(WKE,NI,NJ,SPARX,6,WKK)
366 C
367 C FORM Y-DERIV. OF LAPLACIAN OF TAU
368 C
369 CALL DXDYDS(WKF,WKE,1.ODO,.FALSE.,.TRUE.,NI,NJ)
370 CALL PSOLVE(WKK,WKF,NI,NJ,.FALSE.,.TRUE.,.FALSE.)
371 C
372 C FORM TERM3 OF EQN. 1
373 C
374 CALL NLLOOP(WKI,WKC,WKK,NI,NJ)
375 CALL PSOLVE(WKJ,WKI,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
376 C
377 C ADD TO R.H.S. OF EQN. 1
378 C
379 CALL GDADGD(WKG,1.ODO,WKG,-L2,WKJ,NI,NJ)
380 C
381 C FORM TERM 3 OF EQN. 2
382 C
383 CALL NLLOOP(WKI,WKA,WKK,NI,NJ)
384 CALL PSOLVE(WKJ,WKI,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
385 C
386 C ADD TO R.H.S. OF EQN. 2
387 C
388 CALL GDADGD(WKH,1.ODO,WKH,-L2,WKJ,NI,NJ)
389 CALL GDADGD(WKL,1.ODO,WKL,-OMCO,WKJ,NI,NJ)
390 C
391 C FORM X-DERIV. OF LAPLACIAN OF TAU
392 C
393 CALL DXDYDS(WKF,WKE,1.ODO,.TRUE.,.FALSE.,NI,NJ)

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```

394      CALL PSOLVE(WKJ,WKF,NI,NJ,.TRUE.,.FALSE.,.FALSE.)
395  C
396  C FORM TERM 4 OF EQN. 1
397  C
398      CALL NLLOOP(WKI,WKD,WKJ,NI,NJ)
399      CALL PSOLVE(WKK,WKI,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
400  C
401  C ADD TO R.H.S. OF EQN. 1
402  C
403      CALL GDADGD(WKG,1.ODO,WKG,L2,WKK,NI,NJ)
404  C
405  C FORM TERM 4 OF EQN. 2
406  C
407      CALL NLLOOP(WKI,WKB,WKJ,NI,NJ)
408      CALL PSOLVE(WKJ,WKI,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
409  C
410  C ADD TO R.H.S. OF EQN. 2
411  C
412      CALL GDADGD(WKH,1.ODO,WKH,L2,WKJ,NI,NJ)
413      CALL GDADGD(WKL,1.ODO,WKL,OMCO,WKJ,NI,NJ)
414  C
415  C FIND VERTICAL VELOCITY
416  C
417      CALL NLLOOP(WKI,WKA,WKD,NI,NJ)
418      CALL PSOLVE(WKJ,WKI,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
419      CALL NLLOOP(WKI,WKB,WKC,NI,NJ)
420      CALL PSOLVE(WKK,WKI,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
421      CALL GDADGD(WKF,1.ODO,WKJ,-1.ODO,WKK,NI,NJ)
422      IF(II.EQ.1) GO TO 748
423      IF(I3.NE.O.OR.I4.NE.O) GO TO 748
424      CALL D2XYS(WKE,WKF,1.ODO,.TRUE.,.FALSE.,NI,NJ)
425      CALL PSOLVE(WKI,WKE,NI,NJ,.TRUE.,.FALSE.,FOURTH)
426      CALL D2XYS(WKJ,WKF,1.ODO,.FALSE.,.TRUE.,NI,NJ)
427      CALL PSOLVE(WKK,WKJ,NI,NJ,.FALSE.,.TRUE.,FOURTH)
428      CALL GDADGD(WKE,1.ODO,WKI,1.ODO,WKK,NI,NJ)
429      CALL GDADGD(WKL,1.ODO,WKL,OMCO,WKE,NI,NJ)
430      CALL PROJN(WKI,WKL,FOURTH,NI,NJ,.TRUE.,.TRUE.)
431      DO 226 J=1,NJ
432      DO 226 I=1,NI
433      226 WKI(I,J)=ALP*WKI(I,J)
434      CALL RELHEL(OMEG,WKI,NI,NJ,LX,HELMCF,RELPH,XTOLH)
435      748 CONTINUE
436  C
437  C ADD TERM 6 TO R.H.S. OF EQN. 2
438  C
439      CALL GDADGD(WKH,1.ODO,WKH,OMCT,WKF,NI,NJ)
440  C
441      WRITE(7,999) II
442      999 FORMAT(I6)
443      IF(I4.EQ.O.AND.I3.EQ.O) CALL OUTPUT(II,XCI,TAU,OMEG,XI,
444      1 YI,X,Y,NI,NJ,NIU,NJU,FX,FY,FX,Y,D,WORK,NIP,NJP,OUT,DT)
445      IF(I3.EQ.O) WRITE(8) XCI
446      IF(I3.EQ.O) WRITE(8) TAU
447      IF(II.EQ.M) GO TO 1000
448  C
449  C THE R.H.S. OF EQN. 1 AND 2 HAVE NOW BEEN FORMED AND MAY PROCEED
450  C WITH THE SOLUTION.
451  C
452  C
453      CALL PROJN(WKI,WKG,FOURTH,NI,NJ,.TRUE.,.TRUE.)
454      CALL PROJN(WKF,WKH,FOURTH,NI,NJ,.TRUE.,.TRUE.)
455  C
456  C SOLVE EQN. 1
457  C
458      DO 225 J=1,NJ
459      DO 225 I=1,NI

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460      WKI(I,J)=ALP*WKI(I,J)
461      225 WKF(I,J)=ALP*WKF(I,J)
462      CALL REL(XCI1,WKI,NI,NJ,LX,RELF,XTOL)
463      C
464      C SOLVE EQN. 2
465      C
466      CALL RELHEL(TAU1,WKF,NI,NJ,LX,HELMCF,RELFH,XTOLH)
467      IF(II.EQ.1) GO TO 240
468      C
469      C EXTRAPOLATE IN TIME USING AOAMS-BASHFORTH METHOD.
470      C
471      DO 250 J=1,NJ
472      DO 250 I=1,NI
473      XCI(I,J)=XCI(I,J)+OT1*(3.000*XCI1(I,J)-XCT2(I,J))
474      TAU(I,J)=TAU(I,J)+OT1*(3.000*TAU1(I,J)-TAT2(I,J))
475      250 CONTINUE
476      GO TO 260
477      C
478      C EXTRAPOLATE FIRST TIME STEP USING FORWARD DIFFERENCE.
479      C
480      240 DO 200 J=1,NJ
481      DO 200 I=1,NI
482      XCI(I,J)=XCI(I,J)+DT*XCI1(I,J)
483      TAU(I,J)=TAU(I,J)+DT*TAU1(I,J)
484      200 CONTINUE
485      260 DO 220 J=1,NJ
486      DO 220 I=1,NI
487      XCT2(I,J)=XCI1(I,J)
488      TAT2(I,J)=TAU1(I,J)
489      220 CONTINUE
490      1000 CONTINUE
491      CALL PLOT(0.0,0.0,999)
492      STOP
493      END
494      SUBROUTINE SMY(F,NI,NJ,SPAR,K,FSM)
495      IMPLICIT REAL*8 (A-H,O-Z)
496      DIMENSION F(NI,NJ),FSM(NI,NJ),SPAR(K)
497      COMMON / HYMESH / HY(1)
498      C
499      C SIMPLE 3 PT. SMOOTHER IN Y-DIRECTION
500      C
501      C SPAR=SMOOTHING PARAMETERS FOR 6 GRID POINTS
502      C
503      NJM=NJ-1
504      S2=SPAR(1)
505      S1=1.000-S2
506      DO 10 I=1,NI
507      FSM(I,1)=S1*F(I,1)+S2*F(I,2)
508      10 FSM(I,NJ)=S1*F(I,NJ)+S2*F(I,NJM)
509      DO 20 J=2,6
510      S1=1.000-SPAR(J)
511      J1=J+1
512      J2=J-1
513      DO 20 I=1,NI
514      20 FSM(I,J)=S1*F(I,J)+SPAR(J)*(F(I,J1)+F(I,J2))
515      DO 30 J=2,6
516      JP=NJ-J+1
517      S1=1.000-SPAR(J)
518      J1=JP+1
519      J2=JP-1
520      DO 30 I=1,NI
521      30 FSM(I,JP)=S1*F(I,JP)+SPAR(JP)*(F(I,J1)+F(I,J2))
522      NJ5=NJ-5
523      DO 60 J=1,6
524      DO 60 I=1,NI
525      60 F(I,J)=FSM(I,J)

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526      DO 70 J=NJ5,NJ
527      DO 70 I=1,NI
528      70 F(I,J)=FSM(I,J)
529      RETURN
530      END
531      SUBROUTINE SMX(F,NI,NJ,SPAR,K,FSM)
532      IMPLICIT REAL*8 (A-H,O-Z)
533      DIMENSION F(NI,NJ),FSM(NI,NJ),SPAR(K)
534      COMMON / HXMESH / HX(1)
535      C
536      C SIMPLE 3-POINT SMOOTHER IN X-DIRECTION.
537      C
538      NIM=NI-1
539      DO 10 J=1,6
540      S2=SPAR(J)
541      S1=1.0DO-SPAR(J)
542      FSM(NI,J)=S1*F(NI,J)+S2*F(NIM,J)
543      FSM(1,J)=S1*F(1,J)+S2*F(2,J)
544      JP=NJ-J+1
545      FSM(1,JP)=S1*F(1,JP)+S2*F(2,JP)
546      10 FSM(NI,JP)=S1*F(NI,JP)+S2*F(NIM,JP)
547      DO 20 I=2,NIM
548      I1=I-1
549      I2=I+1
550      DO 20 J=1,6
551      S1=1.0DO-SPAR(J)
552      FSM(I,J)=S1*F(I,J)+SPAR(J)*(F(I2,J)+F(I1,J))
553      JP=NJ-J+1
554      FSM(I,JP)=S1*F(I,JP)+SPAR(J)*(F(I2,JP)+F(I1,JP))
555      20 CONTINUE
556      NJ5=NJ-5
557      DO 60 J=1,6
558      DO 60 I=1,NI
559      60 F(I,J)=FSM(I,J)
560      DO 70 J=NJ5,NJ
561      DO 70 I=1,NI
562      70 F(I,J)=FSM(I,J)
563      RETURN
564      END

```

END OF FILE


```

1      C
2      C
3      C   FILE:  SPEC
4      C
5      C   FILES:
6      C       OBJECT   - SPEC.O
7      C               - SPLATM.O   (USES FOR OUTPUT)
8      C       INPUT
9      C               3 = VMESH
10     C               4 = MESHD
11     C               5 = SPECD   - INITIAL AMPLITUDES AND CONTROL DATA
12     C       OUTPUT
13     C               6 = *PRINT* - OUTPUT OF MAPS AND GRAPHS
14     C               7 = *PRINT* - OUTPUT OF AMPLITUDES
15     C               8 = *PRINT* - ENERGY AND ENSTROPY
16     C               12 = FEMIN   - INITIAL DATA FOR FEM MODEL
17     C               10 = *PRINT* - THEORETICAL CALCULATIONS
18     C               11 = SPDIAG - BINARY OUTPUT FOR DIAGNOSTICS
19     C
20     C   SUBROUTINES:
21     C       - MAIN
22     C       - OUTPT
23     C
24     C   MAIN ROUTINE FOR THE SPECTRAL MODEL.
25     C
26     C
27     C       IMPLICIT REAL*8 (A-H,O-Z)
28     C       DIMENSION A(200,9),EKE(200),APE(200),TEN(200),PE(200)
29     C       DIMENSION PHIH(200),PHIT(200),DPHI(200)
30     C
31     C   PROPER DIMENSIONS FOR ARRAYS:
32     C
33     C   DIM X(NIU),Y(NJU),XCIB(NIU,NJU),TAUB(NIU,NJU),OMEGB(NIU,NJU)
34     C   DIM XI(NI),YI(NI)
35     C
36     C       DIMENSION X(281),Y(45),XCIB(281,45),TAUB(281,45),OMEGB(281,45)
37     C       DIMENSION XI(63),YI(27),HX1(63),HY1(27)
38     C       DIMENSION F(63,27),HX(63),HY(27)
39     C       REAL*8 L,LX,LY,LXPR,LYPR
40     C       COMMON TAU1(200),TAU2(200),TAU3(200),XCI1(200),XCI2(200)
41     C       COMMON XCI3(200),OMEG1(200),OMEG2(200),OMEG3(200)
42     C       COMMON C4,C5,DT
43     C       LOGICAL*1 LFMT(1)/'*/
44     C
45     C       READ INITIAL VALUES
46     C
47     C       READ(5,LFMT) XCI1(1),XCI2(1),XCI3(1),TAU1(1),TAU2(1),TAU3(1)
48     C
49     C       N=NUMBER OF WAVES, M=NUMBER OF TIME STEPS, DT=TIME STEP
50     C       IDT2 = NO. OF HOURS BETWEEN OUTPUT OF MAPS.
51     C
52     C       READ(5,LFMT) N,M,DT,IDT2
53     C       WRITE(7,22) N
54     C 22   FORMAT('      NUMBER OF WAVES IS      ',I4)
55     C
56     C       SET VALUE OF CONSTANTS
57     C
58     C       FN=DFLOAT(N)
59     C       ICNT=0
60     C       CALL PLOTS
61     C       CALL ORGEP(1.0,1.0,1.0)
62     C       PI=3.141592700
63     C       FO=1.0312D-4

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64      BETA=1.62D-11
65      DP=5.00D4
66      SIG=2.8D-6
67      LY=8.80D6
68      LX=2.80D7
69      L=LX/2.00D0/PI
70      LXPR=2.00D0*PI*FN/LX
71      LYPR=2.*PI/LY
72      ALPHA=LX/LY
73      A2=ALPHA*ALPHA
74      BSTAR=L*BETA/FO
75      SIGST=DP*DP*SIG/(2.00D0*FO*FO*L*L)
76      G=9.806D0
77      GAM=1.2004D0
78      ITIME=0
79      GAN=GAM*ALPHA*FN
80      TO=A2+FN*FN
81      T1=1.00D0/A2+SIGST
82      T2=SIGST+1.00D0/TO
83      T3=A2+1.00D0/SIGST
84      T4=TO+1.00D0/SIGST
85      A2=ALPHA*ALPHA
86      C1=2.00D0*PI*PI*DP*FO*FO*L**4/(G*ALPHA*SIGST)
87      C6=L**6*PI**4*DP*FO*FO*8.00D0/(G*ALPHA)
88      C3=DP*2.00D0*PI*L*L/G/ALPHA
89      C4=L*L*FO*FO/G
90      C5=DP*FO*10.00D0
91      M1=M+1
92      Q2=2.00D0*FO*FO/(SIG*DP*DP)
93      C7=4.00D0*DSQRT(2.00D0)*L*L*FO/LY
94      C8=2.00D0*PI*FN/LX
95      CR=BETA/C8/C8
96      C9=2.00D0*(1.00D0+Q2/C8/C8)
97      WRITE(12,474) N,M,DT,IDT2
98 474  FORMAT(2I6,4X,E12.6,I5)
99  C
100  C  GET PARAMETERS FOR F.E.M. MODEL TO MAKE OUPUT CONSISTENT
101  C
102      READ(4,LFMT) NIU,NJU
103      READ(4,LFMT) NI,NJ
104      READ(4,LFMT) NILA,NJLA
105      READ(4,LFMT) HXU,HYU
106      READ(3,LFMT)(X(I),I=1,NIU)
107      READ(3,LFMT)(Y(I),I=1,NJU)
108      READ(3,LFMT)(XI(I),I=1,NI)
109      READ(3,LFMT)(YI(I),I=1,NJ)
110      CALL CALH(HX1,XI,NI)
111      CALL CALH(HY1,YI,NJ)
112      NIM=NI-1
113      NJM=NJ-1
114      HX1(NI)=HX1(NIM)
115      HY1(NI)=HY1(NJM)
116      NIP=NI
117      NJP=NJ
118      NIM=NI-1
119      NJM=NJ-1
120      DO 60 I=1,NIP
121 60  HX(I)=HX1(I)
122      DO 65 J=1,NJP
123 65  HY(J)=HY1(J)
124      WRITE(10,475) CR
125 475  FORMAT(' ROSSBY PHASE SPEED = ',E12.4,/)
126      WRITE(10,470)
127 470  FORMAT(4X,'ITIME',3X,'THERMAL WIND      MEAN WIND ',10X,
128      1'REAL PHASE SPEEDS',6X,'IMAGINARY SPEED',4X,'E-TIME')
129      WRITE(8,12)

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130      12  FORMAT(3X,'TIME',10X,'APE',13X,'KE',14X,'TEN',13X,'PE')
131      WRITE(7,13)
132      13  FORMAT('      TIME',6X,'XCI1',9X,'XCI2',9X,'XCI3',9X,
133      1  'TAU1',9X,'TAU2',9X,'TAU3',9X,'OMEG1',8X,'OMEG2',8X,'OMEG3')
134      DO 500 I=1,M
135          ITIME=I-1
136      C
137      C      FIND THE VERTICAL VELOCITIES
138      C
139          OMEG1(I)=-GAN/T1*(TAU2(I)*XCI3(I)-TAU3(I)*XCI2(I))
140          OMEG2(I)=(((TAU1(I)*XCI3(I)+TAU3(I)*XCI1(I))*FN*FN/TO
141      1  -(TAU3(I)*XCI1(I)-TAU1(I)*XCI3(I)))*GAN
142      2  -BSTAR*FN*TAU3(I)/TO)/T2
143          OMEG3(I)=(((TAU2(I)*XCI1(I)-TAU1(I)*XCI2(I))
144      1  -(TAU1(I)*XCI2(I)+TAU2(I)*XCI1(I))*FN*FN/TO)*GAN
145      2  +BSTAR*FN*TAU2(I)/TO)/T2
146          IF(I.EQ.1) WRITE(12,1000) XCI1(1),XCI2(1),XCI3(1),TAU1(1),
147      1  TAU2(1),TAU3(1),OMEG1(1),OMEG2(1),OMEG3(1)
148      1000 FORMAT(9(1X,E12.4))
149          WRITE(7,10) ITIME,XCI1(I),XCI2(I),XCI3(I),TAU1(I),TAU2(I),
150      1  TAU3(I),OMEG1(I),OMEG2(I),OMEG3(I)
151      10  FORMAT(3X,I4,9(1X,E12.4))
152      C
153      C      CALCULATE PHASES
154      C
155          PHIH(I)=DATAN2(XCI2(I),XCI3(I))*180.0DO/PI
156          PHIT(I)=DATAN2(TAU2(I),TAU3(I))*180.0DO/PI
157          DPHI(I)=PHIH(I)-PHIT(I)
158      C
159      C      CALCULATE THEORETICAL PARAMETERS
160      C
161          UT=C7*TAU1(I)
162          UST=C7*XCI1(I)
163          D=(Q2*CR/C8/C8)**2+4.0DO*(1.0DO-(Q2/C8/C8)**2)*UT*UT
164          CTEMP=-(2.0DO+Q2/C8/C8)*CR
165          IF(D.GE.0.0DO) GO TO 400
166      C
167      C      IMAGINARY PHASE VELOCITY
168      C
169          CREA2=0.0DOO
170          CIM=DSQRT(-D)/C9
171          CREA1=CTEMP/C9
172          ETAU=1.0DO/CIM/C8
173          GO TO 450
174      400 CREA1=(CTEMP+DSQRT(D))/C9
175          CREA2=(CTEMP-DSQRT(D))/C9
176          CIM=0.0DO
177          ETAU=0.0DO
178      450 WRITE(10,460) I,UT,UST,CREA1,CREA2,CIM,ETAU
179      460 FORMAT(3X,I4,6(4X,E12.4))
180      C
181      C      CALCULATE ENERGIES.
182      C
183          APE(I)=TAU1(I)*TAU1(I)+TAU2(I)*TAU2(I)+TAU3(I)*TAU3(I)
184          APE(I)=APE(I)*C1
185          EKE(I)=XCI2(I)*XCI2(I)+XCI3(I)*XCI3(I)+TAU2(I)*TAU2(I)
186      1  +TAU3(I)*TAU3(I)
187          EKE(I)=(EKE(I)*(1.0DO/LY/LY+FN*FN/LX/LX)+(XCI1(I)*XCI1(I)
188      1  +TAU1(I)*TAU1(I))/LY/LY)
189          EKE(I)=EKE(I)*C6
190          TEN(I)=APE(I)+EKE(I)
191      C
192      C      CALCULATE POTENTIAL ENSTROPHY.
193      C
194          Q1=(DABS(A2*XCI1(I)+T3*TAU1(I)))**2
195      2  +(DABS(TO*XCI2(I)+T4*TAU2(I)))**2

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196      1 +(DABS(TO*XCI3(I)+T4*TAU3(I)))*2
197      Q1=LX*LY*Q1/2.ODO
198      Q3=(DABS(A2*XCI1(I)-T3*TAU1(I)))*2
199      2 +(DABS(TO*XCI2(I)-T4*TAU2(I)))*2
200      1 +(DABS(TO*XCI3(I)-T4*TAU3(I)))*2
201      Q3=LX*LY*Q3/2.ODO
202      PE(I)=(Q1+Q3)*C3
203      WRITE(8,11)ITIME, APE(I),EKE(I),TEN(I),PE(I)
204      11  FORMAT(3X,I4,4(4X,E12.4))
205      I2=(I-1)/IDT2*IDT2
206      I3=I2-(I-1)
207      IF(I3.EQ.O) ICNT=ICNT+1
208      I4=(ICNT-1)/4*4-(ICNT-1)
209      IF(I4.EQ.O.AND.I3.EQ.O) CALL OUTPT(I,X,Y,NIU,NJU,LXPR,
210      1 LYPR,HXU,HYU,XCIB,TAUB,OMEGB,DT)
211      IF(I3.EQ.O) CALL FOUT(F,NI,NJ,HX,HY,XCI1(I),XCI2(I),XCI3(I),
212      1 LXPR,LYPR)
213      IF(I3.EQ.O) CALL FOUT(F,NI,NJ,HX,HY,TAU1(I),TAU2(I),TAU3(I),
214      1 LXPR,LYPR)
215      C
216      C      BEGIN INTEGRATION
217      C
218      TIME=I*DT
219      IF(I.NE.1) GO TO 100
220      C
221      C      FIRST STEP:  USE EULER
222      TAU1(2)=TAU1(1)-OMEG1(1)*DT/ALPHA/ALPHA
223      T1BAR=(TAU1(1)+TAU1(2))/2.ODO
224      TAU2(2)=TAU2(1)+((T1BAR*XCI3(1)+TAU3(1)*XCI1(1))*FN*FN*GAN/TO
225      1 -BSTAR*FN*TAU3(1)/TO-OMEG2(1)/TO)*DT
226      T2BAR=(TAU2(1)+TAU2(2))/2.ODO
227      TAU3(2)=TAU3(1)+(-(T1BAR*XCI2(1)+T2BAR*XCI1(1))*FN*FN*GAN/TO
228      1 +BSTAR*FN*T2BAR/TO-OMEG3(1)/TO)*DT
229      XCI1(2)=XCI1(1)
230      T3BAR=(TAU3(1)+TAU3(2))/2.ODO
231      XCI2(2)=XCI2(1)+((XCI1(1)*XCI3(1)+T1BAR*T3BAR)
232      1 *GAN*FN*FN/TO-BSTAR*FN*XCI3(1)/TO)*DT
233      X2BAR=(XCI2(1)+XCI2(2))/2.ODO
234      XCI3(2)=XCI3(1)+(-(XCI1(1)*X2BAR+T1BAR*T2BAR)
235      1 *GAN*FN*FN/TO+BSTAR*FN*X2BAR/TO)*DT
236      GO TO 500
237      100 CONTINUE
238      I1=I+1
239      I2=I-1
240      C
241      C      USE ADAMS-BASHFORTH METHOD FOR SUBSEQUENT TIME STEPS
242      C
243      XCI1(I1)=XCI1(I)
244      TAU1(I1)=TAU1(I)-(3.ODO*OMEG1(I)/ALPHA/ALPHA/2.ODO
245      1 -OMEG1(I2)/ALPHA/ALPHA/2.)*DT
246      C=(TAU1(I)*XCI3(I)+TAU3(I)*XCI1(I))*GAN*FN*FN/TO
247      1 -BSTAR*FN*TAU3(I)/TO-OMEG2(I)/TO
248      C2=(TAU1(I2)*XCI3(I2)+TAU3(I2)*XCI1(I2))*GAN*FN*FN/TO
249      1 -BSTAR*FN*TAU3(I2)/TO-OMEG2(I2)/TO
250      TAU2(I1)=TAU2(I)+(3.ODO*C/2.ODO-C2/2.ODO)*DT
251      C=-(TAU1(I)*XCI2(I)+TAU2(I)*XCI1(I))*GAN*FN*FN/TO
252      1 +BSTAR*FN*TAU2(I)/TO-OMEG3(I)/TO
253      C2=-(TAU1(I2)*XCI2(I2)+TAU2(I2)*XCI1(I2))*GAN*FN*FN/TO
254      1 +BSTAR*FN*TAU2(I2)/TO-OMEG3(I2)/TO
255      TAU3(I1)=TAU3(I)+(3.ODO*C/2.ODO-C2/2.ODO)*DT
256      C=XCI1(I)*XCI3(I)*GAN*FN*FN/TO-BSTAR*FN*XCI3(I)/TO
257      C2=XCI1(I2)*XCI3(I2)*GAN*FN*FN/TO-BSTAR*FN*XCI3(I2)/TO
258      T=(TAU1(I)*TAU3(I)+TAU1(I1)*TAU3(I1))*GAN*FN*FN/TO/2.
259      XCI2(I1)=XCI2(I)+(3.ODO*C/2.ODO-C2/2.ODO+T)*DT
260      C      FROM EQN. 12
261      C=XCI1(I)*XCI2(I)+XCI1(I1)*XCI2(I1)

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262      1 +TAU1(I)*TAU2(I)+TAU1(I1)*TAU2(I1)
263      C2=XCI2(I)+XCI2(I1)
264      XCI3(I1)=XCI3(I)-C*GAN*FN*FN/TO*DT/2.ODO+
265      1 C2*BSTAR*FN/TO*DT/2.ODO
266      500 CONTINUE
267      DXC1=(XCI1(2)-XCI1(1))/DT
268      DXC2=(XCI2(2)-XCI2(1))/DT
269      DXC3=(XCI3(2)-XCI3(1))/DT
270      DTA1=(TAU1(2)-TAU1(1))/DT
271      DTA2=(TAU2(2)-TAU2(1))/DT
272      DTA3=(TAU3(2)-TAU3(1))/DT
273      WRITE(12,1000) DXC1,DXC2,DXC3,DTA1,DTA2,DTA3
274      C
275      C PLOT OUTPUT PARAMETERS
276      C
277      DO 600 I=1,M
278      A(I,1)=XCI1(I)
279      A(I,2)=XCI2(I)
280      A(I,3)=XCI3(I)
281      A(I,4)=TAU1(I)
282      A(I,5)=TAU2(I)
283      A(I,6)=TAU3(I)
284      A(I,7)=OMEG1(I)
285      A(I,8)=OMEG2(I)
286      A(I,9)=OMEG3(I)
287      600 CONTINUE
288      CALL SPLATM(A,200,9,9,M1,2,-.5D-1,.5D-1)
289      C
290      C PLOT ENERGYS
291      C
292      DO 620 I=1,M
293      A(I,1)=EKE(I)
294      A(I,2)=APE(I)
295      A(I,3)=TEN(I)
296      620 CONTINUE
297      CALL SPLATM(A,200,9,3,M1,2,0.ODO,0.ODO)
298      C
299      C PLOT PHASES
300      C
301      DO 640 I=1,M1
302      A(I,1)=PHIH(I)
303      A(I,2)=PHIT(I)
304      A(I,3)=DPHI(I)
305      640 CONTINUE
306      CALL SPLATM(A,200,9,3,M1,2,-180.ODO,180.ODO)
307      CALL PLOT(0.0,0.0,999)
308      STOP
309      END
310      SUBROUTINE OUTPT(I,X,Y,NIU,NJU,LXPR,LYPR,HXU,HYU,XCIB,TAUB,
311      1 OMEGB,DT)
312      C
313      C SET UP FIELDS TO BE PLOTTED IN MANNER CONSISTENT WITH FEM
314      C
315      IMPLICIT REAL*8 (A-H,O-Z)
316      DIMENSION XCIB(NIU,NJU),TAUB(NIU,NJU),OMEGB(NIU,NJU)
317      DIMENSION X(NIU),Y(NJU)
318      REAL*8 LXPR,LYPR
319      COMMON TAU1(200),TAU2(200),TAU3(200),XCI1(200),XCI2(200)
320      COMMON XCI3(200),OMEG1(200),OMEG2(200),OMEG3(200)
321      COMMON C4,C5
322      RT2=DSQRT(2.ODO)
323      DO 20 KI=1,NIU
324      XX=LXPR*(X(KI)-HXU)
325      CX=DCOS(XX)
326      SX=DSIN(XX)
327      DO 20 KJ=1,NJU

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328      YY=LYPR*(Y(KJ)-HYU)
329      CY=DCOS(YY)
330      SY=DSIN(YY)
331      S=SX*SY
332      C=CX*SY
333      XCIB(KI,KJ)=RT2*CY*XC11(I)+2.ODO*S*XC12(I)+2.ODO*C*XC13(I)
334      TAUB(KI,KJ)=RT2*CY*TAU1(I)+2.ODO*S*TAU2(I)+2.ODO*C*TAU3(I)
335      OMEGB(KI,KJ)=RT2*CY*OMEG1(I)+2.ODO*S*OMEG2(I)+2.ODO*C*OMEG3(I)
336      XCIB(KI,KJ)=C4*XCIB(KI,KJ)
337      TAUB(KI,KJ)=C4*TAUB(KI,KJ)
338      OMEGB(KI,KJ)=C5*OMEGB(KI,KJ)
339      20 CONTINUE
340      CALL FPLLOT(XCIB,NIU,NJU,1)
341      CALL FPLLOT(TAUB,NIU,NJU,1)
342      CALL FPLLOT(OMEGB,NIU,NJU,0)
343      RETURN
344      END
345      SUBROUTINE FOUT(F,NI,NJ,HX,HY,F1,F2,F3,LXPR,LYPR)
346      IMPLICIT REAL*8 (A-H,O-Z)
347      C
348      C  OUTPUT FIELDS FOR DIAGNOSTIC PROGRAM
349      C
350      DIMENSION F(NI,NJ),HX(NI),HY(NJ)
351      REAL*8 LXPR,LYPR
352      F11=F1*DSQRT(2.ODO)
353      F22=F2*2.ODO
354      F33=F3*2.ODO
355      SY=0.ODO
356      DO 100 J=1,NJ
357      J1=J-1
358      IF(J1.EQ.0) GO TO 110
359      SY=SY+HY(J1)
360      110 Y2=SY*LYPR
361      SIY=DSIN(Y2)
362      COY=DCOS(Y2)
363      SX=0.ODO
364      DO 100 I=1,NI
365      I1=I-1
366      IF(I1.EQ.0) GO TO 120
367      SX=SX+HX(I1)
368      120 X2=SX*LXPR
369      SIX=DSIN(X2)
370      COX=DCOS(X2)
371      S=SIY*SIX
372      C=SIY*COX
373      F(I,J)=F11*COY+F22*S+F33*C
374      100 CONTINUE
375      WRITE(11) F
376      RETURN
377      END
END OF FILE

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1      C      FILE:  DIFF
2      C      OBJECT FILE:  DIFF.O
3      C      SUBROUTINES:
4      C          - DXDYDS
5      C          - CALH
6      C          - SETD2
7      C          - D2XYS
8      C
9      C
10     C
11     SUBROUTINE DXDYDS(R,U,CON,ALONGX,ALONGY,NI,NJ)
12     C
13     C      S/R : DXDYDS - CALCULATES A DIFFERENCE IN X AND Y DIRECTIONS
14     C
15     C      AUTHDR : ANDREW STANIFORTH
16     C
17     C      REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
18     C
19     C      ARGUMENTS :
20     C          OUT - R          - RESULT
21     C          IN  - U          - FIELD TO BE DIFFERENCED
22     C              - CDN        - MULTIPLICATIVE CONSTANT (MULTIPLIES DERIVATIVE)
23     C              - ALONGX     - IF .TRUE. DIFFERENCE IN X DIRECTION
24     C              - ALONGY     - IF .TRUE. DIFFERENCE IN Y DIRECTION
25     C              - NI        - X DIMENSION
26     C              - NJ        - Y DIMENSION
27     C
28     C
29     C
30     IMPLICIT REAL*8 (A-H,O-Z)
31     DIMENSION R(NI,NJ),U(NI,NJ)
32     LDGICAL ALONGX,ALONGY
33     CDN=.5ODO*CON
34     NIM=NI-1
35     NJM=NJ-1
36     IF(ALONGX) GO TO 100
37     IF(ALONGY) GO TO 200
38     RETURN
39     C
40     C      X - DERIVATIVE
41     C
42     100 CONTINUE
43     IT=NI-2
44     DO 150 J=1,NJ
45     DD 140 I=1,IT
46     I1=I+1
47     I2=I+2
48     140 R(I1,J)=CDNX*(U(I2,J)-U(I,J))
49     R(1,J)=CONX*(U(2,J)-U(1,J))
50     R(NI,J)=CONX*(U(NI,J)-U(NIM,J))
51     150 CONTINUE
52     RETURN
53     C
54     C      Y - DERIVATIVE
55     C
56     200 CONTINUE
57     IT=NJ-2
58     DD 250 I=1,NI
59     DD 240 J=1,IT
60     J1=J+1
61     J2=J+2
62     240 R(I,J1)=CONX*(U(I,J2)-U(I,J))
63     R(I,1)=CONX*(U(I,2)-U(I,1))

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64      R(I,NJ)=CONX*(U(I,NJ)-U(I,NJM))
65      250 CONTINUE
66      RETURN
67      END
68      SUBROUTINE CALH(H,X,M)
69      C
70      C S/R: CALH - CALCULATION OF GRID LENGTHS FROM POINTS X.
71      C           H(I)=X(I+1)-X(I), I=1,M-1
72      C
73      C AUTHOR: A. STANIFORTH - 1973
74      C
75      C REVISION OO1: A. STANIFORTH - C. THIBEAULT JAN 79 DOCUMENTATION
76      C REVISION OO2: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
77      C
78      C CALL: CALL CALH(H,X,M)
79      C
80      C ARGUMENTS:
81      C OUT - H - GRID-LENGTHS
82      C IN - X - ARRAY OF POINTS IN ASCENDING ORDER
83      C - M - LENGTH OF ARRAYS H AND X
84      C
85      C NOTES: - H(M) IS NOT CALCULATED
86      C
87      C
88      C IMPLICIT REAL*8 (A-H,O-Z)
89      C DIMENSION X(M),H(M)
90      C DO 10 I=2,M
91      C I1=I-1
92      C H(I1)=X(I)-X(I1)
93      C 10 CONTINUE
94      C RETURN
95      C END
96      C SUBROUTINE SETD2(A,B,C,CON,H,N)
97      C
98      C S/R : SETD2 - SET-UP ELEMENTS OF TRI-DIAGONAL MATRIX USED IN S/R
99      C D2XYS FOR TAKING SECOND DERIVATIVES.
100     C
101     C AUTHOR - ANDREW STANIFORTH - JAN 79
102     C
103     C REVISION OO1: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
104     C
105     C ARGUMENTS:
106     C OUT - A - LOWER-DIAGONAL ELEMENTS
107     C - B - DIAGONAL ELEMENTS
108     C - C - UPPER-DIAGONAL ELEMENTS
109     C - CON - MULTIPLICATIVE FACTOR (CON*SECOND DERIVATIVE)
110     C - H - MESH-LENGTHS
111     C - N - NUMBER OF NODAL POINTS
112     C
113     C
114     C
115     C IMPLICIT REAL*8 (A-H,O-Z)
116     C DIMENSION A(N),B(N),C(N),H(N)
117     C HR=CON/H(1)
118     C B(1)=0.0DO
119     C NM=N-1
120     C DO 100 I=1,NM
121     C I1=I+1
122     C HR=CON/H(I)
123     C A(I1)=HR
124     C B(I)=B(I)-HR
125     C B(I1)=-HR
126     C 100 C(I)=HR
127     C RETURN
128     C END
129     C SUBROUTINE D2XYS(R,U,CON,ALONGX,ALONGY,NI,NJ)

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130      C
131      C  S/R: D2XYS - CALCULATES 2ND DERIVATIVE IN X- OR Y-
132      C                OIRECTION DEPENOING UPON WHETHER ALONGX OR
133      C                ALONGY IS TRUE.
134      C
135      C  AUTHOR:  A. STANIFORTH - OCTOBER 1977
136      C
137      C  REVISION 001:  A. STANIFORTH - C. THIBEAULT JAN 79 DOCUMENTATION
138      C  REVISION 002:  T. GOOS - 1979-80 AOAATED FOR AMDAHL AT U OF A
139      C
140      C  CALL:  CALL D2XYS(R,U,CON,ALONGX,ALONGY,NI,NJ)
141      C
142      C  ARGUMENTS:
143      C    OUT    - R      - RESULT
144      C    IN     - U      - INPUT FIELD TO BE TWICE DIFFERENCIATED
145      C              - CON   - CONSTANT
146      C              - ALONGX - IF TRUE DIFFERENTIATE IN X-OIRECTION
147      C              - ALONGY - IF TRUE OIFFERENTIATE IN Y-OIRECTION
148      C              - NI     - X-DIMENSION
149      C              - NJ     - Y-OIMENSION
150      C
151      C
152      C    IMPLICIT REAL*8 (A-H,O-Z)
153      C    DIMENSION R(NI,NJ),U(NI,NJ)
154      C    COMMON / WKS1D4 / WK4(1)
155      C    COMMON / WKS1D5 / A(1)
156      C    COMMON / WKS1D6 / B(1)
157      C    COMMON / WKS1D7 / C(1)
158      C    COMMON / HX MESH / HX(1)
159      C    COMMON / HY MESH / HY(1)
160      C    LOGICAL ALONGX,ALONGY
161      C
162      C
163      C
164      C    NIM=NI-1
165      C    NJM=NJ-1
166      C    IF(ALONGX) GO TO 100
167      C    IF(ALONGY) GO TO 200
168      C    RETURN
169      C
170      C  X-DIRECTION
171      C
172      C  100 CONTINUE
173      C    CALL SETD2(A,B,C,CON,HX,NI)
174      C    DO 130 J=1,NJ
175      C      WK4(NI)=B(NI)*U(NI,J)
176      C      DO 120 I=1,NIM
177      C        I1=I+1
178      C      120 WK4(I)=C(I)*U(I1,J)+B(I)*U(I,J)
179      C        DO 125 I=1,NIM
180      C          I1=I+1
181      C        125 R(I1,J)=A(I1)*U(I,J)+WK4(I1)
182      C        R(1,J)=WK4(1)
183      C      130 CONTINUE
184      C      RETURN
185      C
186      C  Y-DIRECTION
187      C
188      C  200 CONTINUE
189      C    CALL SETO2(A,B,C,CON,HY,NJ)
190      C    DO 230 I=1,NI
191      C      WK4(NJ)=B(NJ)*U(I,NJ)
192      C      DO 220 J=1,NJM
193      C        J1=J+1
194      C      220 WK4(J)=C(J)*U(I,J1)+B(J)*U(I,J)
195      C      DO 225 J=1,NJM

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196         J1=J+1
197     225 R(I,J1)=A(J1)*U(I,J)+WK4(J1)
198         R(I,1)=WK4(1)
199     230 CONTINUE
200         RETURN
201         END
202     SUBROUTINE GDADGD(R,CONG,G,CONH,H,NI,NJ)
203     C
204     C S/R:  GDADGD - R(I,J)=CONG*G(I,J)+CONH*H(I,J)
205     C
206     C AUTHOR: A. STANIFORTH DEC 78
207     C
208     C REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
209     C
210     C ARGUMENTS:
211     C OUT      - R      - RESULTING GRID
212     C IN       - CONG   - MULTIPLICATIVE CONSTANT FOR FIRST INPUT GRID
213     C          - G      - FIRST INPUT GRID
214     C          - CONH   - MULTIPLICATIVE CONSTANT FOR SECOND INPUT GRID
215     C          - H      - SECOND INPUT GRID
216     C          - NI     - DIMENSION OF X-DIRN
217     C          - NJ     - DIMENSION OF Y-DIRN
218     C
219     C - R AND G MAY SHARE THE SAME SPACE IF G NOT REQUIRED SUBSEQUENTLY
220     C - R AND H MAY SHARE THE SAME SPACE IF H NOT REQUIRED SUBSEQUENTLY
221     C
222     C
223         IMPLICIT REAL*8 (A-H,O-Z)
224         DIMENSION R(NI,NJ),G(NI,NJ),H(NI,NJ)
225         DO 20 J=1,NJ
226         DO 20 I=1,NI
227             R(I,J)=CONG*G(I,J)+CONH*H(I,J)
228     20 CONTINUE
229         RETURN
230         END
END OF FILE

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1  C
2  C
3  C FILE: INTERP
4  C OBJECT FILE: INTERP.O
5  C SUBROUTINES:
6  C   - INTRPT
7  C   - DINT
8  C   - ROSSR3
9  C   - SPD
10 C   - FD1
11 C   - FDM
12 C
13 C
14 C
15 C   SUBROUTINE INTRPT(FI,IFI,JFI,F,IF,JF,XI,YI,X,Y,FX,FY,FXY,
16 C   1 HX,HY,KL,P,S,C,A,O,WORK,KOER)
17 C
18 C ORIGINALLY WRITTEN AT ORPN, OORVAL P.Q.
19 C
20 C REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
21 C
22 C INTRPT DOES 2-DIMENSIONAL SPLINE INTERPOLATION. THE INPUT F(IF,JF) IS
23 C DEFINED ON THE PTS ((X(I),Y(J),I=1,IF),J=1,JF). THE OUTPUT
24 C FI(4,IFI,JFI) IS CALCULATED AT PTS ((XI(I),YI(J),I=1,IFI),J=1,JFI)
25 C USING BI-CUBIC SPLINES IN TERMS OF THE PARTIAL DERIVATIVES OF F,FX,FY,
26 C FXY. THESE P.D.S ARE CALCULATED BY SPD. (IF KL.EQ.1, FXY=DX(DY(F)),
27 C WHILE IF KL.NE.1, FXY=DY(DX(F))).
28 C FD1 AND FDM RETURN END PT DERIVATIVES USED BY SPD.
29 C THE VECTORS XI,YI,X,Y,(HX(I)=X(I+1)-X(I),I=1,IF-1),
30 C (HY(J)=Y(J+1)-Y(J),J=1,JF-1), ARE ADDITIONAL INPUT.
31 C THE ACTUAL INTERPOLATION IS DONE BY THE ROUTINE DINT, AFTER THE
32 C EVALUATION OF FX,FY,FXY.
33 C IF KOER=FALSE, THEN ONLY INTERPOLATED VALUES OF THE FUNCTION ARE
34 C RETURNED.
35 C IF KDER=TRUE, THE VALUES OF THE X,Y AND XY DERIVATIVES AT THE
36 C INTERPOLATED PTS WILL ALSO BE RETURNED.
37 C
38 C   IMPLICIT REAL*8 (A-H,O-Z)
39 C   DIMENSION FI(4,IFI,JFI)
40 C   DIMENSION F (IF,JF)
41 C   DIMENSION FX(IF,JF)
42 C   DIMENSION FY(IF,JF)
43 C   DIMENSION FXY(IF,JF)
44 C
45 C THE PROPER DIMENSIONS HERE ARE IFI,JFI,IF,JF,IF,JF
46 C
47 C   DIMENSION XI(IFI)
48 C   DIMENSION YI(JFI)
49 C   DIMENSION X(IF)
50 C   DIMENSION Y(JF)
51 C   DIMENSION HX(IF)
52 C   DIMENSION HY(JF)
53 C
54 C WORKING STORAGE FOR SPD,ROSSR3 (LENGTH MAXO(IF,JF) IS SUFFICIENT)
55 C
56 C   DIMENSION P(IF)
57 C   DIMENSION S(IF)
58 C   DIMENSION A(IF)
59 C   DIMENSION C(IF)
60 C   DIMENSION D(IF)
61 C
62 C WORKING STORAGE FOR DINT
63 C

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64      DIMENSION WORK(JFI,8)
65      LOGICAL KDER
66      C
67      C CHECK TO ENSURE EXTRAPOLATION IS NOT BEING ATTEMPTED A MACHINE
68      C DEPENDENT INCREMENT IS ADDED IN TEST TO ACCOUNT FOR MACHINE ROUND-OFF.
69      C
70      X1=XI(1)+0ABS(XI(1)*1.0-12)-X(1)
71      IF(X1.LT.0.) WRITE(6,900) XI(1),X(1)
72      IF(X1.LT.0.) STDP
73      X1=YI(1)+0ABS(YI(1)*1.0-12)-Y(1)
74      IF(X1.LT.0.) WRITE(6,901) YI(1),Y(1)
75      IF(X1.LT.0.) STOP
76      X1= X(1)+0ABS( X(1)*1.0-12)-XI(1)
77      X1=X(IF)+DABS(X(IF)*1.0-12)-XI(IFI)
78      IF(X1.LT.0.) WRITE(6,902) IF,X(IF),IFI,XI(IFI)
79      IF(X1.LT.0.) STDP
80      X1=Y(JF)+DABS(Y(JF)*1.0-12)-YI(JFI)
81      IF(X1.LT.0.) WRITE(6,903) JF,Y(JF),JFI,YI(JFI)
82      IF(X1.LT.0.) STOP
83      900 FDRMAT(1H1,' EXTRAPDLATION ATTEMPTED, DETECTED BY INTRPT. XI(1)=' ,
84      1E16.8,'LT X(1)=' ,E16.8)
85      901 FORMAT(1H1,' EXTRAPDLATIDN ATTEMPTED, DETECTED BY INTRPT. YI(1)=' ,
86      1E16.8,'LT Y(1)=' ,E16.8)
87      902 FORMAT(1H1,' EXTRAPOLATION ATTEMPTED, DETECTED BY INTRPT. X(' ,I5,
88      1')=' ,E16.8,'LT XI(' ,I5,' )=' ,E16.8)
89      903 FORMAT(1H1,' EXTRAPOLATION ATTEMPTED, DETECTED BY INTRPT. Y(' ,I5,
90      1')=' ,E16.8,'LT YI(' ,I5,' )=' ,E16.8)
91      CMU1=0.0DO
92      CLMDAM = 0.0DO
93      IF(KL.NE.1) GO TO 150
94      10 DO 100 I=1,IF
95      DD 50 J=1,JF
96      S(J)=F(I,J)
97      50 CONTINUE
98      C1=FD1(S,HY,JF)
99      CM = FDM(S,HY,JF)
100     CALL SPD (P,S,JF,HY,CMU1,C1,CLMOAM,CM,A,C,D)
101     DD 60 J=1,JF
102     FY(I,J)=P(J)
103     60 CONTINUE
104     IF(KL.EQ.1) GO TO 100
105     DD 70 J=1,JF
106     S(J)=FX(I,J)
107     70 CONTINUE
108     C1=FD1(S,HY,JF)
109     CM=FDM(S,HY,JF)
110     CALL SPD(P,S,JF,HY,CMU1,C1,CLMDAM,CM,A,C,D)
111     DD 75 J=1,JF
112     FXY(I,J)=P(J)
113     75 CONTINUE
114     100 CDNTINUE
115     IF(KL.NE.1) GO TO 600
116     150 DD 500 J=1,JF
117     DO 200 I=1,IF
118     S(I)=F(I,J)
119     200 CONTINUE
120     C1=FD1(S,HX,IF)
121     CM=FDM(S,HX,IF)
122     CALL SPD(P,S,IF,HX,CMU1,C1,CLMOAM,CM,A,C,D)
123     DD 210 I=1,IF
124     FX(I,J)=P(I)
125     210 CONTINUE
126     IF(KL.NE.1) GO TO 500
127     DD 260 I=1,IF
128     S(I)=FY(I,J)
129     260 CDNTINUE

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130      C1=FD1(S,HX,IF)
131      CM=FDM(S,HX,IF)
132      CALL SPD(P,S,IF,HX,CMU1,C1,CLMDAM,CM,A,C,D)
133      DO 280 I=1,IF
134      FXY(I,J)=P(I)
135 280 CONTINUE
136 500 CONTINUE
137      IF(KL.NE.1) GD TD 10
138 600 CALL DINT(FI,IFI,JFI,F,IF,JF,FX,FY,FXI,YI,X,Y,HX,HY,
139      1 WORK(1,1),WORK(1,2),WDRK(1,3),WDRK(1,4),WDRK(1,5),WORK(1,6),
140      2 WORK(1,7),WORK(1,8),KDER)
141      RETURN
142      END
143      SUBROUTINE DINT(FI,IFI,JFI,F,IF,JF,FX,FY,FXI,YI,X,Y,HX,HY,
144      1 ZA,ZB,ZC,ZD,ZAY,ZBY,ZCY,ZDY,KDER)
145      IMPLICIT REAL*8 (A-H,O-Z)
146      DIMENSION FI(4,IFI,JFI)
147      DIMENSION F(IF,JF)
148      DIMENSION FX(IF,JF)
149      DIMENSION FY(IF,JF)
150      DIMENSION FXY(IF,JF)
151      DIMENSION XI(IFI)
152      DIMENSION YI(JFI)
153      DIMENSION X(IF)
154      DIMENSION Y(JF)
155      DIMENSION HX(IF)
156      DIMENSION HY(JF)
157      DIMENSION ZA(JFI),ZB(JFI),ZC(JFI),ZD(JFI)
158      DIMENSION ZAY(JFI),ZBY(JFI),ZCY(JFI),ZDY(JFI)
159      LOGICAL KDER
160      C
161      C DRIGINALLY WRITTEN AT DRPN, DORVAL P.Q.
162      C REVISIDN 001: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
163      C
164      C THE DISCRETE FN F(I,J) IS ASSUMED KNOWN AT THE PTS WHOSE CODRDS ARE
165      C (X(I),Y(J)), WHERE I=1,...IF ,J=1,...JF.
166      C IT IS ASSUMED THAT X(I).LT.X(I+1) AND Y(J).LT.Y(J+1).
167      C HX(I) MUST BE PREVIDUSLY DEFINED AS HX(I)=X(I+1)-X(I) FDR I=1,...(IF-1).
168      C SIM. FDR HY(J).
169      C IF KDER=TRUE THEN
170      C THIS SUBROUTINE RETURNS IN FI(K,I,J) THE INTERPDLATED VALUES OF THE
171      C FUNCTION AND ITS X,Y,XY DERIVATIVES AT PTS (XI(I),YI(J)),
172      C WHERE I=1,...IFI ,J=1,...JFI.
173      C IT IS ASSUMED THAT XI(I).LT.XI(I+1) AND YI(J).LT.YI(J+1).
174      C HERE K=1 REFERS TO THE FUNCTION,
175      C K=2 REFERS TO THE X-DER OF THE FUNCTION,
176      C K=3 REFERS TO THE Y-DER OF THE FUNCTION,
177      C K=4 REFERS TO THE XY-DER OF THE FUNCTION.
178      C NOTE THAT X(1).LE.XI(1).LE.XI(IFI).LE.XI(IF).
179      C SIM. FOR Y AND YI.
180      C IF KDER=FALSE THEN ONLY THE INTERPOLATED VALUES OF THE FUNCTION ARE
181      C RETURNED. (I.E. CALCULATIONS ARE PERFORMED FOR K=1 ONLY).
182      C THE APPROXIMATION USED IS A BICUBIC SPLINE OF INTERPOLATION IN
183      C TERMS OF THE PARTIAL DERIVATIVES FX(I,J),FY(I,J) AND FXY(I,J)
184      C WHERE I=1,...IF ,J=1,...JF.
185      C THESE DERIVATIVES MUST BE CALCULATED OUTSIDE OF THE RDUTINE.
186      C E.G. BY USING SUBROUTINE SPD IN A SUITABLE MANNER - SUBROUTINE INTD
187      C DDES THIS AND CALLS THIS SUBRDUTINE. (INTD IS CALLABLE FROM INTRPT).
188      C WORKING STORAGE ARRAYS ZA,ZB,ZC,ZD HAVE DIMENSION JFI.
189      C
190      C
191      LL=2
192      DD 15 J=1,JFI
193      DO 5 L=LL,JF
194      IF(YI(J).LE.Y(L)) GO TD 8
195      5 CONTINUE

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196      L=JF
197      8 L1=L-1
198      LL=L
199      WN=YI(J)-Y(L1)
200      WD=1.ODO/HY(L1)
201      WE=WN*WD
202      WE1=1.ODO-WE
203      WE2=WE1*WE1
204      WW=2.ODO*WE
205      ZA(J)=WE2*WN
206      ZB(J)=WE1*WN*WE
207      ZC(J)=WE2*(1.ODO+WW)
208      ZD(J)=WE*WE*(3.ODO-WW)
209      ZAY(J)=-WW*WE1+WE2
210      ZBY(J)=-WE*WE+WW*WE1
211      ZCY(J)=-6.ODO*WD*WE1*WE
212      ZDY(J)=-ZCY(J)
213      15 CONTINUE
214      KK=2
215      DO 500 I=1,IFI
216      DO 40 K=KK,IF
217      IF(XI(I).LE.X(K)) GO TO 45
218      40 CONTINUE
219      K=IF
220      45 K1=K-1
221      KK=K
222      WM=XI(I)-X(K1)
223      WDD=1.ODO/HX(K1)
224      WZ=WM*WDD
225      WZ1=1.ODO-WZ
226      WZ2=WZ1*WZ1
227      ZZ=2.ODO*WZ
228      ZE=WZ2*WM
229      ZF=WZ1*WM*WZ
230      ZG=WZ2*(1.ODO+ZZ)
231      ZL=WZ*WZ*(3.ODO-ZZ)
232      IF(.NOT.KDER) GO TO 46
233      ZEX=-ZZ*WZ1+WZ2
234      ZFX=-WZ*WZ+ZZ*WZ1
235      ZGX=-6.ODO*WDD*WZ1*WZ
236      ZLX=-ZGX
237      46 LL=2
238      DO 400 J=1,JFI
239      DO 420 L=LL,JF
240      IF(YI(J).LE.Y(L)) GO TO 450
241      420 CONTINUE
242      L=JF
243      450 L1=L-1
244      LL=L
245      Z1=ZE*FX(K1,L1)-ZF*FX(K,L1)+ZG*FY(K1,L1)+ZL*FY(K,L1)
246      Z2=ZE*FX(K1,L)-ZF*FX(K,L)+ZG*FY(K1,L)+ZL*FY(K,L)
247      Z3=ZE*FX(K1,L1)-ZF*FX(K,L1)+ZG*F(K1,L1)+ZL*F(K,L1)
248      Z4=ZE*FX(K1,L)-ZF*FX(K,L)+ZG*F(K1,L)+ZL*F(K,L)
249      FI(1,I,J)=ZA(J)*Z1-ZB(J)*Z2+ZC(J)*Z3+ZD(J)*Z4
250      IF(.NOT.KDER) GO TO 400
251      Z1X=ZEX*FX(K1,L1)-ZFX*FX(K,L1)+ZGX*FY(K1,L1)+ZLX*FY(K,L1)
252      Z2X=ZEX*FX(K1,L)-ZFX*FX(K,L)+ZGX*FY(K1,L)+ZLX*FY(K,L)
253      Z3X=ZEX*FX(K1,L1)-ZFX*FX(K,L1)+ZGX*F(K1,L1)+ZLX*F(K,L1)
254      Z4X=ZEX*FX(K1,L)-ZFX*FX(K,L)+ZGX*F(K1,L)+ZLX*F(K,L)
255      FI(2,I,J)=ZA(J)*Z1X-ZB(J)*Z2X+ZC(J)*Z3X+ZD(J)*Z4X
256      FI(3,I,J)=ZAY(J)*Z1-ZBY(J)*Z2+ZCY(J)*Z3+ZDY(J)*Z4
257      FI(4,I,J)=ZAY(J)*Z1X-ZBY(J)*Z2X+ZCY(J)*Z3X+ZDY(J)*Z4X
258      400 CONTINUE
259      500 CONTINUE
260      RETURN
261      END

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262      FUNCTION FDM(F,H,M)
263      C
264      C  FUNCTION: FOM - THIS FUNCTION RETURNS IN FOM THE DERIVATIVE
265      C                OF F AT THE PT X(M) IN TERMS OF
266      C                F(M-3),F(M-2),F(M-1),F(M),H(M-3),H(M-2),H(M-1)
267      C                WHERE H(I)=X(I+1)-X(I) FOR I=M-3,M-2,M-1.
268      C
269      C  AUTHOR:  A. STANIFORTH - 1973
270      C
271      C  REVISION 001: A. STANIFORTH - C. THIBEAULT JAN 79 DOCUMENTATION
272      C  REVISION 002: T. GOOS - 1979-80 ADAPTED FOR AMOHL AT U OF A
273      C
274      C  FUNCTION :  CALL FDM(F,H,M)
275      C
276      C  ARGUMENTS:
277      C      IN      - F - FIELD OF VALUES
278      C              - H - MESH-SPACING
279      C              - M - NO. OF POINTS
280      C
281      C  NOTES: - THE APPROXIMATION USED IS THE DIFFERENTIATED FORM OF
282      C          LAGRANGES CUBIC INTERPOLATION FORMULA FOR NON-UNIFORM
283      C          GRIDS.
284      C
285      C
286      C      IMPLICIT REAL*8 (A-H,O-Z)
287      C      DIMENSION F(M),H(M)
288      C      H1=H(M-1)
289      C      H2=H(M-2)
290      C      H3=H(M-3)
291      C      X1=H1+H2
292      C      X2=H2+H3
293      C      X3=X1+H3
294      C      FDM=-X1*H1*F(M-3)/(H3*X2*X3)+X3*H1*F(M-2)/(H3*H2*X1)
295      C      FDM=FDM-X3*X1*F(M-1)/(X2*H2*H1)+(1.000/X3+1.000/X1+1.000/H1)*F(M)
296      C      RETURN
297      C      END
298      C      FUNCTION FD1(F,H,M)
299      C      IMPLICIT REAL*8 (A-H,O-Z)
300      C      DIMENSION F(M),H(M)
301      C
302      C  ORIGINALLY WRITTEN AT DRPN, DORVAL P.Q.
303      C  REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMOHL AT U OF A
304      C
305      C
306      C  LET F(1),F(2),..... BE THE VALUES OF A FUNCTION F DEFINED AT
307      C  SUCCESSIVE PTS X(1),X(2),..... OF A NON UNIFORM GRID.
308      C  THIS FUNCTION ROUTINE RETURNS IN F01 THE DERIVATIVE OF F AT THE PT
309      C  X(1) IN TERMS OF F(1),F(2),F(3),F(4),H(1),H(2),H(3).
310      C  THE APPROXIMATION USED IS THE DIFFERENTIATED FORM OF LAGRANGES CUBIC
311      C  INTERPOLATION FORMULA FOR NON UNIFORM GRIDS.
312      C
313      C
314      C      H1=H(1)
315      C      H2=H(2)
316      C      H3=H(3)
317      C      X1=H1+H2
318      C      X2=H2+H3
319      C      X3=X1+H3
320      C      FD1=-(1.000/H1+1.000/X1+1.000/X3)*F(1)+X1*X3*F(2)/(H1*H2*X2)
321      C      FD1=FD1-H1*X3*F(3)/(X1*H2*H3)+H1*X1*F(4)/(X3*X2*H3)
322      C      RETURN
323      C      END
324      C      SUBROUTINE ROSSR3(P,A,DELTA,C,O,M)
325      C
326      C  ORIGINALLY WRITTEN AT ORPN, DORVAL P.Q.
327      C  REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMOHL AT U OF A

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328 C
329 C VERSION 3
330 C ANY TRI-DIAGONAL MATRIX SAY Q IN EQUATION QP=D MAY BE
331 C NORMALIZED WITH 1'S ON DIAGONAL AND C(I), I=1,M-1 FOR UPPER DIAGONAL
332 C AND A(I), I=2,M FOR LOWER DIAGONAL.
333 C
334 C DELTA IS A WORKING STORAGE ARRAY OF DIMENSION M
335 C IF THE VECTOR D IS NOT REQUIRED SUBSEQUENTLY THEN THE CALL STATEMENT
336 C CALL ROSSR3(P,A,D,C,D,M)
337 C WILL USE THE ARRAY D AS WORKING STORAGE AND REDUCE THE OVERALL STORAGE
338 C REQUIRED.
339 C IF THE ARRAY C IS NOT REQUIRED SUBSEQUENTLY THEN THE CALL STATEMENT
340 C CALL ROSSR3(C,A,DELTA,C,D,M)
341 C WILL REDUCE CORE STORAGE REQUIREMENTS.
342 C IF BOTH C AND D ARE NOT REQUIRED SUBSEQUENTLY THEN THE CALL STATEMENT
343 C CALL ROSSR3(C,A,D,C,D,M)
344 C WILL FURTHER REDUCE THE CORE REQUIREMENTS.
345 C IMPLICIT REAL*8 (A-H,O-Z)
346 C DIMENSION P(M),A(M),DELTA(M),C(M),D(M)
347 C C(M)=0.0D0
348 C P(1)=-C(1)
349 C DELTA(1)=D(1)
350 C DO 1 I=2,M
351 C I1=I-1
352 C AI=A(I)
353 C X=1.0D0/(1.0D0+AI*P(I1))
354 C P(I)=-C(I)*X
355 C 1 DELTA(I)=(D(I)-AI*DELTA(I1))*X
356 C P(M)=DELTA(M)
357 C DO 2 I=2,M
358 C II=M-I+1
359 C I11=II+1
360 C 2 P(II)=P(II)*P(I11)+DELTA(II)
361 C RETURN
362 C END
363 C SUBROUTINE SPD(P,S,M,H,CMU1,C1,CLMDAM,CM,A,C,D)
364 C
365 C
366 C THIS SUBROUTINE RETURNS IN P THE DERIVATIVES OF THE FUNCTION S DEFINED
367 C AT M CONSECUTIVE DISCRETE DATA POINTS X(1),X(2),....X(M).
368 C THESE DATA POINTS DEFINE A NON-UNIFORM GRID H(I)=X(I+1)-X(I) WHERE I
369 C RUNS FROM 1 TO (M-1)
370 C IT USES SUBROUTINE ROSSR3.
371 C THE APPROXIMATION USED IS A CUBIC SPLINE FIT TO ALL THE DATA POINTS.
372 C IT REQUIRES BOUNDARY CONDITIONS TO BE SPECIFIED OF THE FORM
373 C P(1)+CMU1*P(2)=C1
374 C CLMDAM*P(M-1)+P(M)=CM
375 C FOR A NATURAL SPLINE (THIS IS LEAST RESTRICTIVE CONDITION AND IMPLIES
376 C ZERO CURVATURE AT THE END POINTS), SDD(1)=SDD(M)=0, AND WE HAVE
377 C CMU1=CLMDAM=0.5, C1=1.5*(S(2)-S(1))/H(1), CM=1.5*(S(M)-S(M-1))/H(M-1)
378 C
379 C FOR SPECIFIED SLOPES (I.E. SD(1) AND SD(M) GIVEN) WE HAVE
380 C CMU1=CLMDAM=0. C1=SD(1) CM=SD(M)
381 C
382 C FOR SPECIFIED SECOND DERIVATIVES (I.E. SDD(1) AND SDD(M) GIVEN) WE HAVE
383 C CMU1=CLMDAM=0.5, C1=1.5*(S(2)-S(1))/H(1)-H(1)*SDD(1)*0.25
384 C CM=1.5*(S(M)-S(M-1))/H(M-1)+H(M-1)*SDD(M)*0.25
385 C
386 C NOTE THAT IF THE GRID SIZES H(I) ARE FIXED THEN THE CALCULATION OF A
387 C AND C MAY BE MADE OUTSIDE THE SUBROUTINE AND PASSED VIA A COMMON BLOCK.
388 C
389 C IMPLICIT REAL*8 (A-H,O-Z)
390 C DIMENSION P(M),S(M),A(M),C(M),D(M),H(M)
391 C M1=M-1
392 C DO 1 I=2,M1
393 C I1=I-1

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394      HI=H(I)
395      HI1=H(I1)
396      X1=0.5DO/(HI+HI1)
397      A(I)=HI*X1
398      1 C(I)=HI1*X1
399      DO 2 I=2,M1
400      I1=I-1
401      I2=I+1
402      HI=H(I)
403      HI1=H(I1)
404      SI=S(I)
405      X1=H(I)/H(I-1)
406      2 D(I)=1.5DO*((SI-S(I1))*X1+(S(I2)-SI)/X1)/(HI+HI1)
407      C(1)=CMU1
408      A(M)=CLMDAM
409      D(1)=C1
410      D(M)=CM
411      CALL ROSSR3(P,A,D,C,D,M)
412      RETURN
413      END
END OF FILE
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1      C      FILE:  PROD
2      C      OBJECT FILE:  PROD.O
3      C      SUBROUTINES:
4      C          - NLINX
5      C          - NLLLOOP
6      C
7      C
8      C
9      C      SUBROUTINE NLINX(R,U,V,H6,NI)
10     C
11     C      ADAPTED FROM DRPN DDRVAL P.Q.
12     C      REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
13     C
14     C      IMPLICIT REAL*8 (A-H,D-Z)
15     C      DIMENSION R(NI),U(NI),V(NI)
16     C
17     C      H6(I)=H(I)/6. MUST BE PRESET BEFORE ROUTINE
18     C
19     C      DIMENSION H6(NI)
20     C      NIM=NI-1
21     C      WK50=U(1)*V(1)
22     C      R(1)=0.000
23     C      DD 300 I=1,NIM
24     C      I1=I+1
25     C      H6I=H6(I)
26     C      WK4=(U(I)+U(I1))*(V(I)+V(I1))* .5000
27     C      R(I)=R(I)+(WK50+WK4)*H6I
28     C      WK50=U(I1)*V(I1)
29     C      300 R(I1)=(WK50+WK4)*H6I
30     C      RETURN
31     C      END
32     C      SUBROUTINE NLLDOP(R,U,V,NI,NJ)
33     C
34     C      S/R : NLLDOP - COMPUTE THE HORIZONTAL PRDDUCT R=U*V
35     C
36     C      AUTHDR : A. STANIFDRTH - REVISED DEC 78 TO REMOVE BANK CONFLICTS
37     C      REVISION 001: T. GOOS 1979-80 ADAPTED FOR AMDAHL AT U OF A
38     C
39     C      ARGUMENTS:
40     C      OUT  - R  - RESULTING PRODUCT
41     C      IN   - U  - INPUT FIELD
42     C          - V  - INPUT FIELD
43     C          - NI - X DIMENSION
44     C          - NJ - Y DIMENSION
45     C
46     C
47     C      IMPLICIT REAL*8 (A-H,O-Z)
48     C      DIMENSION R(NI,NJ),U(NI,NJ),V(NI,NJ)
49     C      COMMON / HXMESH / HX(1)
50     C      COMMON / HYMESH / HY(1)
51     C      CDMMON / WKS1D1 / HX6(1)
52     C      CDMMON / WKS1D2 / UO(1)
53     C      CDMMON / WKS1D3 / VO(1)
54     C      COMMON / WKS1D6 / RO(1)
55     C      COMMON / WKS1D7 / WK7(1)
56     C      CDMMON / WKS1D8 / WK8(1)
57     C      CDMMON / WKS1D9 / WK9(1)
58     C      COMMON / WKS110 / HDLD(1)
59     C      TWELFT=1.000/12.000
60     C      SIXTH=1.000/6.000
61     C      NIM=NI-1
62     C      NJM=NJ-1
63     C

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64      C  COMPUTE H6(I)=H(I)/6.
65      C  AND SET RO TO O.
66      C
67          DO 50 I=1,NIM
68          UO(I)=U(I,1)
69          VO(I)=V(I,1)
70          RO(I)=O.ODO
71      50 HX6(I)=HX(I)/6.ODO
72          RO(NI)=O.ODO
73          UO(NI)=U(NI,1)
74          VO(NI)= V(NI,1)
75          J=1
76      C
77      C  COMPUTE ROW PRODUCT U*V ON FIRST ROW
78      C
79          CALL NLINX(WK7,UO,VO,HX6,NI)
80      C
81      C  LOOP OVER ROWS
82      C
83          DO 100 J=1,NJM
84          J1=J+1
85          HYJ1=HY(J)*SIXTH
86          HYJ2=HY(J)*TWELFT
87          DO 75 I=1,NI
88      C
89      C  ADD FIRST CONTRIBUTION TO J TH ROW
90      C
91          HOLD(I)=HYJ1*WK7(I)+RO(I)
92      C
93      C  COMPUTE 2.*U AND 2.*V ON COLLOCATION ROW
94      C
95          WK8(I)=UO(I)+U(I,J1)
96      75 WK9(I)=VO(I)+V(I,J1)
97      C
98      C  COMPUTE ROW PRODUCT HY(J)*U*V/3. ON COLLOCATION ROW
99      C
100         CALL NLINX(WK7,WK8,WK9,HX6,NI)
101         DO 90 I=1,NI
102         WK8(I)=HYJ2*WK7(I)
103      C
104      C  ADD OTHER CONTRIBUTION TO J TH ROW
105      C
106         RO(I)=HOLD(I)+WK8(I)
107      C
108      C  MOVE RESULT RO INTO R(1,J)
109      C
110         R(I,J)=RO(I)
111      C
112      C  ADD FIRST CONTRIBUTION TO (J+1)ST ROW
113      C
114      90 HOLD(I)=WK8(I)
115      C
116      C  COMPUTE ROW PRODUCT U*V ON (J+1)ST ROW
117      C
118         CALL NLINX(WK7,U(1,J1),V(1,J1),HX6,NI)
119      C
120      C  ADD OTHER CONTRIBUTION TO (J+1)ST LEVEL
121      C
122         DO 120 I=1,NI
123         RO(I)=HYJ1*WK7(I)+HOLD(I)
124      C
125      C  MOVE U(I,J+1),V(I,J+1) INTO UO,VO
126      C
127         UO(I)=U(I,J1)
128      120 VO(I)=V(I,J1)
129      100 CONTINUE

```



```
130      C
131      C  MOVE RESULT RO INTO LAST ROW OF R
132      C
133          DO 150 I=1,NI
134      150 R(I,NJ)=RO(I)
135          RETURN
136          END
END OF FILE
```



```

1 C
2 C
3 C FILE: PROJ
4 C OBJECT FILE: PROJ.O
5 C SUBROUTINES:
6 C - PROJN
7 C - SETABC
8 C - SETTRI
9 C - SOLTRI
10 C - PSOLVE
11 C
12 C
13 SUBROUTINE PROJN(R,RHS,FOURTH,NI,NJ,ALONGX,ALONGY)
14 C
15 C S/R: PROJN - THE FINITE-ELEMENT PROJECTION OPERATOR (ALONG ANY
16 C OR ALL DIRECTIONS, AS DECIDED BY THE LOGICAL VARIABLES
17 C ALONGX,ALONGY)
18 C
19 C R = PROJECTION (RHS)
20 C
21 C AUTHOR: A. STANIFORTH - SPRING 1977
22 C
23 C REVISION 001: A. STANIFORTH - JAN 1979 VECTORIZED
24 C REVISION 002: A. STANIFORTH - C. THIBEAULT FEB 79 DOCUMENTATION
25 C REVISION 003: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
26 C
27 C CALL: CALL PROJN(R,RHS,FOURTH,NI,NJ,ALONGX,ALONGY)
28 C
29 C ARGUMENTS:
30 C OUT - R - RESULT OF PROJECTION
31 C IN - RHS - RIGHT-HAND-SIDE TO BE PROJECTED
32 C - FOURTH - LOGICAL SWITCH - TRUE = FOURTH ORDER PROJECTION
33 C IN HORIZONTAL
34 C - FALSE = SECOND ORDER PROJECTION
35 C IN HORIZONTAL
36 C - NI - X-DIMENSION
37 C - NJ - Y-DIMENSION
38 C - ALONGX - IF .TRUE. PROJECT IN X DIRECTION
39 C - ALONGY - IF .TRUE. PROJECT IN Y DIRECTION
40 C
41 C
42 C NOTES: - AT PRESENT THE CODE HANDLES SECOND AND FOURTH ORDER IN
43 C BOTH DIMENSIONS.
44 C - IF ALONGX AND ALONGY ARE ALL .FALSE., A COPY OF RHS IS
45 C RETURNED IN R.
46 C
47 C
48 C
49 C IMPLICIT REAL*8 (A-H,O-Z)
50 C DIMENSION R(NI,NJ), RHS(NI,NJ)
51 C COMMON / HXMESH / HX(1)
52 C COMMON / HYMesh / HY(1)
53 C COMMON / WKS1D5 / A(1)
54 C COMMON / WKS1D6 / B(1)
55 C COMMON / WKS1D7 / C(1)
56 C LOGICAL ALONGX,ALONGY
57 C LOGICAL FOURTH
58 C
59 C
60 C SET UP WEIGHTS FOR PROJECTION OPERATOR
61 C WT = 2. CORRESPONDS TO USUAL PROJECTION, WHILST
62 C WT = 5. CORRESPONDS TO PROJECTION USED IN THE FOURTH-ORDER SOLUTION
63 C OF ELLIPTIC BOUNDARY-VALUE PROBLEMS.

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```

64      C
65      WTX=2.000
66      WTY=2.000
67      IF(FOURTH) WTX=5.000
68      IF(FOURTH) WTY=5.000
69      C
70      C CODE FOR PROJECTION ALONG X DIRECTION
71      C
72      IF(.NOT.ALONGX) GO TO 200
73      CALL SETABC(A,B,C,HX,WTX,NI)
74      NIM=NI-1
75      DO 130 J=1,NJ
76      DO 150 I=1,NIM
77      I1=I+1
78      I2=I+2
79      R(I1,J)=A(I1)*RHS(I,J)+B(I1)*RHS(I1,J)
80      IF(I.LT.NIM) R(I1,J)=R(I1,J)+C(I1)*RHS(I2,J)
81      150 CONTINUE
82      R(1,J)=C(1)*RHS(2,J)+B(1)*RHS(1,J)
83      130 CONTINUE
84      DO 135 J=1,NJ
85      DO 135 I=1,NI
86      RHS(I,J)=R(I,J)
87      135 CONTINUE
88      C
89      200 CONTINUE
90      C
91      C CODE FOR PROJECTION ALONG Y DIRECTION
92      C
93      IF(.NOT.ALONGY) RETURN
94      CALL SETABC(A,B,C,HY,WTY,NJ)
95      NJM=NJ-1
96      DO 230 I=1,NI
97      DO 250 J=1,NJM
98      J1=J+1
99      J2=J+2
100     R(I,J1)=A(J1)*RHS(I,J)+B(J1)*RHS(I,J1)
101     IF(J.LT.NJM) R(I,J1)=R(I,J1)+C(J1)*RHS(I,J2)
102     250 CONTINUE
103     R(I,1)=C(1)*RHS(I,2)+B(1)*RHS(I,1)
104     230 CONTINUE
105     RETURN
106     END
107     SUBROUTINE SETABC(A,B,C,H,WT,N)
108     C
109     C S/R: SETABC - SETS THE TRI-DIAGONAL ELEMENTS OF THE MATRIX
110     C ASSOCIATED WITH THE PROJECTION OPERATOR.
111     C
112     C AUTHOR: A. STANIFORTH - SPRING 1977
113     C
114     C REVISION 001: A. STANIFORTH - C. THIBEAULT FEB 79 DOCUMENTATION
115     C REVISION 002: T. GOOS - 1979-80 ADAPTED FOR AMOHL AT U OF A
116     C
117     C CALL: CALL SETABC(A,B,C,H,WT,N)
118     C
119     C ARGUMENTS:
120     C OUT - A - LOWER-DIAGONAL ELEMENTS
121     C - B - DIAGONAL ELEMENTS
122     C - C - UPPER-DIAGONAL ELEMENTS
123     C IN - H - MESH-SPACING
124     C - WT - USUALLY WT=2., BUT WT=5. WHEN SOLVING EBV PROBLEMS
125     C TO FOURTH ORDER
126     C - N - NO. OF POINTS
127     C
128     C
129     IMPLICIT REAL*8 (A-H,O-Z)

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130      DIMENSION A(N),B(N),C(N)
131      DIMENSION H(N)
132      NM=N- 1
133      C1=0.5DO/(1.ODO+WT)
134      C2=WT*C1
135      A(1)=0.ODO
136      B(1)=H(1)*C2
137      C(1)=H(1)*C1
138      DO 10 I=2,NM
139      HI1=H(I-1)
140      HI=H(I)
141      A(I)=HI1*C1
142      B(I)=(HI1+HI)*C2
143      C(I)=HI*C1
144 10 CONTINUE
145      A(N)=H(NM)*C1
146      B(N)=H(NM)*C2
147      C(N)=0.ODO
148      RETURN
149      END
150      SUBROUTINE SETTRI(BIGE,BIGC,BIGA,A,B,C,N)
151  C
152  C  S/R: SETTRI - SETTRI DOES THE PREPROCESSING PASS TO SOLVE A
153  C                TRI-DIAGONAL SYSTEM OF LINEAR EQUATIONS.
154  C
155  C  AUTHOR: D. ROBERTSON - MARCH 1977
156  C
157  C  REVISION 001: A. STANIFORTH - C. THIBEAULT FEB 79 DOCUMENTATION
158  C  REVISION 002: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
159  C
160  C  CALL:  CALL  SETTRI(BIGE,BIGC,BIGA,A,B,C,N)
161  C
162  C  ALGORITHM:
163  C      - SOLVES           $M \cdot P = D$ 
164  C
165  C      WHERE M IN THE N BY N TRI-DIAGONAL MATRIX
166  C
167  C      *****
168  C      *B(1),C(1),O,O.....*
169  C      *A(2),B(2),C(2),O,O....*
170  C      *      .      *
171  C      *      .      *
172  C      *      .      *
173  C      * ... O,A(I),B(I),C(I),O...*
174  C      *      .      *
175  C      *      .      *
176  C      *      ...O,O.A(N),B(N)*
177  C      *****
178  C
179  C      - A(1) AND C(N) ARE NOT DEFINED BY THE MATRIX, BUT ARRAYS A,C
180  C      SHOULD BE OF FULL SIZE, N.  FOR CONVENIENCE THIS ROUTINE
181  C      ZEROES C(N).  THE METHOD IS GAUSSIAN ELIMINATION WITHOUT
182  C      PIVOTING, FOLLOWED BY BACK SUBSTITUTION.
183  C
184  C  ARGUMENTS
185  C      OUT      - BIGE - PROCESSED ARRAY TO BE USED IN S/R SOLTRI
186  C                - BIGC - PROCESSED ARRAY TO BE USED IN S/R SOLTRI
187  C                - BIGA - PROCESSED ARRAY TO BE USED IN S/R SOLTRI
188  C      IN       - A      - LOWER-DIAGONAL ELEMENTS OF TRI-DIAGONAL MATRIX
189  C                - B      - DIAGONAL ELEMENTS OF TRI-DIAGONAL MATRIX
190  C      IN       - C      - UPPER-DIAGONAL ELEMENTS OF TRI-DIAGONAL MATRIX
191  C                - N      - NO. OF POINTS
192  C
193  C  NOTES: - SETTRI IS CALLED TO COMPUTE BIGE,BIGC,BIGA FROM A,B,C
194  C          AFTER THAT A CALL TO SOLTRI WILL COMPLETE THE SOLUTION FOR
195  C          A GIVEN RHS D.  THIS METHOD IS EFFICIENT WHEN SOLVING WITH

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196 C          SEVERAL DIFFERENT RHS BUT THE SAME MATRIX M, AS ALL
197 C          DIVISIONS ARE DONE ONLY ONCE, AND SOLTRI PERFORMS N
198 C          VECTORIZABLE MULTIPLICATIONS FOLLOWED BY TWO SIMILAR
199 C          RECURSIVELY DEFINED, NON-VECTORIZABLE , LOOPS, EACH HAVING
200 C          N MULTIPLIES AND N SUBTRACTIONS, PER RHS.  EXAMINATION OF
201 C          SOLTRI SHOWS BIGA(1), BIGC(N) ARE NOT USED.
202 C
203 C
204 C
205 C          IMPLICIT REAL*8 (A-H,O-Z)
206 C          DIMENSION BIGE(N),BIGC(N),BIGA(N),A(N),B(N),C(N)
207 C          BIGE(1)=1.0DO/B(1)
208 C          BIGC(1)=C(1)*BIGE(1)
209 C          C(N)=0.0DO
210 C          DO 50 I=2,N
211 C             I1=I-1
212 C             AI=A(I)
213 C             BIGE(I)=1.0/(B(I)-AI*BIGC(I1))
214 C             BIGC(I)=BIGE(I)*C(I)
215 C             BIGA(I)=AI*BIGE(I)
216 C          50 CONTINUE
217 C          RETURN
218 C          END
219 C          SUBROUTINE SOLTRI(N,P,D,BIGD,BIGE,BIGC,BIGA)
220 C
221 C S/R:  SOLTRI - USED AFTER SETTRI TO COMPLETE THE SOLUTION TO THE
222 C          TRI-DIAGONAL MATRIX PROBLEM.  SEE SETTRI FOR DETAILS.
223 C
224 C
225 C          IMPLICIT REAL*8 (A-H,O-Z)
226 C          DIMENSION P(N),D(N),BIGD(N),BIGE(N),BIGC(N),BIGA(N)
227 C          DO 20 I=1,N
228 C             BIGD(I)=D(I)*BIGE(I)
229 C          20 CONTINUE
230 C          DO 50 I=2,N
231 C             I1=I-1
232 C             BIGD(I)=BIGD(I)-BIGA(I)*BIGD(I1)
233 C          50 CONTINUE
234 C          P(N)=BIGD(N)
235 C          DO 100 I=2,N
236 C             IREV=N-I+1
237 C             IREV1=IREV+1
238 C             P(IREV)=BIGD(IREV)-BIGC(IREV)*P(IREV1)
239 C          100 CONTINUE
240 C          RETURN
241 C          END
242 C          SUBROUTINE PSOLVE(R,RHS,NI,NJ,ALONGX,ALONGY,FOURTH)
243 C
244 C S/R:  PSOLVE : SOLUTION OF THE MATRIX PROBLEMS ASSOCIATED WITH THE
245 C          PROJECTION OPERATOR (ALONG ANY OR ALL DIRECTIONS, AS
246 C          DECIDED BY THE LOGICAL VARIABLES ALONGX,ALONGY)
247 C
248 C          R = INVERSE PROJECTION ( RHS )
249 C
250 C          AUTHOR: A. STANIFORTH - REVISED DEC 78 TO EXECUTE FASTER
251 C          REVISION 001: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A
252 C
253 C          ARGUMENTS:
254 C             OUT      - R          - RESULT
255 C             IN       - RHS        - RIGHT-HAND SIDE
256 C             NI       - NI         - X DIMENSION
257 C             NJ       - NJ         - Y DIMENSION
258 C             ALONGX   - IF .TRUE. , INVERSE PROJECT IN X DIRECTION
259 C             ALONGY   - IF .TRUE. , INVERSE PROJECT IN Y DIRECTION
260 C
261 C

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262      C
263      IMPLICIT REAL*8 (A-H,O-Z)
264      DIMENSION R(NI,NJ),RHS(NI,NJ)
265      COMMON / HXMESH / HX(1)
266      COMMON / HYMesh / HY(1)
267      COMMON / WKS101 / A(1)
268      COMMON / WKS103 / B(1)
269      COMMON / WKS104 / C(1)
270      COMMON / WKS105 / CF1(1)
271      COMMON / WKS106 / CF2(1)
272      COMMON / WKS107 / CF3(1)
273      COMMON / WKS109 / TEMP2(1)
274      LOGICAL ALONGX,ALONGY,FOURTH
275      C
276      C SET UP WTS ASSOCIATED WITH PROJN OPERATOR.
277      C WT=2. CORRESPONDS TO NORMAL USAGE.
278      C
279      WTX=2.0DO
280      WTY=2.0DO
281      IF(FOURTH) WTX=5.0DO
282      IF(FOURTH) WTY=5.0DO
283      C
284      C CODE FOR SOLVING ALONG X DIRECTION
285      C
286      IF(.NOT.ALONGX) GO TO 200
287      CALL SETABC(A,B,C,HX,WTX,NI)
288      CALL SETTRI(CF1,CF2,CF3,A,B,C,NI)
289      DO 130 J=1,NJ
290      CALL SOLTRI(NI,R(1,J),RHS(1,J),TEMP2,CF1,CF2,CF3)
291      130 CONTINUE
292      GO TO 250
293      200 CONTINUE
294      DO 450 J=1,NJ
295      DO 450 I=1,NI
296      450 R(I,J)=RHS(I,J)
297      C
298      C CODE FOR SOLVING ALONG Y DIRECTION
299      C
300      250 IF(.NOT.ALONGY) RETURN
301      CALL SETABC(A,B,C,HY,WTY,NJ)
302      CALL SETTRI(CF1,CF2,CF3,A,B,C,NJ)
303      DO 460 I=1,NI
304      R(I,1)=CF1(1)*R(I,1)
305      460 CONTINUE
306      DO 215 J=2,NJ
307      J1=J-1
308      CF3J=CF3(J)
309      CF1J=CF1(J)
310      DO 215 I=1,NI
311      R(I,J)=-CF3J*R(I,J1)+CF1J*R(I,J)
312      215 CONTINUE
313      DO 225 J=2,NJ
314      JREV=NJ-J+1
315      JREV1=JREV+1
316      CF2J=CF2(JREV)
317      DO 225 I=1,NI
318      R(I,JREV)=-CF2J*R(I,JREV1)+R(I,JREV)
319      225 CONTINUE
320      RETURN
321      ENO
END OF FILE

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1      SUBROUTINE OUTPUT(II,XCI,TAU,OMEG,XI,YI,X,Y,NI,NJ,NIU,NJU,
2      1 FX,FY,FX,Y,D,WORK,NIP,NJP,OUT,DT)
3      C
4      C ROUTINE TO PRODUCE MAPS OF FORECAST FIELDS.
5      C
6      IMPLICIT REAL*8 (A-H,O-Z)
7      DIMENSION XCI(NI,NJ),TAU(NI,NJ),OMEG(NI,NJ)
8      DIMENSION XI(NIP),YI(NJP),X(NIU),Y(NJU)
9      DIMENSION D(4,NIU,NJU)
10     DIMENSION FX(NIP,NJP),FY(NIP,NJP),FX,Y(NIP,NJP),WORK(NJU,8)
11     DIMENSION OUT(NIP,NJP)
12     DIMENSION E(281,45)
13     COMMON / HXMES1 / HX1(1)
14     COMMON / HYMES1 / HY1(1)
15     COMMON / WKS1X1 / WK1(1)
16     COMMON / WKS1X2 / WK2(1)
17     COMMON / WKS1X3 / WK3(1)
18     COMMON / WKS1X4 / WK4(1)
19     COMMON / WKS1X5 / WK5(1)
20     COMMON / CONSTA / C4,C5
21     LOGICAL KDER
22     C
23     C USES ROUTINE INTRPT TO PRODUCE VALUES OF FIELDS ON UNIFORM GRID
24     C
25     KDER=.FALSE.
26     KL=2
27     ITIME=DFLOAT(II-1)*DT/.370D0+.50D0
28     CALL INTRPT(D,NIU,NJU,XCI,NIP,NJP,X,Y,XI,YI,FX,
29     1 FY,FX,Y,HX1,HY1,KL,WK1,WK2,WK3,WK4,WK5,WORK,KDER)
30     DO 100 J=1,NJU
31     DO 100 I=1,NIU
32     100 E(I,J)=C4*D(1,I,J)
33     CALL FPLLOT(E,NIU,NJU,1)
34     CALL INTRPT(D,NIU,NJU,TAU,NIP,NJP,X,Y,XI,YI,FX,
35     1 FY,FX,Y,HX1,HY1,KL,WK1,WK2,WK3,WK4,WK5,WORK,KDER)
36     DO 300 J=1,NJU
37     DO 300 I=1,NIU
38     300 E(I,J)=C4*D(1,I,J)
39     CALL FPLLOT(E,NIU,NJU,1)
40     CALL INTRPT(D,NIU,NJU,OMEG,NIP,NJP,X,Y,XI,YI,FX,
41     1 FY,FX,Y,HX1,HY1,KL,WK1,WK2,WK3,WK4,WK5,WORK,KDER)
42     DO 475 J=1,NJU
43     DO 475 I=1,NIU
44     475 E(I,J)=C5*D(1,I,J)
45     CALL FPLLOT(E,NIU,NJU,0)
46     RETURN
47     END
48     SUBROUTINE MSHFLT(FIELD,WKX,WKY,CON,NI,NJ,IPOW)
49     C
50     C S/R: MSHFLT - REDUCES AMPLITUDE OF FIELD (IN-PLACE) FOR REGIONS OF
51     C POOR HORIZONTAL RESOLUTION
52     C FIELD(I,J) = FAC*(FIELD(I,J)-CON)+CON WHERE
53     C FAC = (HMIN(I,J)/H(I,J))*IPOW,
54     C H(I,J) = SQRT(HX(I)*HX(I) + HY(J)*HY(J)),
55     C HMIN(I,J)= SMALLEST(H(I,J)).
56     C
57     C AUTHOR: ROGER DALEY - SUMMER 1977
58     C
59     C REVISION 001: A. STANIFORTH - C. THIBEAULT JAN 79 DOCUMENTATION
60     C REVISION 002: T. GOOS - 1979-80 ADAPTED FOR AMDAHL AT U OF A.
61     C
62     C CALL: CALL MSHFLT(FIELD,WKX,WKY,CON,NI,NJ,IPOW)
63     C

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64      C  ARGUMENTS:
65      C  IN-OUT - FIELD - FIELD USED FOR IN-PLACE AMPLITUDE REDUCTION
66      C  IN      - XI      - X-SPECIFICATION OF GRID
67      C          - YI      - Y-SPECIFICATION OF GRID
68      C          - WKX      - WORK(NI)
69      C          - WKY      - WORK(NJ)
70      C          - CON      - PARAMETER CONTROLLING MEAN OF FIELD
71      C          - NI      - X-DIMENSION
72      C          - NJ      - Y-DIMENSION
73      C          - IPOW     - PARAMETER OF AMPLITUDE REDUCTION
74      C
75      C
76      C  IMPLICIT REAL*8 (A-H,O-Z)
77      C  DIMENSION FIELD(NI,NJ)
78      C  DIMENSION WKX(NI),WKY(NJ)
79      C  FMIN=1.D12
80      C  DO 50 J=1,NJ
81      C  DO 50 I=1,NI
82      C  SX = DSQRT(WKX(I)*WKX(I) + WKY(J)*WKY(J))
83      C  IF(SX.LE.FMIN) FMIN=SX
84      C  50 CONTINUE
85      C  DO 100 J=1,NJ
86      C  DO 100 I=1,NI
87      C  SX = DSQRT(WKX(I)*WKX(I) + WKY(J)*WKY(J))
88      C  FAC = (FMIN/SX)**IPOW
89      C  100 FIELD(I,J)=FAC*(FIELD(I,J)-CON)+CON
90      C  RETURN
91      C  END
92      C  SUBROUTINE REL(PHIG,PG,NI,NJ,ALX,RELF,XTOL)
93      C  IMPLICIT REAL *8(A-H,O-Z)
94      C
95      C  PERFORMS RELAXATION SOLUTION FOR POISSONS EQUATION.
96      C
97      C  DIMENSION PG(NI,NJ),PHIG(NI,NJ)
98      C  COMMON / HXMESH / HX(1)
99      C  COMMON / RATIO1 / RAT1(63,27)
100     C  COMMON / RATIO1 / RAT1(63,27)
101     C  COMMON / RATIO4 / RAT4(63,27)
102     C  FIVE=5.0DO
103     C  NIM=NI-1
104     C  NJM=NJ-1
105     C  NI2=NI-2
106     C  NJ2=NJ-2
107     C  NJ4=NJ-4
108     C  DNI3=2.0DO*ALX
109     C  DNI3=1.0DO/DNI3
110     C  RELFP=1.0DO-RELF
111     C  DO 1000 K=1,140
112     C  PHI=0.0DO
113     C  SUM=0.0
114     C  DO 20 I=2,NIM
115     C  I1=I-1
116     C  H=HX(I)+HX(I1)
117     C  20 SUM=SUM+H*PHIG(I,2)
118     C  H=HX(1)+HX(NIM)
119     C  SUM=SUM+PHIG(1,2)*H
120     C  SUM=SUM*DNI3
121     C  DO 25 I=1,NI
122     C  25 PHIG(I,1)=SUM
123     C  DO 300 J=2,NJM
124     C  J1=J-1
125     C  SUP=(PHIG(1,J1)+PHIG(NI,J1))/2.0DO
126     C  PHIG(1,J1)=SUP
127     C  PHIG(NI,J1)=SUP
128     C  J2=J+1
129     C  DO 300 I=2,NIM

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130      Z=RELF*PHIG(I,J)
131      I1=I-1
132      I2=I+1
133      RA11=RAT1(I1,J1)
134      RAO1=RAT1(I,J1)
135      RAI0=RAT1(I1,J)
136      RA=RAT1(I,J)
137      RAI11=RATI1(I1,J1)
138      RAI01=RATI1(I,J1)
139      RAI10=RATI1(I1,J)
140      RAI=RATI1(I,J)
141      SUM=(RA11+RAI11)*PHIG(I1,J1)
142      SUM=SUM+(RAO1+RAI01)*PHIG(I2,J1)
143      SUM=SUM+(RAI0+RAI10)*PHIG(I1,J2)
144      SUM=SUM+(RA+RAI)*PHIG(I2,J2)
145      SUM=SUM+(FIVE*(RAI+RAI10)-RA-RAI0)*PHIG(I,J2)
146      SUM=SUM+(FIVE*(RAI11+RAI01)-RA11-RAO1)*PHIG(I,J1)
147      SUM=SUM+(FIVE*(RAI0+RAI1)-RAI10-RAI11)*PHIG(I1,J)
148      SUM=SUM+(FIVE*(RA+RAO1)-RAI-RAI01)*PHIG(I2,J)
149      SUM=SUM/FIVE
150      SUM=(SUM+PG(I,J))*RAT4(I,J)
151  C      WRITE(7,806) SUM,RAT4(I,J),PG(I,J),PHIG(I,J)
152      PHIO=PHIG(I,J)
153      PHIG(I,J)=Z+RELF*SUM
154      IF(I.LT.18.OR.I.GT.46) GO TO 805
155      IF(J.LE.5) GO TO 805
156      IF(J.GE.NJ4) GO TO 805
157      IF(PHIO.GT.-1.OD-50.AND.PHIO.LT.1.OD-50) GO TO 805
158      PHIO=DABS((PHIG(I,J)-PHIO)/PHIO)
159      PHI=DMAX1(PHIO,PHI)
160  805  CONTINUE
161  806  FORMAT(4(2X,E12.6))
162  300  CONTINUE
163      SUM=0.ODO
164      DO 40 I=2,NIM
165          I1=I-1
166          H=HX(I)+HX(I1)
167      40  SUM=SUM+H*PHIG(I,NJM)
168          H=HX(1)+HX(NIM)
169          SUM=SUM+PHIG(1,NJM)*H
170          SUM=SUM*DNI3
171          DO 45 I=1,NI
172      45  PHIG(I,NJ)=SUM
173          IF(K.EQ.1) GO TO 1000
174          IF(PHI.LT.XTOL) GO TO 1505
175  1000 CONTINUE
176  1505 CONTINUE
177      WRITE(7,1049) K,PHI
178  1049 FORMAT('    K = ',I5,4X,E12.6)
179      RETURN
180      END
181      SUBROUTINE TRANS(OUT,XCI,NI,NJ,NIP,NJP)
182      IMPLICIT REAL*8 (A-H,O-Z)
183      DIMENSION OUT(NIP,NJP),XCI(NI,NJ)
184      NIM=NI-1
185      NJM=NJ-1
186      DO 50 J=2,NJM
187          J1=J-1
188          DO 50 I=2,NIM
189              I1=I-1
190      50  OUT(I1,J1)=XCI(I,J)
191      RETURN
192      END
193      SUBROUTINE RELHEL(PHIG,PG,NI,NJ,ALX,ALAM,RELF,XTOL)
194      IMPLICIT REAL *8(A-H,O-Z)
195  C

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196 C PERFORMS RELAXATION SOLUTION FOR HELMHOLTZ PROBLEM.
197 C
198     DIMENSION PG(NI,NJ),PHIG(NI,NJ)
199     COMMON / HXMESH / HX(1)
200     COMMON / RATIO1 / RAT1(63,27)
201     COMMON / RATIO11 / RATI1(63,27)
202     COMMON / RATIO4 / RAT4(63,27)
203     COMMON / RATIOO / RATO(63,27)
204     COMMON / RATIO2 / RAT2(63,27)
205     COMMON / RATIO3 / RAT3(63,27)
206     COMMON / RATIO5 / RAT5(63,27)
207     FIVE=5.000
208     ALAMP=ALAM/60.00
209     NIM=NI-1
210     NJM=NJ-1
211     NI2=NI-2
212     NJ2=NJ-2
213     NJ4=NJ-4
214     DNI3=2.000*ALX
215     DNI3=1.000/ONI3
216     RELFP=1.000-RELF
217     DO 1000 K=1,140
218     PHI=0.000
219     SUM=0.0
220     DO 20 I=2,NIM
221     I1=I-1
222     H=HX(I)+HX(I1)
223     20 SUM=SUM+H*PHIG(I,2)
224     H=HX(1)+HX(NIM)
225     SUM=SUM+PHIG(1,2)*H
226     SUM=SUM*ONI3
227     DO 25 I=1,NI
228     25 PHIG(I,1)=SUM
229     DO 300 J=2,NJM
230     J1=J-1
231     SUP=(PHIG(1,J1)+PHIG(NI,J1))/2.000
232     PHIG(1,J1)=SUP
233     PHIG(NI,J1)=SUP
234     J2=J+1
235     DO 300 I=2,NIM
236     828 CONTINUE
237     Z=RELP*PHIG(I,J)
238     I1=I-1
239     I2=I+1
240     P11=PHIG(I1,J1)
241     P21=PHIG(I2,J1)
242     P12=PHIG(I1,J2)
243     P22=PHIG(I2,J2)
244     PO2=PHIG(I,J2)
245     PO1=PHIG(I,J1)
246     P1O=PHIG(I1,J)
247     P2O=PHIG(I2,J)
248     RA11=RAT1(I1,J1)
249     RAO1=RAT1(I,J1)
250     RA1O=RAT1(I1,J)
251     RA=RAT1(I,J)
252     RAI11=RATI1(I1,J1)
253     RAO1=RATI1(I,J1)
254     RAI1O=RATI1(I1,J)
255     RAI=RATI1(I,J)
256     SUM=(RA11+RAI11)*P11
257     SUM=SUM+(RAO1+RAIO1)*P21
258     SUM=SUM+(RA1O+RAI1O)*P12
259     SUM=SUM+(RA+RAI)*P22
260     SUM=SUM+(FIVE*(RAI+RAI1O)-RA-RA1O)*PO2
261     SUM=SUM+(FIVE*(RAI11+RAIO1)-RA11-RAO1)*PO1

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262      SUM=SUM+(FIVE*(RA10+RA11)-RA110-RA111)*P10
263      SUM=SUM+(FIVE*(RA+RAO1)-RAI-RAIO1)*P20
264      SUM=SUM/5.0DO
265      SUM5=RAT2(I,J)*PO2 +RAT3(I,J)*P20
266      SUM5=SUM5+RAT3(I1,J)*P10 + RAT2(I,J1)*PO1
267      SUM5=SUM5*FIVE+RATO(I,J)*P22+RATO(I1,J)*P12
268      SUM5=SUM5+RATO(I1,J1)*P11+RATO(I,J1)*P21
269      SUM5=SUM5*ALAMP
270      SUM=SUM-SUM5
271      210 SUM=(SUM+PG(I,J))*RAT5(I,J)
272      C      WRITE(7,600) SUM,SUM5,RAT5(I,J),PG(I,J),PHIG(I,J)
273      PHIO=PHIG(I,J)
274      PHIG(I,J)=Z+RELF*SUM
275      IF(I.LT.18.OR.I.GT.46) GO TO 805
276      IF(J.LE.5) GO TO 805
277      IF(J.GE.NJ4) GO TO 805
278      IF(PHIO.GT.-1.0D-50.AND.PHIO.LT.1.0D-50) GO TO 805
279      PHIO=DABS((PHIG(I,J)-PHIO)/PHIO)
280      PHI=DMAX1(PHIO,PHI)
281      805 CONTINUE
282      827 FORMAT(3I6,E12.6,3X,E12.6)
283      300 CONTINUE
284      SUM=0.0DO
285      DO 40 I=2,NIM
286      I1=I-1
287      H=HX(I)+HX(I1)
288      40 SUM=SUM+H*PHIG(I,NJM)
289      H=HX(1)+HX(NIM)
290      SUM=SUM+PHIG(1,NJM)*H
291      SUM=SUM*DNI3
292      DO 45 I=1,NI
293      45 PHIG(I,NJ)=SUM
294      IF(K.EQ.1) GO TO 1000
295      IF(PHI.LT.XTOL) GO TO 1505
296      1000 CONTINUE
297      1505 CONTINUE
298      WRITE(7,1049) K,PHI
299      1049 FORMAT(' HELM. K=',I5,4X,E12.6)
300      610 FORMAT(1X,8(1X,E12.6))
301      600 FORMAT(8(2X,E12.6))
302      RETURN
303      END
304      SUBROUTINE FPL0T(Z,NIU,NJU,IFLAG)
305      REAL*8 Z(NIU,NJU)
306      C
307      C PLOTS FIELDS IN 2 SECTIONS; ONE OVER THE NON-UNIFORM PORTION
308      C OF THE GRID AND ONE INCLUDING THE UNIFORM PORTION OF THE
309      C GRID.
310      C
311      DIMENSION F(92,45)
312      DIMENSION CVAL1(19),CVAL2(19),VOP(8),IOP(8)
313      C
314      C CONTOUR VALUES FOR HEIGHT AND THICKNESS FIELDS.
315      C
316      DATA CVAL1/540.0,480.0,420.0,360.0,300.0,240.0,180.0,
317      1 120.0,60.0,0.0,-60.0,-120.0,-180.0,-240.0,-300.0,
318      2 -360.0,-420.0,-480.0,-540.0/
319      C
320      C CONTOUR VALUES FOR OMEGA FIELD.
321      C
322      DATA CVAL2/3.6,3.2,2.8,2.4,2.0,1.6,1.2,.8,.4,0.0,
323      1 -.4,-.8,-1.2,-1.6,-2.0,-2.4,-2.8,-3.2,-3.6/
324      DATA VOP/0.0,0.0,0.0,0.0,0.0,-1.0,0.0,0.0,0.0/
325      IOP(1)=1
326      IOP(2)=1
327      IOP(3)=0

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328         IOP(4)=1
329         IOP(6)=0
330         IOP(7)=1
331         IOP(8)=0
332         IOP(5)=IFLAG
333         NC=19
334         CALL ORIGIN(999,6.5,3.5,1.0,1.0)
335     C
336     C PLOT VARIABLE PORTION
337     C
338         DO 100 J=1,45
339         DO 100 I=1,92
340     100 F(I,J)=Z(I,J)
341         IF(IFLAG.EQ.0) CALL CONTUR(5.5,2.75,F,92,92,45,CVAL2,NC,IOP,VOP)
342         IF(IFLAG.EQ.1) CALL CONTUR(5.5,2.75,F,92,92,45,CVAL1,NC,IOP,VOP)
343         CALL ORIGIN(999,6.5,3.5,1.0,1.0)
344     C
345     C PLOT UNIFORM PORTION
346     C
347         DO 150 J= 1,45
348         DO 150 I=92,183
349         I1=I-91
350     150 F(I1,J)=Z(I,J)
351         IF(IFLAG.EQ.0) CALL CONTUR(5.5,2.75,F,92,92,45,CVAL2,NC,IOP,VOP)
352         IF(IFLAG.EQ.1) CALL CONTUR(5.5,2.75,F,92,92,45,CVAL1,NC,IOP,VOP)
353         RETURN
354         END
END OF FILE

```



```

1      C
2      C
3      C   FILE:  DIAG
4      C
5      C   FILES
6      C       OBJECT - DIAG
7      C
8      C       INPUT
9      C           1 = FEDIAG
10     C           2 = SPDIAG
11     C           3 = VMESH
12     C           4 = MESHD
13     C           5 = SPECN
14     C       OUTPUT
15     C           6 = ENERGIES AND ENSTROPY
16     C           7 = ERRORS IN MAP FORM
17     C           8 = MEAN ERRORS AND S1 SCORES
18     C           9 = -PDIAG
19     C
20     C   COMPUTES DIAGNOSTIC VALUES FROM THE SPECTRAL AND FEM FORECASTS.
21     C   THESE VALUES INCLUDE:
22     C       KINETIC ENERGIES
23     C       AVAILABLE POTENTIAL ENERGIES
24     C       POTENTIAL ENSTROPHYS
25     C       MEAN DIFFERENCE BETWEEN SOLUTIONS
26     C       MEAN ABSOLUTE DIFFERENCE BETWEEN SOLUTIONS
27     C       S1 SCORES
28     C
29     C   USES SYSTEM ROUTINES TO PLOT GRAPHS OF THESE VALUES.
30     C
31     C       IMPLICIT REAL*8 (A-H,O-Z)
32     C       DIMENSION XCSP(63,27),TASP(63,27),XCFE(63,27),TAFE(63,27)
33     C       DIMENSION XCSP0(63,27),TASPO(63,27),XCFE0(63,27),TAFE0(63,27)
34     C       DIMENSION X(281),Y(45),XI(63),YI(27)
35     C       DIMENSION HX1(63),HY1(27)
36     C       DIMENSION APE(9),EKE(9),TOT(9),PE(9)
37     C       DIMENSION APEF(9),EKEF(9),TOTF(9),PEF(9)
38     C       DIMENSION DEXC(9),ADEXC(9),DETA(9),ADETA(9)
39     C       DIMENSION S1XC(9),S1TA(9)
40     C       REAL ZZ(11,9),T(11)
41     C       DIMENSION ZKESP(9),EKESP(9),ZAPESP(9),EAPESP(9)
42     C       DIMENSION ZKEFE(9),EKEFE(9),ZAPEFE(9),EAPPEFE(9)
43     C       DIMENSION APET(9),APETF(9),EKET(9),EKETF(9)
44     C       DIMENSION TOT(9),TOTF(9),PET(9),PETF(9)
45     C       REAL TI1(3),TI2(4),TI3(2),TI4(5)
46     C       LOGICAL*1 LFMT(1) /'*/
47     C       DATA T/0.0,6.0,12.0,18.0,24.0,30.0,36.0,42.0,48.0,0.0,0.0/
48     C       DATA TI1/'ENER','GY','(J)'/
49     C       DATA TI2/'ENST','ROPH','Y(KG','S)'/
50     C       DATA TI3/'S1 S','CORE'/
51     C       DATA TI4/'MEAN','DIF','FERE','NCE','(M)'/
52     C       COMMON / HXMESH / HX(63)
53     C       COMMON / HYMESH / HY(27)
54     C       COMMON / WKS1D1 / WK1(63)
55     C       COMMON / WKS1D2 / WK2(63)
56     C       COMMON / WKS1D3 / WK3(63)
57     C       COMMON / WKS1D4 / WK4(63)
58     C       COMMON / WKS1D5 / WK5(63)
59     C       COMMON / WKS1D6 / WK6(63)
60     C       COMMON / WKS1D7 / WK7(63)
61     C       COMMON / WKS1D8 / WK8(63)
62     C       COMMON / WKS1D9 / WK9(63)
63     C       COMMON / WKS110 / WK10(63)
64     C       COMMON / WORKA / WKA(63,27)

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65      COMMON / WORKB / WKB(63,27)
66      COMMON / WORKC / WKC(63,27)
67      COMMON / WORKD / WKD(63,27)
68      COMMON / WORKE / WKE(63,27)
69      COMMON / WORKF / WKF(63,27)
70      C
71      C SET CONSTANTS
72      C
73      C
74      C CXCI=L*L*FO*FO/G
75      C
76      CXCI=2.15357266D4
77      C
78      C CPE=2.0D0*FO*FO/(DP*DP*SIG)
79      C
80      CPE=3.0382098D-12
81      CEKE=2.1383390D22
82      CAPE=6.4967228D10
83      CCPE=1.0691695E22
84      C
85      C READ PARAMETERS OF GRID.
86      C
87      READ(4,LFMT) NIU,NJU
88      READ(4,LFMT) NI,NJ
89      READ(3,LFMT)(X(I),I=1,NIU)
90      READ(3,LFMT)(Y(J),J=1,NJU)
91      READ(3,LFMT)(XI(I),I=1,NI)
92      READ(3,LFMT)(YI(J),J=1,NJ)
93      NIP=NI
94      NJP=NJ
95      C
96      C CALCULATE GRID LENGTHS.
97      C
98      CALL CALH(HX1,XI,NIP)
99      CALL CALH(HY1,YI,NJP)
100     NJPM=NJP-1
101     NIPM=NIP-1
102     HX1(NIP)=HX1(NIPM)
103     HY1(NJP)=HY1(NJPM)
104     NIM=NI-1
105     NJM=NJ-1
106     DO 60 I=1,NIP
107     60 HX(I)=HX1(I)
108     DO 65 J=1,NJP
109     65 HY(J)=HY1(J)
110     C
111     C READ INITIAL AMPLITUDES
112     C
113     READ(5,LFMT) X1,X2,X3,T1,T2,T3
114     C
115     C N = NO. OF WAVES, M = NO. OF TIME STEPS
116     C DT = TIMESTEP, IPT2 = NO. OF HOURS BETWEEN MAPS
117     C
118     READ(5,LFMT) N,M,DT,IDT2
119     C
120     C CALCULATE NO. OF RECORDS ON FILE
121     C
122     NR=(M-1)/IDT2+1
123     C
124     C SET BOUNDARIES OF COMPARISON GRID(S)
125     C
126     IL1=23
127     IR1=40
128     JT1=18
129     JB1=10
130     C

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131      C   MAIN LOOP
132      C
133          H=HX(30)
134          DO 1000 II=1,NR
135              S1XC(II)=0.ODO
136              S1TA(II)=0.ODO
137      C
138      C   READ SPECTRAL FIELDS
139      C
140          READ(2) XCSP
141          READ(2) TASP
142      C
143      C   READ FEM FIELDS.
144      C
145          READ(1) XCFE
146          READ(1) TAFE
147          IF(II.NE.1) GO TO 200
148          DO 180 J=1,NJ
149              DO 180 I=1,NI
150                  XCSP0(I,J)=XCSP(I,J)
151                  TASPO(I,J)=TASP(I,J)
152                  XCFEO(I,J)=XCFE(I,J)
153                  TAFE0(I,J)=TAFE(I,J)
154      180      CONTINUE
155      200      CONTINUE
156      C
157      C   CALCULATE ENERGIES AND ENSTROPHIES.
158      C
159          CALL APOTEN(APE(II),IL1,IR1,JT1,JB1,TASP,NI,NJ,
160      1 H,APET(II),1)
161          CALL APOTEN(APEF(II),IL1,IR1,JT1,JB1,TAFE,NI,NJ,
162      1 H,APETF(II),1)
163          CALL EKINEN(XCSP,TASP,NI,NJ,H,EKE(II),IL1,IR1,JT1,
164      1 JB1,EKET(II),1)
165          CALL EKINEN(XCFE,TAFE,NI,NJ,H,EKEF(II),IL1,IR1,JT1,
166      1 JB1,EKETF(II),1)
167          CALL POTENS(XCSP,TASP,NI,NJ,H,PE(II),IL1,IR1,JT1,
168      1 JB1,CPE,PET(II),1)
169          CALL POTENS(XCFE,TAFE,NI,NJ,H,PEF(II),IL1,IR1,JT1,
170      1 JB1,CPE,PETF(II),1)
171          PE(II)=PE(II)*CCPE
172          APE(II)=APE(II)*CAPE
173          EKE(II)=EKE(II)*CEKE
174          TOT(II)=APE(II)+EKE(II)
175          PEF(II)=PEF(II)*CCPE
176          APEF(II)=APEF(II)*CAPE
177          EKEF(II)=EKEF(II)*CEKE
178          TOTF(II)=APEF(II)+EKEF(II)
179          PET(II)=PET(II)*CCPE
180          APET(II)=APET(II)*CAPE
181          EKET(II)=EKET(II)*CEKE
182          TOTT(II)=APET(II)+EKET(II)
183          PETF(II)=PETF(II)*CCPE
184          APETF(II)=APETF(II)*CAPE
185          EKETF(II)=EKETF(II)*CEKE
186          TOTTF(II)=APETF(II)+EKETF(II)
187          IF(II.EQ.1) GO TO 1001
188      C
189      C   CALCULATE DIFFERENCE IN LAST (II-1) OUTPUTS
190      C
191      C   TRUE DIFF. FOR XCI FIELD
192      C
193          CALL GDADGD(WKA,1.ODO,XCSP,-1.ODO,XCSP0,NI,NJ)
194      C
195      C   FORECAST DIFF. FOR XCI FIELD
196      C

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197      CALL GDADGD(WKB,1.ODO,XCFE,-1.ODO,XCFEO,NI,NJ)
198      C
199      C FIND ERRDR IN FDRECAST.
200      C
201      CALL GDADGD(WKC,1.ODO,WKA,-1.ODO,WKB,NI,NJ)
202      CALL MEAN(WKC,NI,NJ,DEXC(II),IL1,IR1,JT1,JB1)
203      CALL AMEAN(WKC,NI,NJ,ADEXC(II),IL1,IR1,JT1,JB1)
204      C
205      C CALCULATE DIFFERENCE IN LAST (II-1) DUTPUTS
206      C
207      C TRUE DIFF. FDR TAU FIELD.
208      C
209      CALL GDADGD(WKA,1.ODO,TASP,-1.ODO,TASPO,NI,NJ)
210      C
211      C FDRECAST DIFF. FDR TAU FIELD.
212      C
213      CALL GDADGD(WKB,1.ODO,TAFE,-1.ODO,TAFEO,NI,NJ)
214      C
215      C FIND ERRDR IN FDRECAST.
216      C
217      CALL GDADGD(WKC,1.ODO,WKA,-1.ODO,WKB,NI,NJ)
218      CALL MEAN(WKC,NI,NJ,DETA(II),IL1,IR1,JT1,JB1)
219      CALL AMEAN(WKC,NI,NJ,ADETA(II),IL1,IR1,JT1,JB1)
220      C
221      C CALCULATE S1 SCORES
222      C
223      C FDR XCI
224      C
225      CALL DXDYDS(WKB,XCSP,1.ODO,.TRUE...FALSE.,NI,NJ)
226      CALL PSDLVE(WKA,WKB,NI,NJ,.TRUE...FALSE...FALSE.)
227      CALL DXDYDS(WKC,XCSP,1.ODO,.FALSE...TRUE.,NI,NJ)
228      CALL PSDLVE(WKB,WKC,NI,NJ,.FALSE...TRUE...FALSE.)
229      CALL DXDYDS(WKD,XCFE,1.ODO,.TRUE...FALSE.,NI,NJ)
230      CALL PSDLVE(WKC,WKD,NI,NJ,.TRUE...FALSE...FALSE.)
231      CALL DXDYDS(WKE,XCFE,1.ODO,.FALSE...TRUE.,NI,NJ)
232      CALL PSDLVE(WKD,WKE,NI,NJ,.FALSE...TRUE...FALSE.)
233      C
234      C ERROR IN X-PRESS. GRAD - WKE
235      C
236      CALL GDADGD(WKE,1.ODO,WKC,-1.ODO,WKA,NI,NJ)
237      C
238      C ERRDR IN Y-PRESS. GRAD - WKF
239      C
240      CALL GDADGD(WKF,1.ODO,WKD,-1.ODO,WKB,NI,NJ)
241      C
242      C FORM SUM DF ABS. OF ERRORS
243      C
244      CALL SUM(WKE,NI,NJ,SUM1,IL1,IR1,JT1,JB1)
245      CALL SUM(WKF,NI,NJ,SUM2,IL1,IR1,JT1,JB1)
246      SUM3=SUM1+SUM2
247      CALL SUMB(WKC,WKA,NI,NJ,SUM1,IL1,IR1,JT1,JB1)
248      CALL SUMB(WKD,WKB,NI,NJ,SUM2,IL1,IR1,JT1,JB1)
249      SUM4=SUM1+SUM2
250      IF(DABS(SUM4).LT.1.OD-50) GD TO 435
251      S1XC(II)=100.ODO*SUM3/SUM4
252      C
253      C CALCULATE S1 SCDRE FDR TAU.
254      C
255      435 CALL DXDYDS(WKB,TASP,1.ODO,.TRUE...FALSE.,NI,NJ)
256      CALL PSDLVE(WKA,WKB,NI,NJ,.TRUE...FALSE...FALSE.)
257      CALL DXDYDS(WKC,TASP,1.ODO,.FALSE...TRUE.,NI,NJ)
258      CALL PSOLVE(WKB,WKC,NI,NJ,.FALSE...TRUE...FALSE.)
259      CALL DXDYDS(WKD,TAFE,1.ODO,.TRUE...FALSE.,NI,NJ)
260      CALL PSDLVE(WKC,WKD,NI,NJ,.TRUE...FALSE...FALSE.)
261      CALL DXDYDS(WKE,TAFE,1.ODO,.FALSE...TRUE.,NI,NJ)
262      CALL PSOLVE(WKD,WKE,NI,NJ,.FALSE...TRUE...FALSE.)

```



```

263      C
264      C  ERROR IN X-PRESS. GRAD - WKE
265      C
266          CALL GDADGD(WKE,1.ODO,WKC,-1.ODO,WKA,NI,NJ)
267      C
268      C  ERROR IN Y-PRESS. GRAD - WKF
269      C
270          CALL GDADGD(WKF,1.ODO,WKD,-1.ODO,WKB,NI,NJ)
271      C
272      C  FORM SUM OF ABS. OF ERRORS
273      C
274          CALL SUM(WKE,NI,NJ,SUM1,IL1,IR1,JT1,JB1)
275          CALL SUM(WKF,NI,NJ,SUM2,IL1,IR1,JT1,JB1)
276          SUM3=SUM1+SUM2
277          CALL SUMB(WKC,WKA,NI,NJ,SUM1,IL1,IR1,JT1,JB1)
278          CALL SUMB(WKD,WKB,NI,NJ,SUM2,IL1,IR1,JT1,JB1)
279          SUM4=SUM1+SUM2
280          IF(DABS(SUM4).LT.1.OD-50) GO TO 436
281          S1TA(II)=100.ODO*SUM3/SUM4
282      436 CONTINUE
283      1001 CONTINUE
284      C
285      C  CALCULATE ZONAL AND EDDY ENERGIES.
286      C
287          CALL ZEENE(XCSP,TASP,NI,NJ,JT1,JB1,IL1,IR1,H,
288      1 ZKESP(II),EKESP(II),ZAPESP(II),EAPESP(II))
289          CALL ZEENE(XCFE,TAFE,NI,NJ,JT1,JB1,IL1,IR1,H,
290      1 ZKEFE(II),EKEFE(II),ZAPEFE(II),EAPFEFE(II))
291      1000 CONTINUE
292      C
293      C  INITIALIZE PLOTTING PROGRAM.
294      C
295          CALL PLOTS
296          CALL ORGEP(1.0,1.0,1.0)
297      C
298      C  PLOT ENERGIES
299      C
300          DO 650 IT=1,NR
301              ZZ(IT,1)=EKE(IT)
302              ZZ(IT,2)=EKEF(IT)
303              ZZ(IT,3)=APE(IT)
304              ZZ(IT,4)=APEF(IT)
305              ZZ(IT,5)=TOT(IT)
306              ZZ(IT,6)=TOTF(IT)
307      650 CONTINUE
308          CALL LPLOT(ZZ,T,NR+2,6,3.0,3.0,TI1,3)
309      C
310      C  PLOT KINETIC ENERGIES.
311      C
312          DO 690 IT=1,NR
313              ZZ(IT,1)=EKE(IT)
314              ZZ(IT,2)=EKEF(IT)
315              ZZ(IT,3)=ZKESP(IT)*CEKE
316              ZZ(IT,4)=ZKEFE(IT)*CEKE
317              ZZ(IT,5)=EKESP(IT)*CEKE
318              ZZ(IT,6)=EKEFE(IT)*CEKE
319      690 CONTINUE
320          CALL LPLOT(ZZ,T,NR+2,6,3.0,3.0,TI1,3)
321      C
322      C  PLOT POTENTIAL ENERGIES.
323      C
324          DO 695 IT=1,NR
325              ZZ(IT,1)=APE(IT)
326              ZZ(IT,2)=APEF(IT)
327              ZZ(IT,3)=ZAPESP(IT)*CAPE
328              ZZ(IT,4)=ZAPEFE(IT)*CAPE

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329      ZZ(IT,5)=EAPESP(IT)*CAPE
330      ZZ(IT,6)=EAPEFE(IT)*CAPE
331      695 CONTINUE
332      CALL LPLLOT(ZZ,T,NR+2,6,3.0,3.0,TI1,3)
333      C
334      C PLOT ENSTROPHY
335      C
336      DO 655 IT=1,NR
337      ZZ(IT,1)=PE(IT)
338      ZZ(IT,2)=PEF(IT)
339      655 CONTINUE
340      CALL LPLLOT(ZZ,T,NR+2,2,3.0,3.0,TI2,4)
341      C
342      C PLOT S1 SCORES
343      C
344      DO 660 IT=1,NR
345      ZZ(IT,1)=S1XC(IT)
346      ZZ(IT,2)=S1TA(IT)
347      660 CONTINUE
348      CALL LPLLOT(ZZ,T,NR+2,2,3.0,3.0,TI3,2)
349      C
350      C PLOT MEAN ENERGIES
351      C
352      DO 665 IT=1,NR
353      ZZ(IT,1)=DEXC(IT)*CXCI
354      ZZ(IT,2)=ADEXC(IT)*CXCI
355      ZZ(IT,3)=DETA(IT)*CXCI
356      ZZ(IT,4)=ADETA(IT)*CXCI
357      665 CONTINUE
358      CALL LPLLOT(ZZ,T,NR+2,4,3.0,3.0,TI4,5)
359      C
360      C PLOT TOTALS
361      C
362      DO 830 IT=1,NR
363      ZZ(IT,1)=EKET(IT)
364      ZZ(IT,2)=EKETF(IT)
365      ZZ(IT,3)=APET(IT)
366      ZZ(IT,4)=APETF(IT)
367      ZZ(IT,5)=TOTT(IT)
368      ZZ(IT,6)=TOTTF(IT)
369      830 CONTINUE
370      CALL LPLLOT(ZZ,T,NR+2,6,3.0,3.0,TI1,3)
371      DO 840 IT=1,NR
372      ZZ(IT,1)=PET(IT)
373      ZZ(IT,2)=PETF(IT)
374      840 CONTINUE
375      CALL LPLLOT(ZZ,T,NR+2,2,3.0,3.0,TI2,4)
376      CALL PLOT(O.O,O.O,999)
377      STOP
378      END
379      SUBROUTINE FINT(F,NI,NJ,H,SUM,IL,IR,JT,JB)
380      IMPLICIT REAL*8 (A-H,O-Z)
381      DIMENSION F(NI,NJ)
382      C
383      C PERFORMS INTEGRATION OF F OVER THE SQUARE
384      C IL<=X<=IR ; JB<=Y<=JT
385      C ASSUMING A UNIFORM MESH LENGTH H IN BOTH DIRECTIONS.
386      C
387      SUM=O.O
388      C
389      C DO EDGES.
390      C
391      IL1=IL+1
392      IR1=IR-1
393      DO 300 I=IL1,IR1
394      300 SUM=SUM+F(I,JT)

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395      DO 310 I=IL1,IR1
396 310 SUM=SUM+F(I,JB)
397      JT1=JT-1
398      JB1=JB+1
399      DO 320 J=JB1,JT1
400 320 SUM=SUM+F(IL,J)+F(IR,J)
401      SUM=SUM*.5ODO
402 C
403 C DO CORNERS.
404 C
405      SUM1=F(IL,JT)+F(IR,JT)+F(IL,JB)+F(IR,JB)
406      SUM=SUM+.25DO*SUM1
407 C
408 C DO INTERIOR.
409 C
410      SUM1=0.0
411      DO 350 J=JB1,JT1
412      DO 350 I=IL1,IR1
413 350 SUM1=SUM1+F(I,J)
414      SUM=SUM1+SUM
415      SUM=SUM*H*H
416      RETURN
417      ENO
418      SUBROUTINE EKINEN(XCSP,TASP,NI,NJ,H,EKE,IL1,IR1,
419 1 JT1,JB1,EKET,IFLAG)
420      IMPLICIT REAL*8 (A-H,O-Z)
421      DIMENSION XCSP(63,27),TASP(63,27)
422      COMMON / HXMESH / HX(63)
423      COMMON / HYMESH / HY(27)
424      COMMON / WKS101 / WK1(63)
425      COMMON / WKS1D2 / WK2(63)
426      COMMON / WKS1D3 / WK3(63)
427      COMMON / WKS1D4 / WK4(63)
428      COMMON / WKS1D5 / WK5(63)
429      COMMON / WKS106 / WK6(63)
430      COMMON / WKS1D7 / WK7(63)
431      COMMON / WKS108 / WK8(63)
432      COMMON / WKS109 / WK9(63)
433      COMMON / WKS110 / WK10(63)
434      COMMON / WORKA / WKA(63,27)
435      COMMON / WORKB / WKB(63,27)
436      COMMON / WORKC / WKC(63,27)
437      COMMON / WORKD / WKD(63,27)
438 C
439 C FIND KINETIC ENERGY.
440 C
441      CALL DXDYDS(WKA,XCSP,1.ODO,.TRUE...FALSE.,NI,NJ)
442      CALL PSOLVE(WKB,WKA,NI,NJ,.TRUE...FALSE...FALSE.)
443      CALL NLLOOP(WKA,WKB,WKB,NI,NJ)
444      CALL PSOLVE(WKO,WKA,NI,NJ,.TRUE...TRUE...FALSE.)
445      CALL DXDYDS(WKA,XCSP,1.ODO,.FALSE...TRUE.,NI,NJ)
446      CALL PSOLVE(WKB,WKA,NI,NJ,.FALSE...TRUE...FALSE.)
447      CALL NLLOOP(WKA,WKB,WKB,NI,NJ)
448      CALL PSOLVE(WKC,WKA,NI,NJ,.TRUE...TRUE...FALSE.)
449      CALL GDAOGD(WKD,1.ODO,WKD,1.ODO,WKC,NI,NJ)
450      CALL DXDYDS(WKA,TASP,1.ODO,.TRUE...FALSE.,NI,NJ)
451      CALL PSOLVE(WKB,WKA,NI,NJ,.TRUE...FALSE...FALSE.)
452      CALL NLLOOP(WKA,WKB,WKB,NI,NJ)
453      CALL PSOLVE(WKC,WKA,NI,NJ,.TRUE...TRUE...FALSE.)
454      CALL GOADGD(WKD,1.ODO,WKD,1.ODO,WKC,NI,NJ)
455      CALL OXDYDS(WKA,TASP,1.ODO,.FALSE...TRUE.,NI,NJ)
456      CALL PSOLVE(WKB,WKA,NI,NJ,.FALSE...TRUE...FALSE.)
457      CALL NLLOOP(WKA,WKB,WKB,NI,NJ)
458      CALL PSOLVE(WKC,WKA,NI,NJ,.TRUE...TRUE...FALSE.)
459      CALL GOADGO(WKD,1.ODO,WKO,1.ODO,WKC,NI,NJ)
460      CALL FINT(WKO,NI,NJ,H,EKE,IL1,IR1,JT1,JB1)

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461      IF(IFLAG.EQ.1) CALL TINT(WKO,EKET,NI,NJ)
462      RETURN
463      ENO
464      SUBROUTINE POTENS(XCSP,TASP,NI,NJ,H,PE,IL1,IR1,
465 1 JT1,JB1,CPE,PET,IFLAG)
466      IMPLICIT REAL*8 (A-H,O-Z)
467      DIMENSION XCSP(NI,NJ),TASP(NI,NJ)
468      COMMON / HXMESH / HX(63)
469      COMMON / HYMESH / HY(27)
470      COMMON / WKS1D1 / WK1(63)
471      COMMON / WKS1D2 / WK2(63)
472      COMMON / WKS1D3 / WK3(63)
473      COMMON / WKS1D4 / WK4(63)
474      COMMON / WKS1D5 / WK5(63)
475      COMMON / WKS1D6 / WK6(63)
476      COMMON / WKS1D7 / WK7(63)
477      COMMON / WKS1D8 / WK8(63)
478      COMMON / WKS1D9 / WK9(63)
479      COMMON / WKS1D10 / WK10(63)
480      COMMON / WORKA / WKA(63,27)
481      COMMON / WORKB / WKB(63,27)
482      COMMON / WORKC / WKC(63,27)
483      COMMON / WDRKD / WKD(63,27)
484      C
485      C FORM LAPLACE DF XCI
486      C
487      CALL D2XYS(WKA,XCSP,1.ODO,.TRUE.,.FALSE.,NI,NJ)
488      CALL PSDLVE(WKB,WKA,NI,NJ,.TRUE.,.FALSE.,.TRUE.)
489      CALL D2XYS(WKA,XCSP,1.ODO,.FALSE.,.TRUE.,NI,NJ)
490      CALL PSDLVE(WKC,WKA,NI,NJ,.FALSE.,.TRUE.,.TRUE.)
491      CALL GOAOGO(WKO,1.000,WKB,1.000,WKC,NI,NJ)
492      C
493      C FDRM LAPLACE DF TAU
494      C
495      CALL D2XYS(WKA,TASP,1.ODO,.TRUE.,.FALSE.,NI,NJ)
496      CALL PSOLVE(WKB,WKA,NI,NJ,.TRUE.,.FALSE.,.TRUE.)
497      CALL D2XYS(WKA,TASP,1.000,.FALSE.,.TRUE.,NI,NJ)
498      CALL PSOLVE(WKC,WKA,NI,NJ,.FALSE.,.TRUE.,.TRUE.)
499      CALL GOADGO(WKC,1.000,WKC,1.000,WKB,NI,NJ)
500      C
501      C DD SUMS.
502      C
503      CALL GDAOGD(WKA,1.ODO,WKD,1.ODO,WKC,NI,NJ)
504      CALL GDADGO(WKA,1.ODO,WKA,-CPE,TASP,NI,NJ)
505      CALL GDADGO(WKB,1.ODO,WKD,-1.ODO,WKC,NI,NJ)
506      CALL GDAOGO(WKB,1.000,WKB,CPE,TASP,NI,NJ)
507      C
508      C FORM SQUARES.
509      C
510      CALL NLLDOP(WKC,WKA,WKA,NI,NJ)
511      CALL PSDLVE(WKA,WKC,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
512      CALL NLLDDP(WKC,WKB,WKB,NI,NJ)
513      CALL PSDLVE(WKB,WKC,NI,NJ,.TRUE.,.TRUE.,.FALSE.)
514      C
515      C FDRM SUM.
516      C
517      CALL GOAOGO(WKC,1.ODO,WKA,1.ODO,WKB,NI,NJ)
518      C
519      C INTEGRATE.
520      C
521      CALL FINT(WKC,NI,NJ,H,PE,IL1,IR1,JT1,JB1)
522      IF(IFLAG.EQ.1) CALL TINT(WKC,PET,NI,NJ)
523      RETURN
524      ENO
525      SUBROUTINE MEAN(FIE,NI,NJ,AVE,IL,IR,JT,JB)
526      C

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527 C FIND ARITHMETIC MEAN OF FIELD.
528 C
529     IMPLICIT REAL*8 (A-H,O-Z)
530     DIMENSION FIE(NI,NJ)
531     SUM=0.0DO
532     ISUM=0
533     DO 100 J=JB,JT
534     DO 100 I=IL,IR
535     ISUM=ISUM+1
536 100 SUM=SUM+FIE(I,J)
537     AVE=SUM/DFLOAT(ISUM)
538     RETURN
539     END
540     SUBROUTINE AMEAN(FIE,NI,NJ,AVE,IL,IR,JT,JB)
541 C
542 C FINDS ARITHMETIC MEAN OF ABSOLUTE VALUE OF FIELD
543 C
544     IMPLICIT REAL*8 (A-H,O-Z)
545     DIMENSION FIE(NI,NJ)
546     SUM=0.000
547     ISUM=0
548     DO 100 J=JB,JT
549     DO 100 I=IL,IR
550     ISUM=ISUM+1
551 100 SUM=SUM+DABS(FIE(I,J))
552     AVE=SUM/DFLOAT(ISUM)
553     RETURN
554     END
555     SUBROUTINE SUM(FIE,NI,NJ,SUM1,IL,IR,JT,JB)
556     IMPLICIT REAL*8 (A-H,O-Z)
557     DIMENSION FIE(NI,NJ)
558     SUM1=0.0DO
559     DO 100 J=JB,JT
560     DO 100 I=IL,IR
561 100 SUM1=SUM1+OABS(FIE(I,J))
562     RETURN
563     END
564     SUBROUTINE SUMB(FIE1,FIE2,NI,NJ,SUM1,IL,IR,JT,JB)
565     IMPLICIT REAL*8 (A-H,O-Z)
566     DIMENSION FIE1(NI,NJ),FIE2(NI,NJ)
567     SUM1=0.000
568     DO 100 J=JB,JT
569     DO 100 I=IL,IR
570     F1=OABS(FIE1(I,J))
571     F2=OABS(FIE2(I,J))
572 100 SUM1=SUM1+OMAX1(F1,F2)
573     RETURN
574     END
575     SUBROUTINE LPLLOT(PL,T,NT,IFLAG,XSIZE,YSIZE,TITLE,NTITLE)
576     DIMENSION PL(NT,6),T(NT),SCA(4)
577     REAL TITLE(NTITLE)
578     NUM=4*NTITLE
579     NT2=NT-2
580     NT1=NT-1
581 C
582 C SCALE HORIZONTAL AXIS
583 C
584     CALL SCALE(T,XSIZE,NT2,1)
585 C
586 C SCALE VERTICAL AXIS
587 C
588     XZ1=XSIZE+.8
589     YZ1=YSIZE+.8
590     CALL ORIGIN(999,XZ1,YZ1,1.0,1.0)
591     SC=0.0
592     OR=1.0E55

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593      DD 100 I=1,IFLAG
594      DO 100 J=1,NT2
595      SC=AMAX1(SC,PL(J,I))
596      DR=AMIN1(OR,PL(J,I))
597 100 CONTINUE
598      SCA(1)=OR
599      SCA(2)=SC
600      CALL SCALE(SCA,YSIZE,2,1)
601      DD 125 I=1,IFLAG
602      PL(NT,I)=SCA(4)
603 125 PL(NT1,I)=SCA(3)
604 C
605 C PLOT VERTICAL AXIS
606 C
607      CALL AX2EP(1.0,3,2,0,.85)
608      CALL AXIS2(0.0,0.0,TITLE,NUM,YSIZE,90.0,SCA(3),SCA(4),-1.0)
609      CALL AXIS2(XSIZE,0.0,' ',1,-YSIZE,90.0,SCA(3),SCA(4),1.0)
610 C
611 C PLOT HORIZONTAL AXIS
612 C
613      CALL AX2EP(1.0,3,1,0,.85)
614      CALL AXIS2(0.0,0.0,'TIME (HOURS)',-12,XSIZE,0.0,T(NT1),T(NT),1.)
615      CALL AXIS2(0.0,YSIZE,' ',-1,-XSIZE,0.0,T(NT1),T(NT),1.)
616 C
617 C PLDT CURVES
618 C
619      DO 150 I=1,IFLAG
620      CALL LINE(T,PL(1,I),NT2,1,1,I)
621 150 CDNTINUE
622      RETURN
623      END
624 SUBROUTINE APDTEN(APE,IL1,IR1,JT1,JB1,F,NI,NJ,
625 1 H,APET,IFLAG)
626 IMPLICIT REAL*8 (A-H,O-Z)
627 DIMENSION F(NI,NJ)
628 COMMON / HXMESH / HX(63)
629 CDMMDN / HYMesh / HY(27)
630 CDMMDN / WDRKA / WKA(63,27)
631 COMMON / WDRKB / WKB(63,27)
632 CALL NLLDOP(WKA,F,F,NI,NJ)
633 CALL PSDLVE(WKB,WKA,NI,NJ,.TRUE...TRUE...FALSE.)
634 CALL FINT(WKB,NI,NJ,H,APE,IL1,IR1,JT1,JB1)
635 IF(IFLAG.EQ.1) CALL TINT(WKB,APET,NI,NJ)
636 RETURN
637 END
638 SUBROUTINE ZEENE(F1,F2,NI,NJ,JT,JB,IL,IR,H,ZKE,EKE,ZAPE,EAPE)
639 C
640 C SUBROUTINE TO CALCULATE ZDNAL AND EDDY ENERGIES.
641 C
642 C F1 = XCI
643 C F2 = TAU
644 IMPLICIT REAL*8 (A-H,O-Z)
645 DIMENSION F1(NI,NJ),F2(NI,NJ)
646 COMMON / HXMESH / HX(63)
647 COMMON / HYMesh / HY(27)
648 COMMON / WKS1D1 / WK1(63)
649 COMMON / WKS1D2 / WK2(63)
650 COMMON / WKS1D3 / WK3(63)
651 COMMON / WKS1D4 / WK4(63)
652 COMMON / WKS1D5 / WK5(63)
653 COMMON / WKS1D6 / WK6(63)
654 COMMON / WKS1D7 / WK7(63)
655 COMMON / WKS1D8 / WK8(63)
656 CDMMDN / WKS1D9 / WK9(63)
657 COMMON / WKS110 / WK10(63)
658 COMMON / WORKA / WKA(63,27)

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659      COMMON / WORKB / WKB(63,27)
660      COMMON / WORKC / WKC(63,27)
661      COMMON / WORKD / WKD(63,27)
662      COMMON / WORKE / WKE(63,27)
663      COMMON / WORKF / WKF(63,27)
664      C
665      C FORM AVERAGE IN Y-DIRECTION.
666      C
667          ANI=DFLOAT(IR-IL+1)
668          DO 100 J=1,NJ
669              SUM=0.000
670              SUM1=0.000
671              DO 105 I=IL,IR
672                  105 SUM=SUM+F1(I,J)
673                  DO 106 I=IL,IR
674                      106 SUM1=SUM1+F2(I,J)
675                      SUM=SUM/ANI
676                      SUM1=SUM1/ANI
677                      DO 110 I=1,NI
678                          110 WKE(I,J)=SUM
679                          DO 115 I=1,NI
680                              115 WKF(I,J)=SUM1
681                      100 CONTINUE
682      C
683      C CALCULATE ZKE
684      C
685          CALL EKINEN(WKE,WKF,NI,NJ,H,ZKE,IL,IR,
686              1 JT,JB,SUM,O)
687      C
688      C CALCULATE ZAPE
689      C
690          CALL APOTEN(ZAPE,IL,IR,JT,JB,WKF,NI,NJ,
691              1 H,SUM,O)
692          DO 140 J=1,NJ
693          DO 140 I=1,NI
694              WKE(I,J)=F1(I,J)-WKE(I,J)
695              WKF(I,J)=F2(I,J)-WKF(I,J)
696          140 CONTINUE
697      C
698      C CALCULATE EKE
699      C
700          CALL EKINEN(WKE,WKF,NI,NJ,H,EKE,IL,IR,
701              1 JT,JB,SUM,O)
702          CALL APOTEN(EAPE,IL,IR,JT,JB,WKF,NI,NJ,
703              1 H,SUM,O)
704          RETURN
705          ENO
706          SUBROUTINE TINT(F,SUM,NI,NJ)
707              IMPLICIT REAL*8 (A-H,O-Z)
708              DIMENSION F(NI,NJ)
709              COMMON / HXMESH / HX(1)
710              COMMON / HYMESH / HY(1)
711              NJM=NJ-1
712              NIM=NI-1
713              SUM=0.000
714              DO 100 J=2,NJM
715                  J1=J-1
716                  HYB=(HY(J1)+HY(J))/2.000
717                  DO 100 I=2,NIM
718                      I1=I-1
719                      HXB=(HX(I1)+HX(I))/2.000
720                      SUM=SUM+F(I,J)*HYB*HXB
721              100 CONTINUE
722      C
723      C TOP AND BOTTOM
724      C

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725          HYB=HY(1)/2.ODO
726          HYB2=HY(NJM)/2.ODO
727          DO 150 I=2,NIM
728             I1=I-1
729             HXB=(HX(I1)+HX(I))/2.ODO
730             SUM=SUM+F(I,1)*HYB*HXB
731             SUM=SUM+F(I,NJ)*HYB2*HXB
732          150 CONTINUE
733          C
734          C  ENDS
735          C
736          HXB=(HX(1)+HX(NIM))/2.ODO
737          DD 200 J=2,NJM
738             J1=J-1
739             HYB=(HY(J1)+HY(J))/2.ODO
740             SUM=SUM+F(1,J)*HYB*HXB
741          200 CONTINUE
742          C
743          C  CDRNERS
744          C
745          HXB=(HX(1)+HX(NIM))/2.ODO
746          SUM=SUM+F(1,1)*HXB*HY(1)/2.ODO
747          SUM=SUM+F(1,NJ)*HXB*HY(NJM)/2.ODO
748          RETURN
749          END
END OF FILE

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